

Is Titan flat?

Coupling a 3D Monte Carlo radiative transfer model with the Titan PCM

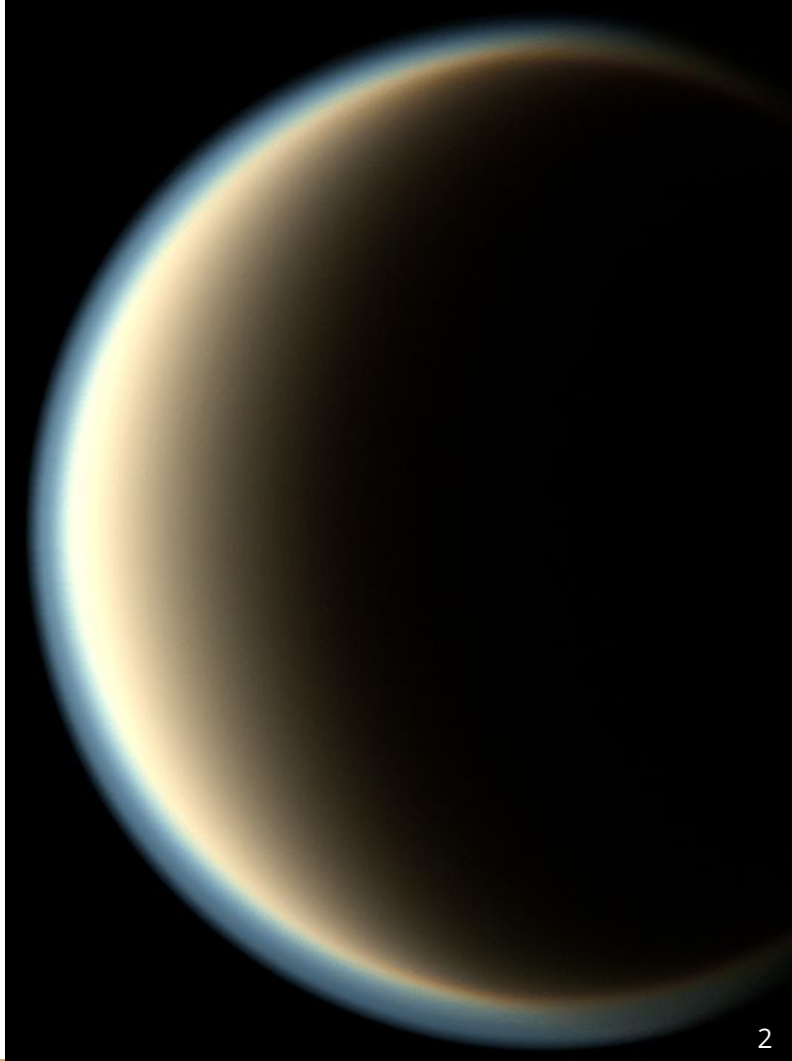
Anthony Arfaux, Sandrine Vinatier, Pascal Rannou, Sébastien Lebonnois, Vincent Eymet, Vincent Forest, Zili He, Ehouarn Millour, Bruno de Batz de Trenquelléon, Clément Petetin

23/06/2026, Journées de la SF2A, Grenoble



Introduction to Titan

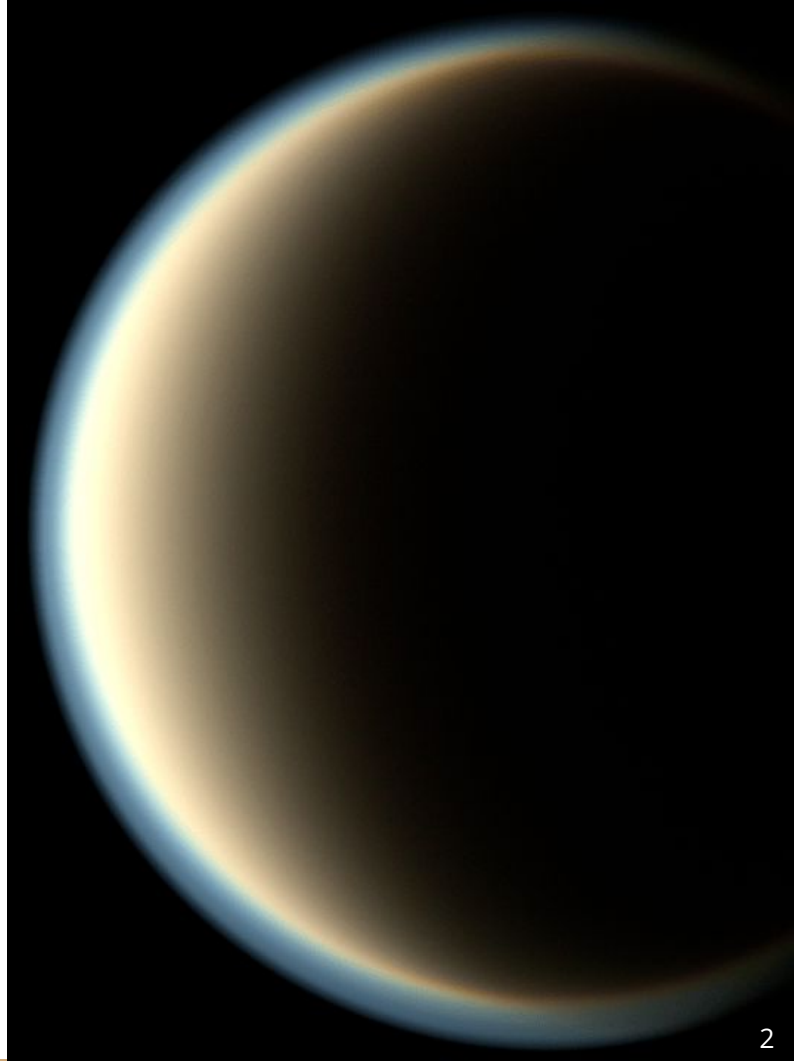
Largest moon of Saturn: 2575 km radius



Introduction to Titan

Largest moon of Saturn: 2575 km radius

Thick atmosphere: ~1000 km



Introduction to Titan

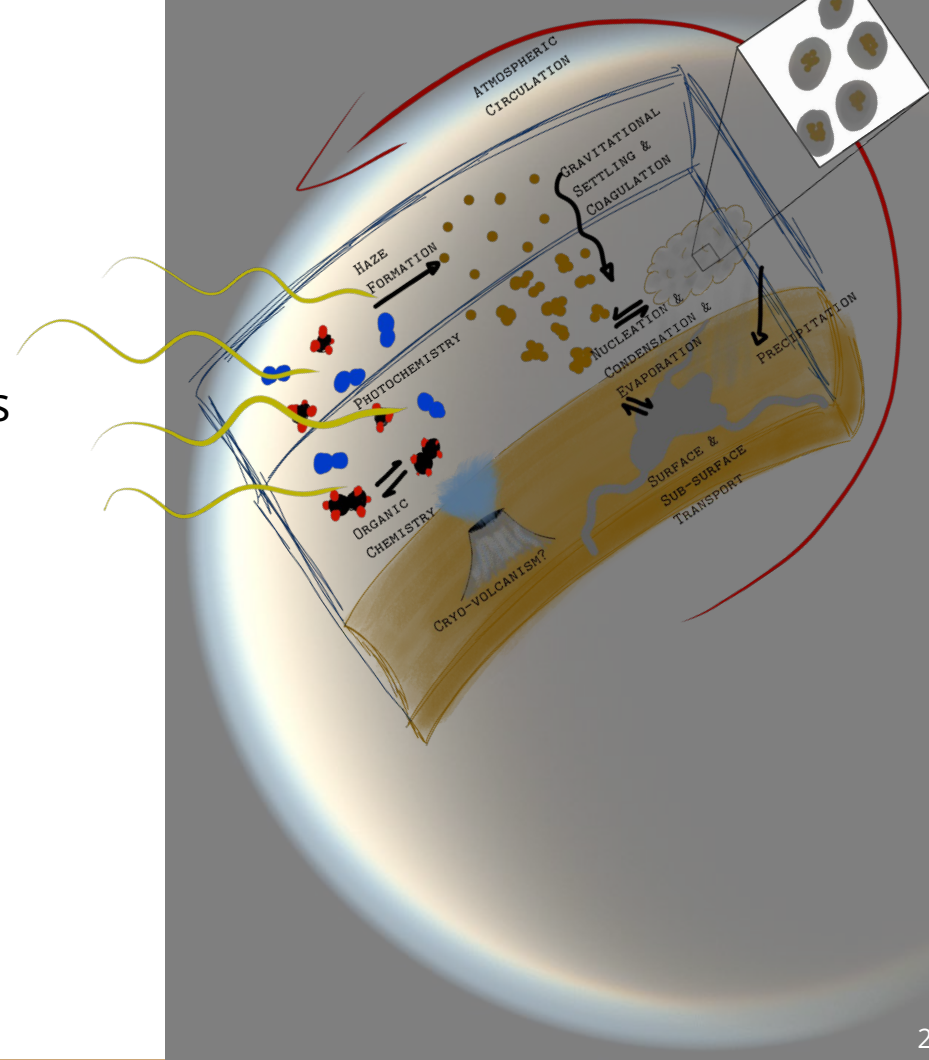
Largest moon of Saturn: 2575 km radius

Thick atmosphere: ~1000 km

Surface pressure: 1.5 bar

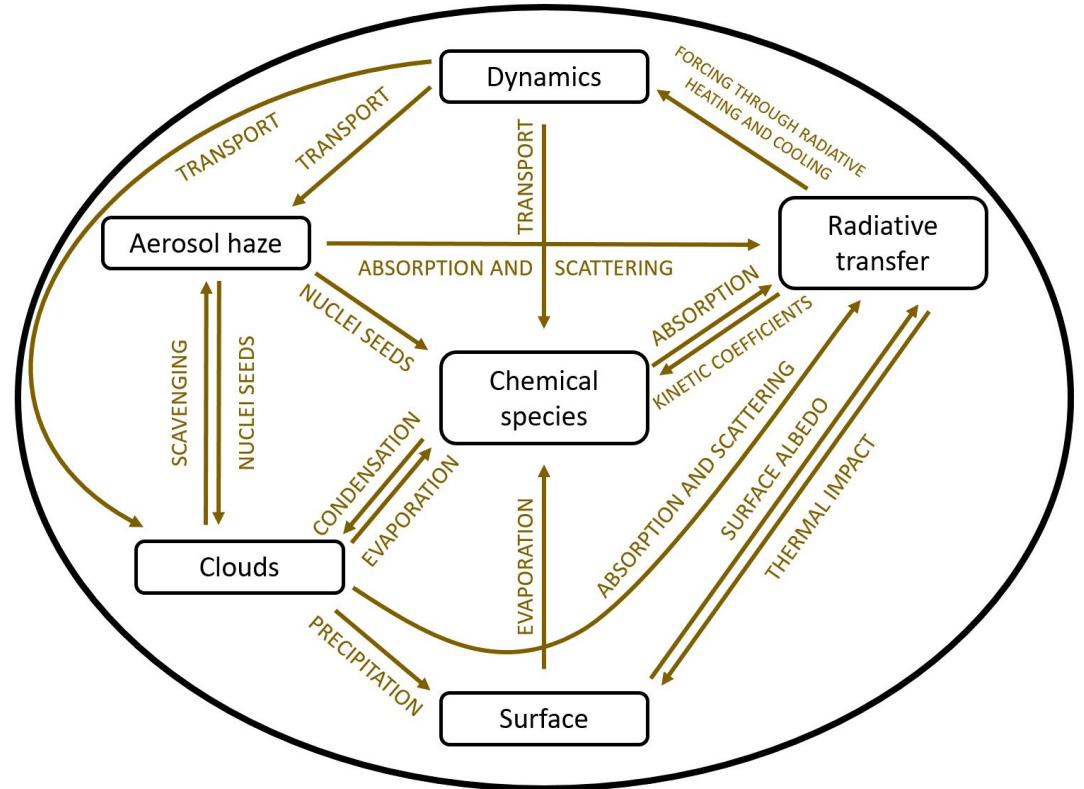
Composition: N_2 (94-98 %), CH_4 (1-5 %)

$CH_4 \Rightarrow$ Haze and clouds



Titan LMDZ Planetary Climate Model

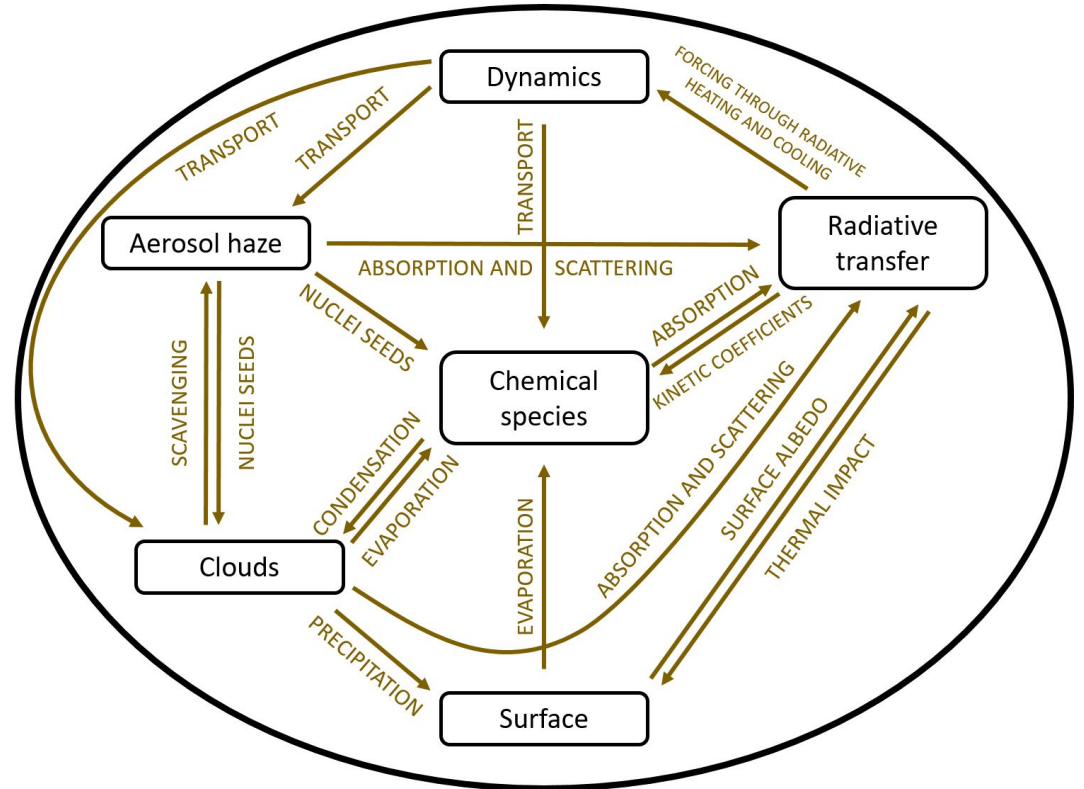
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Titan LMDZ Planetary Climate Model

© Bruno de Batz de Trenquelléon

Highly coupled

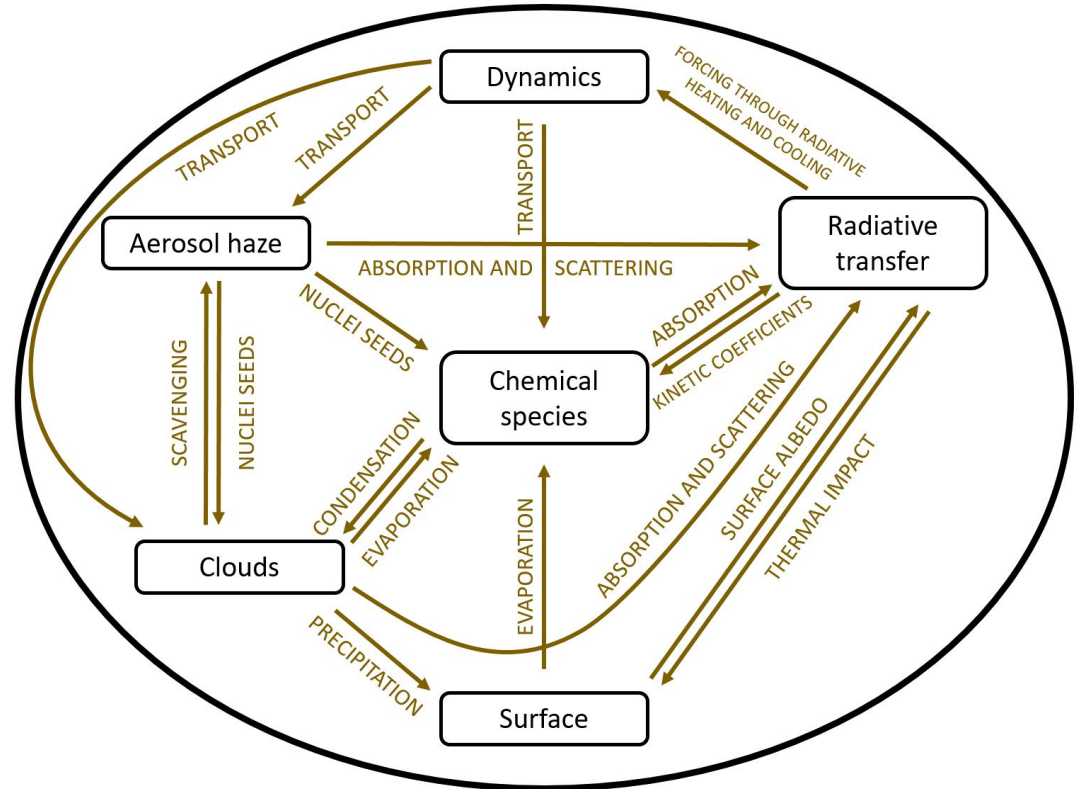


Titan LMDZ Planetary Climate Model

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Highly coupled

Radiative transfer: major impact on whole atmosphere



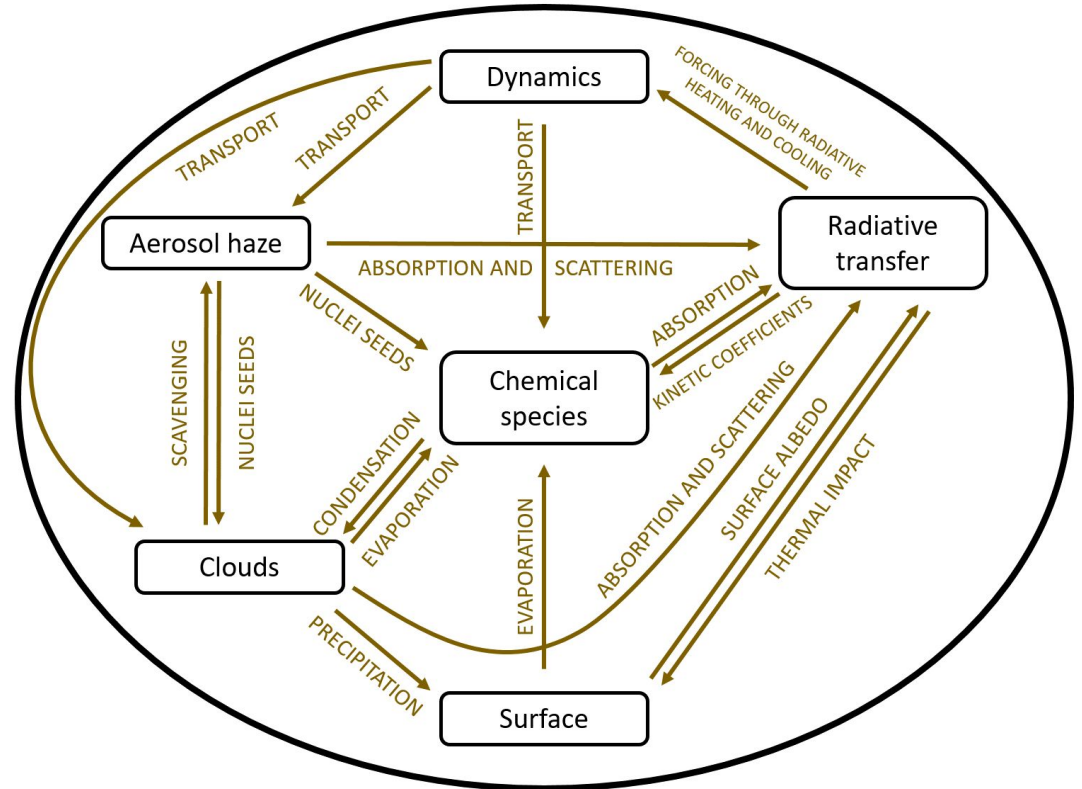
Titan LMDZ Planetary Climate Model

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Highly coupled

Radiative transfer: major impact on whole atmosphere

Radiative transfer: 2-stream plane-parallel (McKay et al. 1989)



htrdr-planets: a nul-collision backward Monte Carlo radiative transfer model

htrdr-planets: model developed by Méso-Star (Toulouse, France)

htrdr-planets: a nul-collision backward Monte Carlo radiative transfer model

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Monte Carlo radiative transfer: ray tracing

htrdr-planets: a nul-collision backward Monte Carlo radiative transfer model

htrdr-planets: model developed by Méso-Star (Toulouse, France)

Monte Carlo radiative transfer: ray tracing

Works in complex geometries: account for sphericity effects

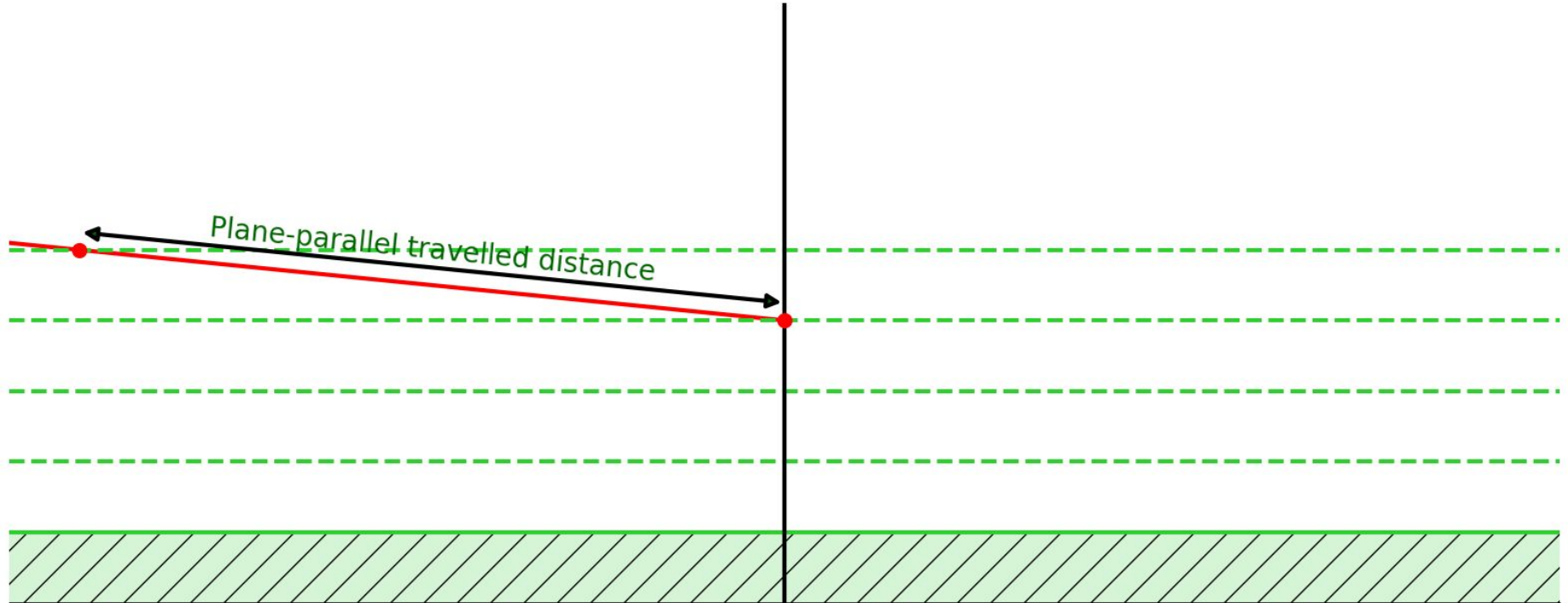
Fully heterogeneous model



Effects of sphericity

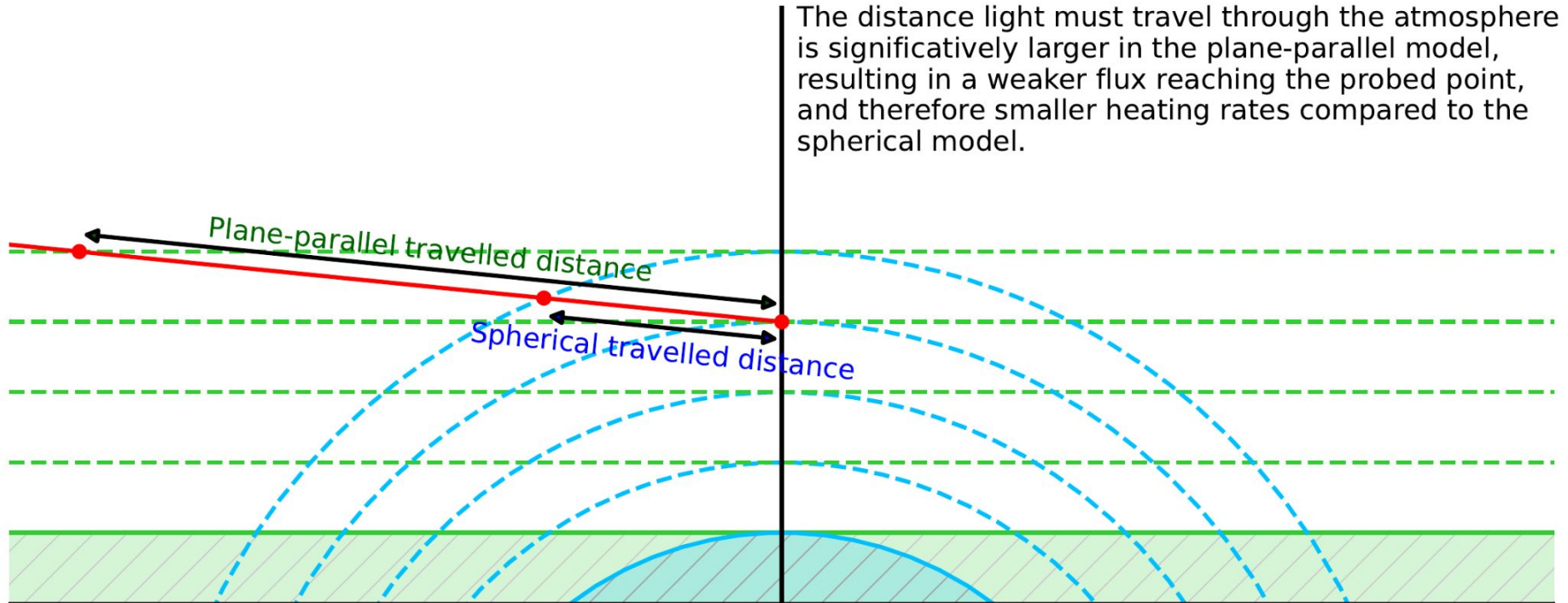
Terminator effects

plane-parallel: long travel distance \Rightarrow strong absorption



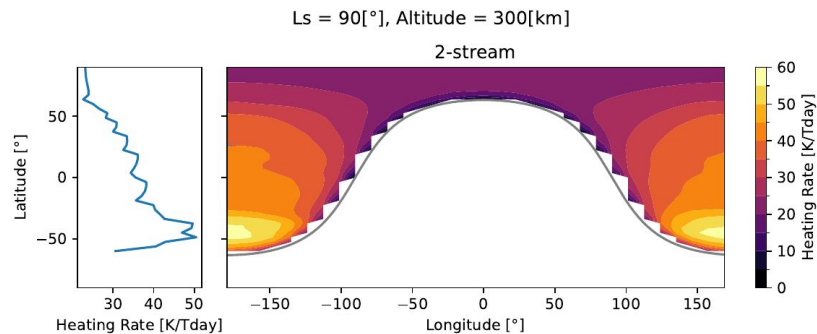
Terminator effects

plane-parallel: long travel distance \Rightarrow strong absorption
spherical: shorter travel distance \Rightarrow weaker absorption

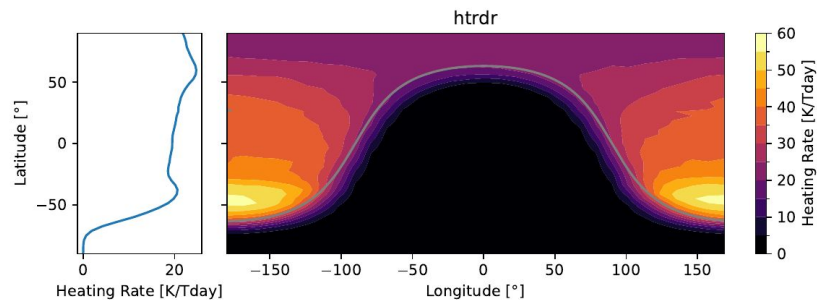


Terminator effects

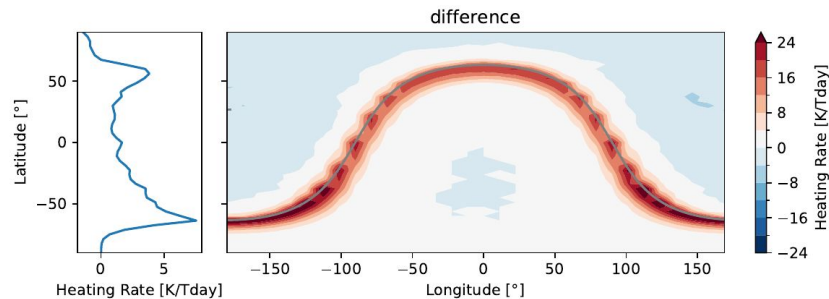
plane-parallel



htrdr
3D



differences



Terminator effects

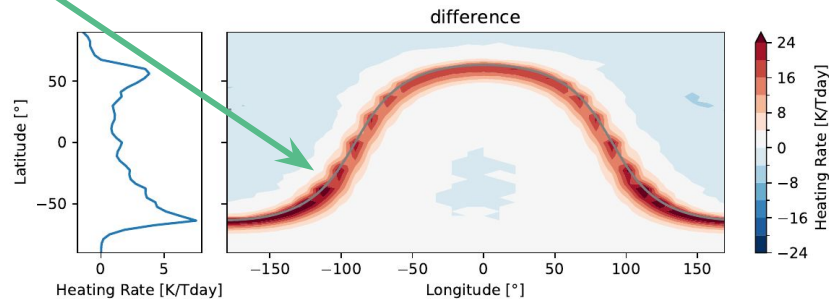
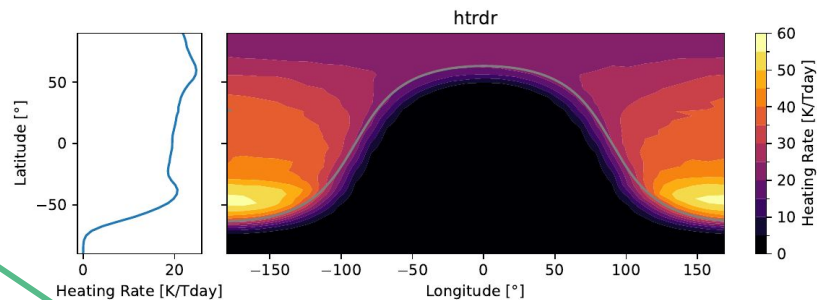
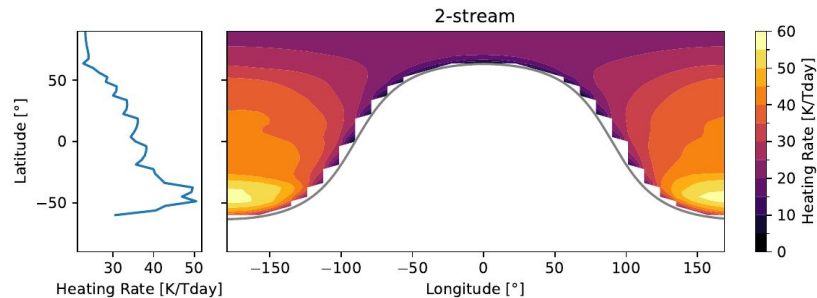
Higher heating rates along the terminator

plane-parallel

htrdr
3D

differences

Ls = 90[°], Altitude = 300[km]



Terminator effects

Higher heating rates along the terminator

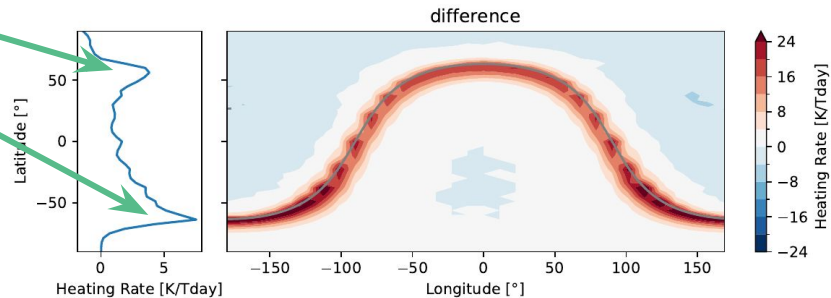
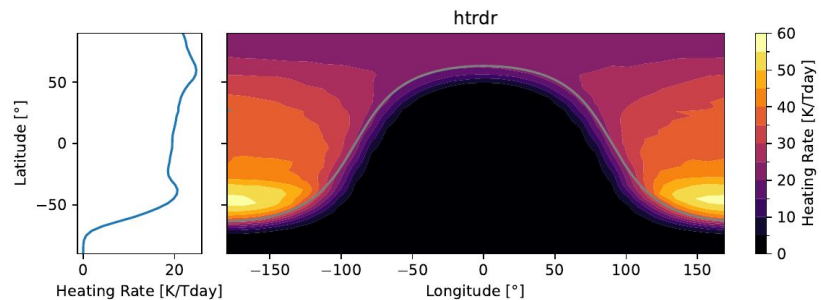
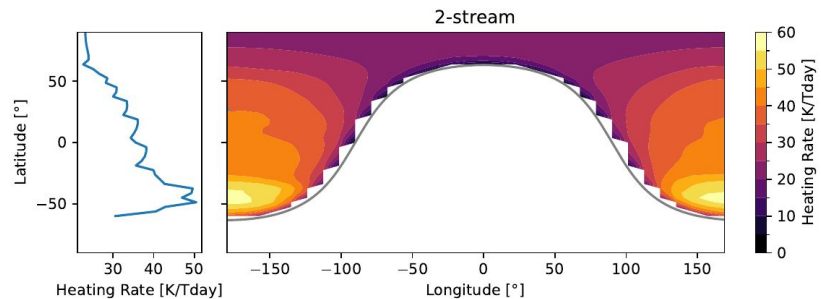
Stronger effects near limits of polar night/day

plane-parallel

htrdr
3D

differences

Ls = 90[°], Altitude = 300[km]



Terminator effects

Higher heating rates along the terminator

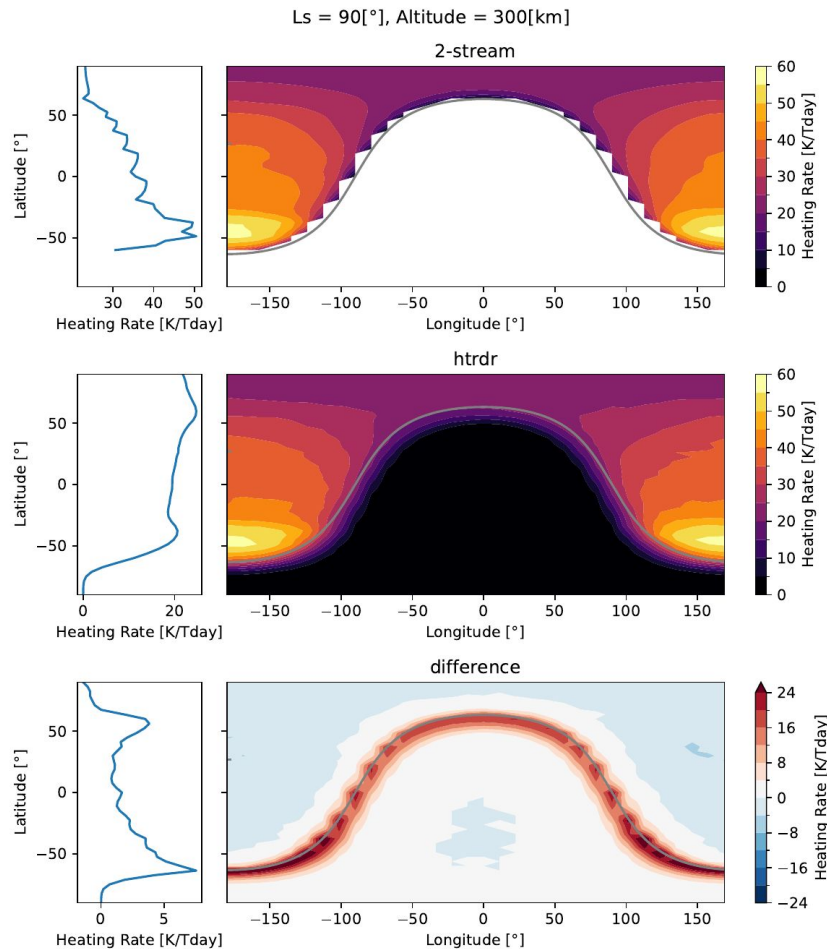
Stronger effects near limits of polar night/day

Zonal averaged differences around 5 K/Titan day

plane-parallel

htrdr
3D

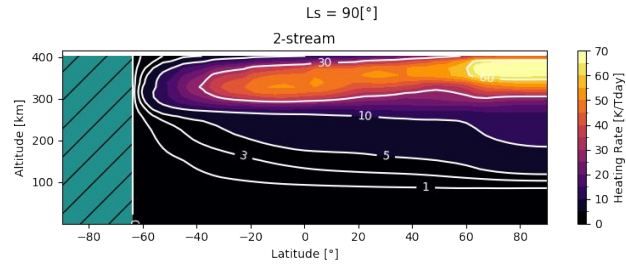
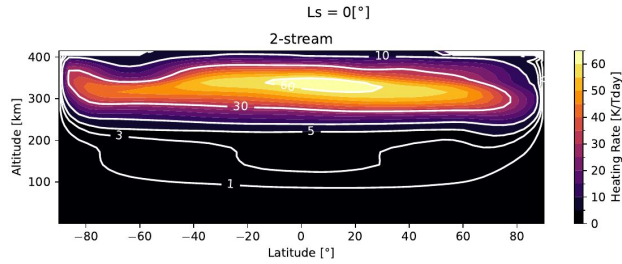
differences



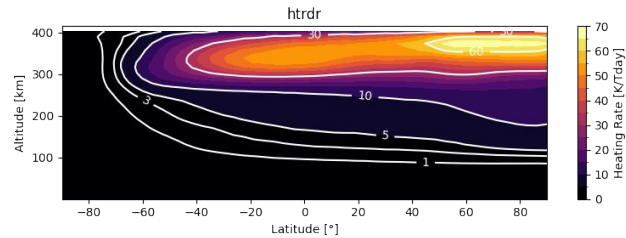
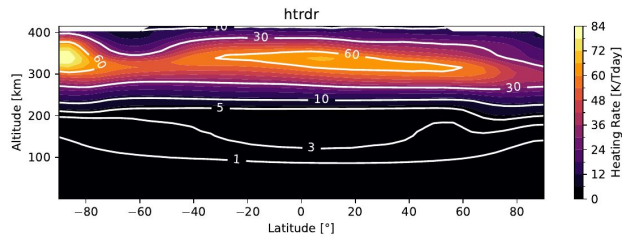
Equinoxe

Solstice

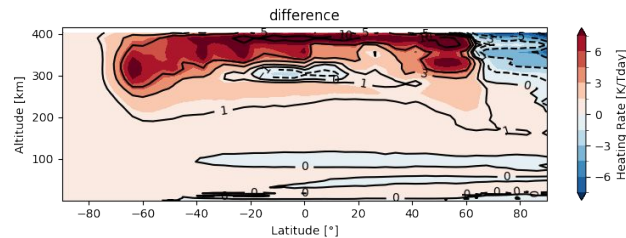
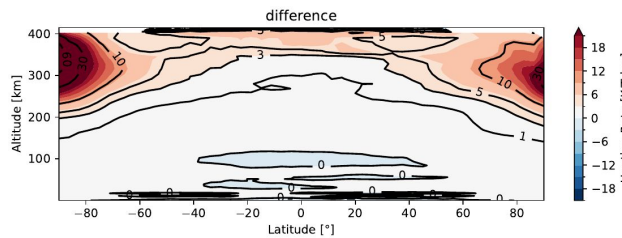
plane-parallel



htrdr
3D



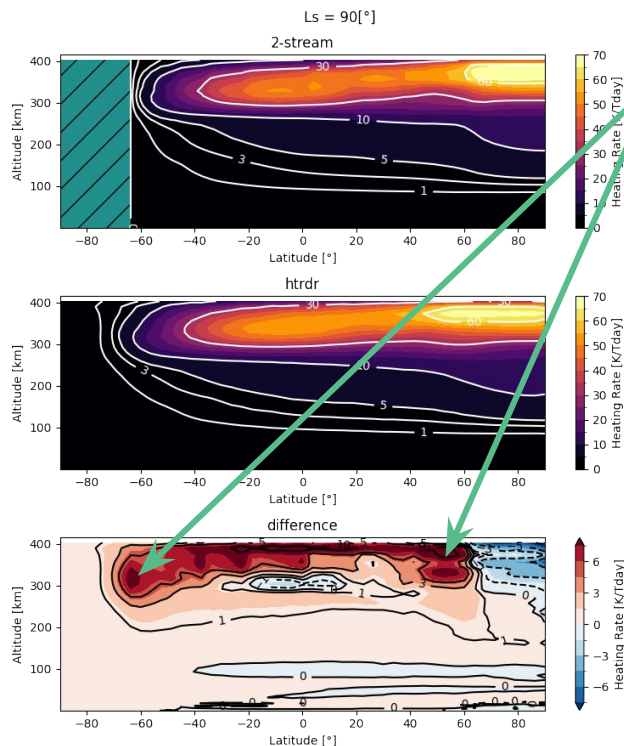
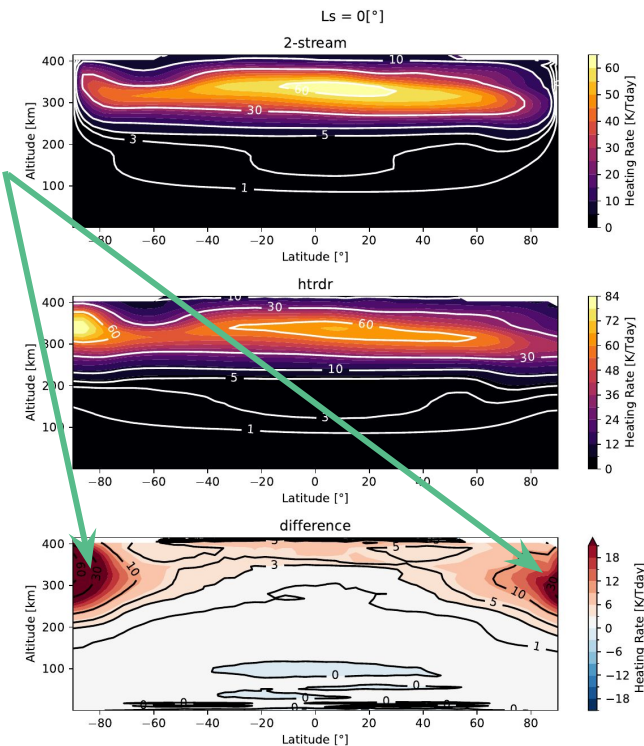
differences



Equinoxes

Solstices

Main differences at the poles



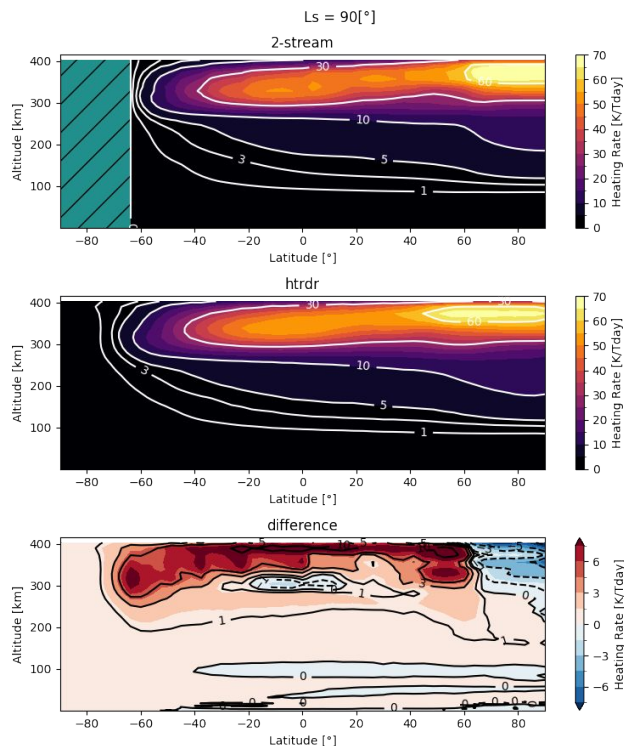
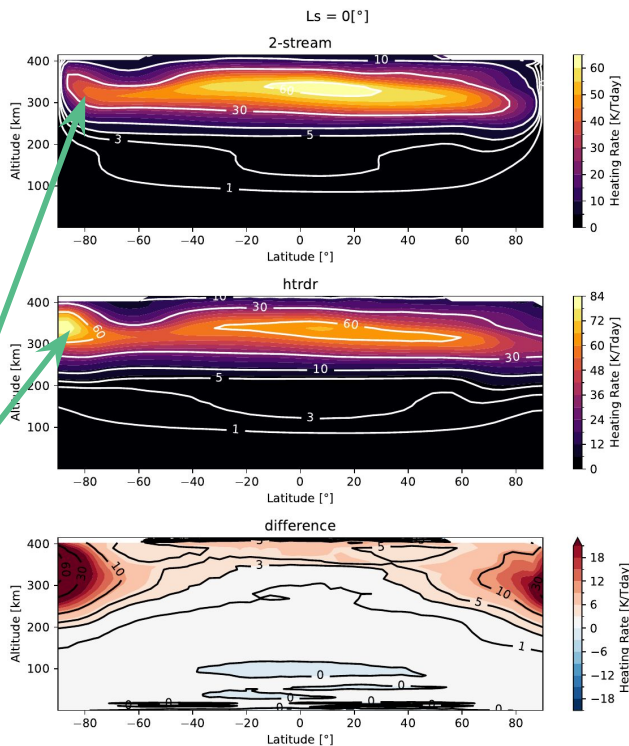
Main differences near 60° latitude: effect of terminator

Equinoxes

Solstices

Main differences at the poles

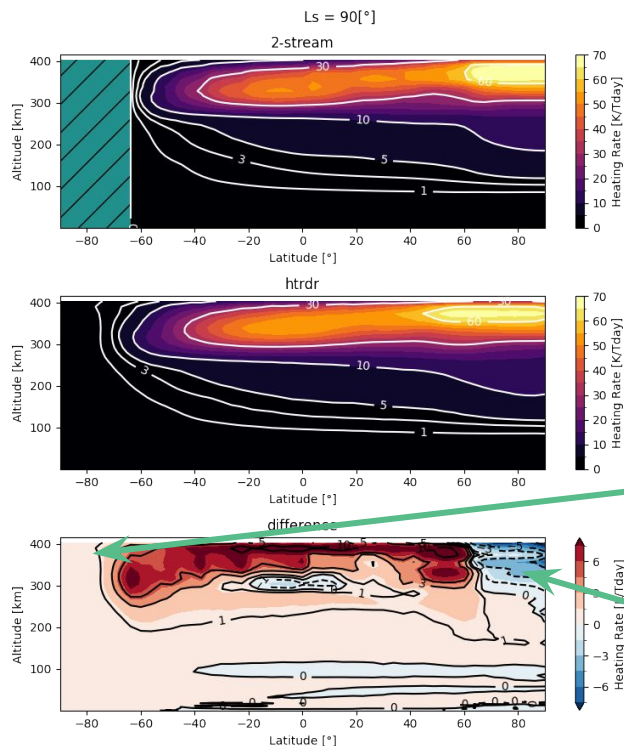
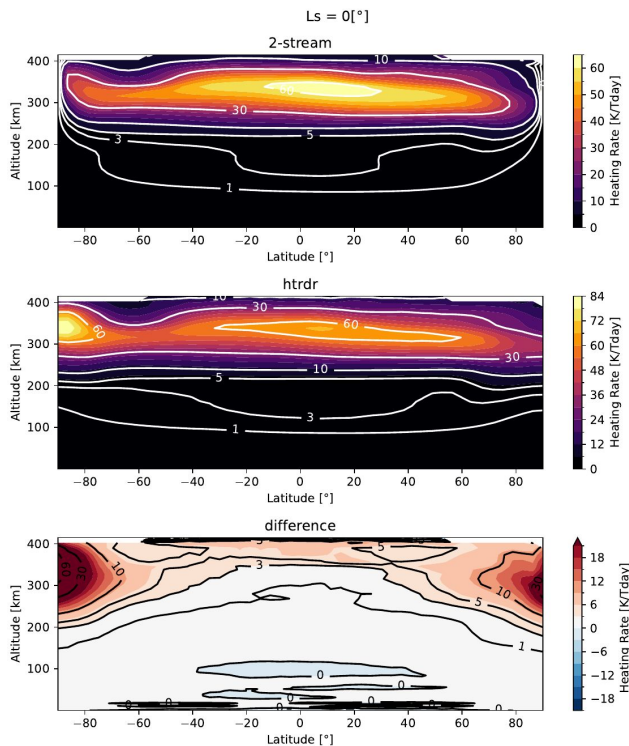
Higher heating rates above former summer poles: detached haze layer at the limb



Main differences near 60° latitude: effect of terminator

Equinoxes

Solstices



Main differences at the poles

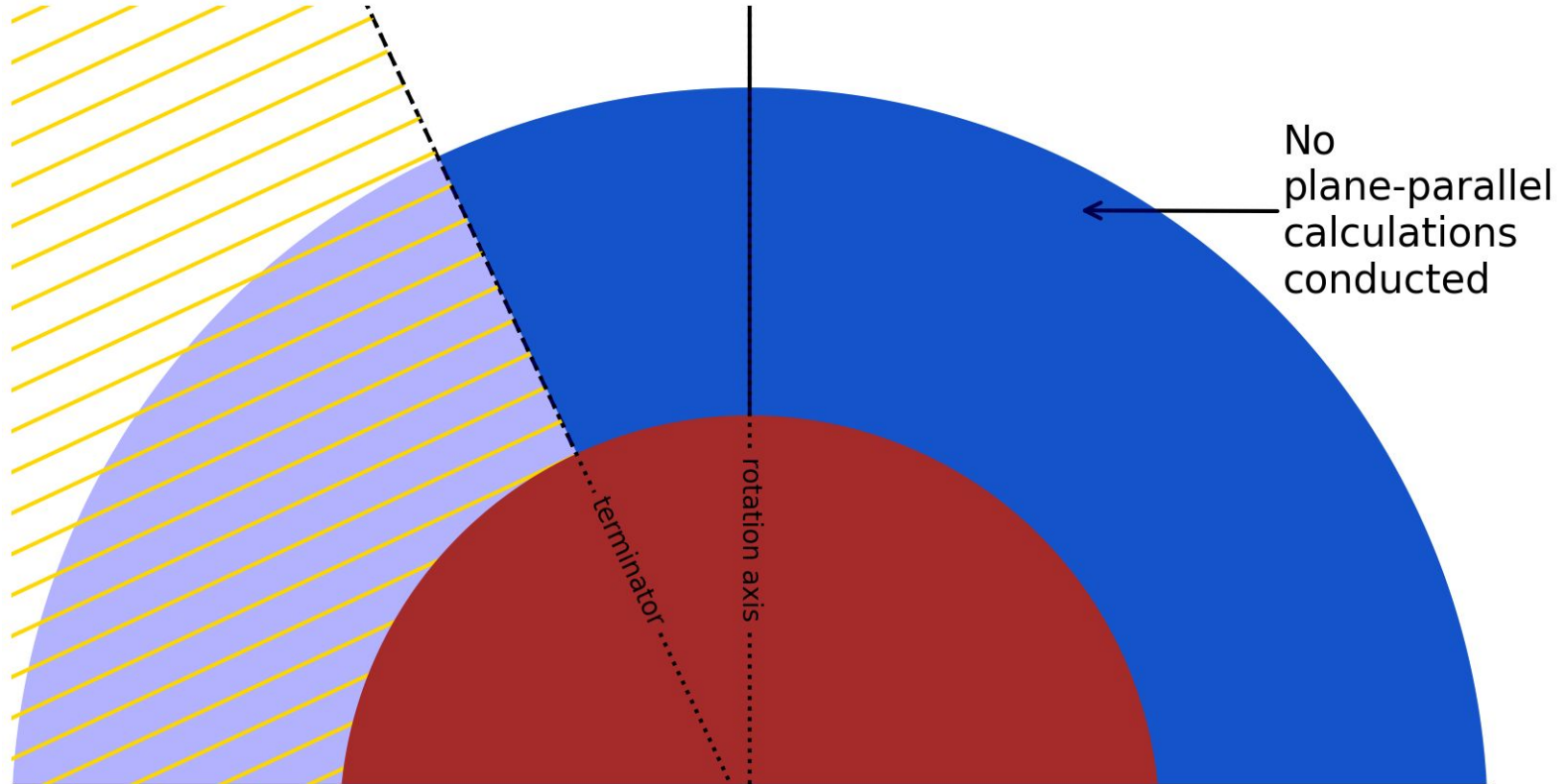
Higher heating rates above former summer poles:
detached haze layer at the limb

Main differences near 60° latitude:
effect of terminator

Hotter polar night, cooler polar day

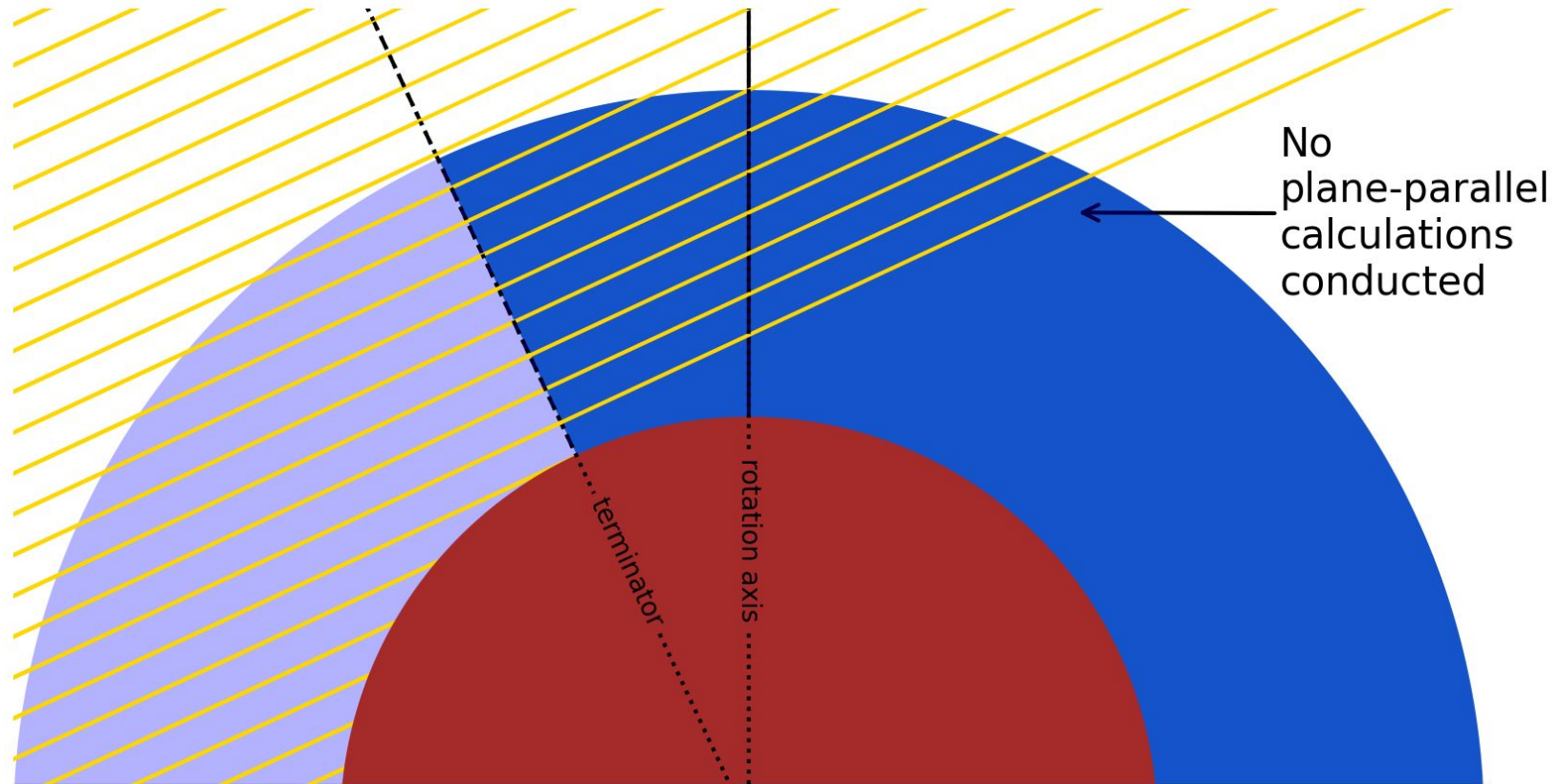
Effects on the polar regions

Plane-parallel



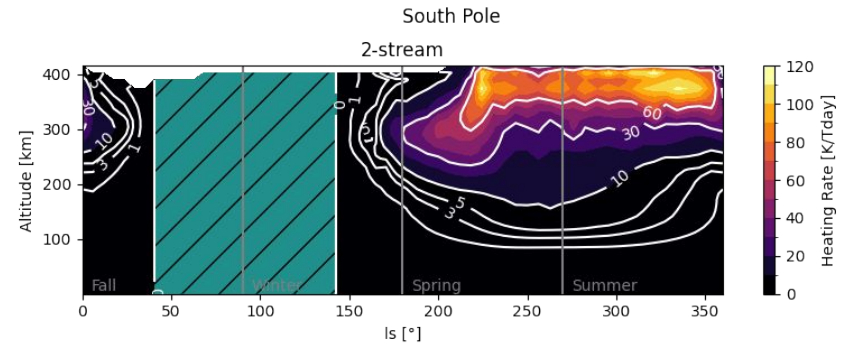
Effects on the polar regions

Spherical model

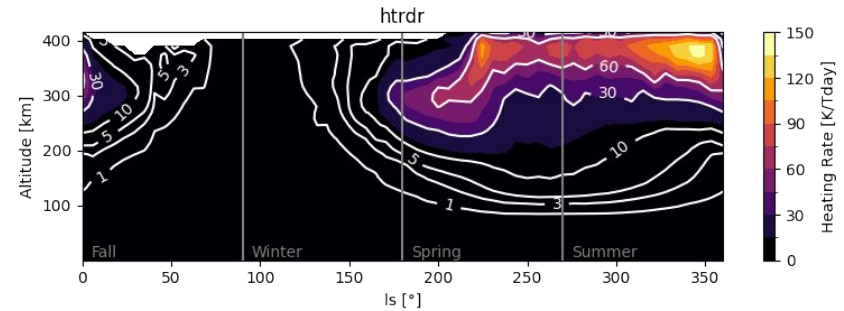


Effects on the polar regions

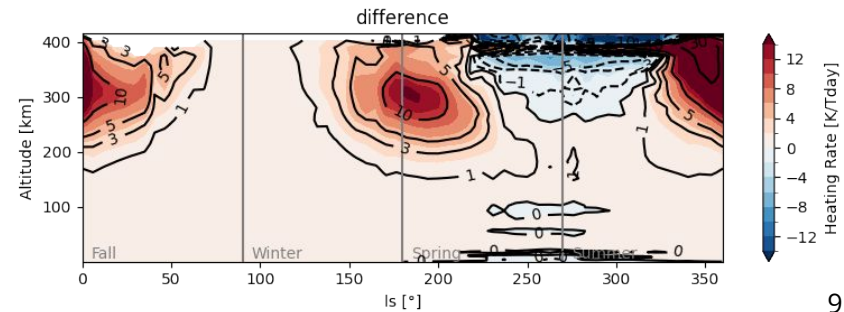
plane-parallel



htrdr
3D



differences



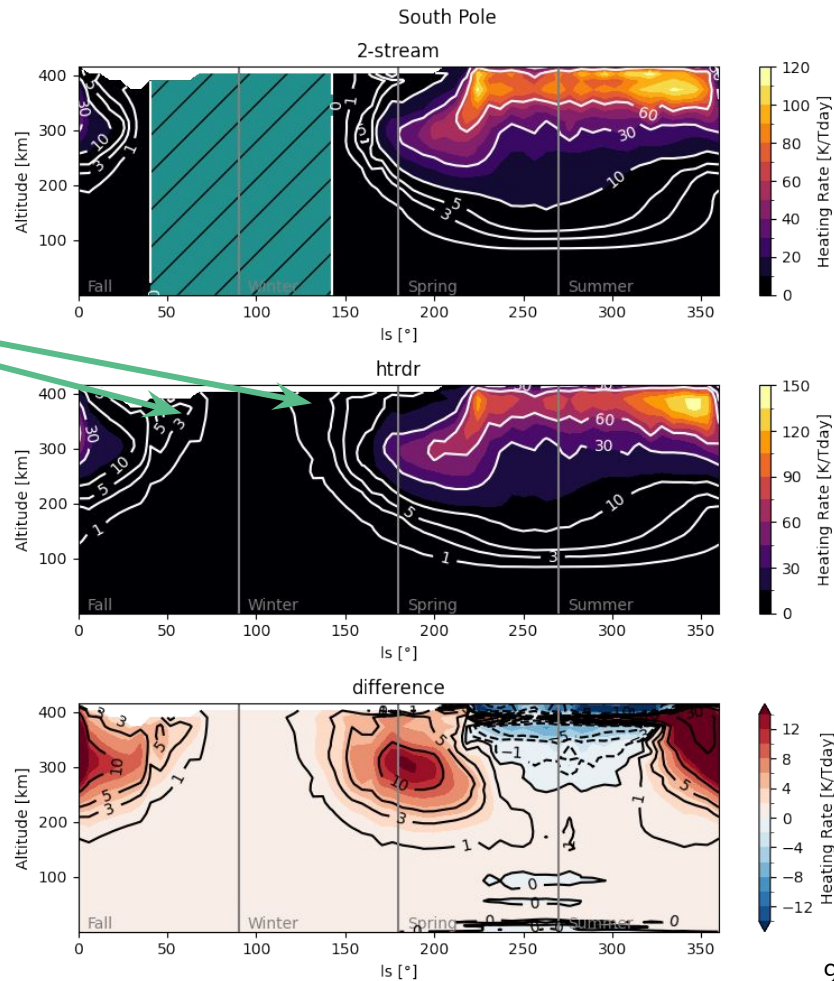
Effects on the polar regions

Heating rates in the polar region during winter

plane-parallel

htrdr
3D

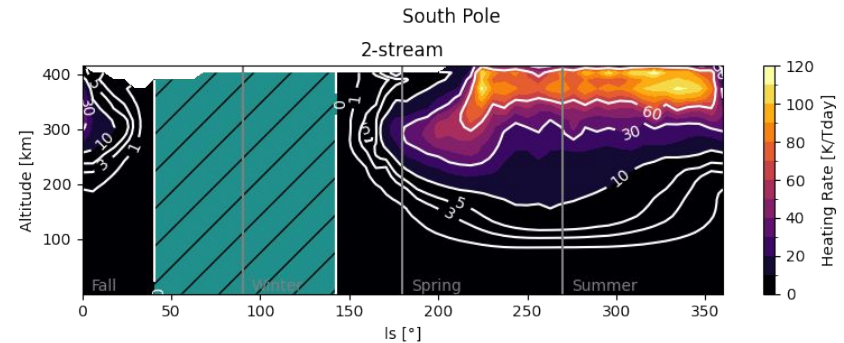
differences



Effects on the polar regions

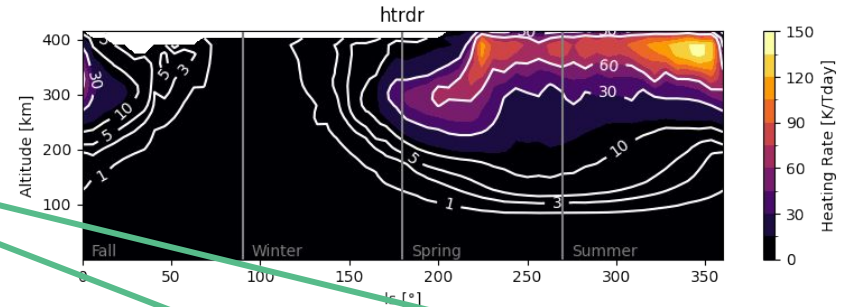
Heating rates in the polar region during winter

plane-parallel

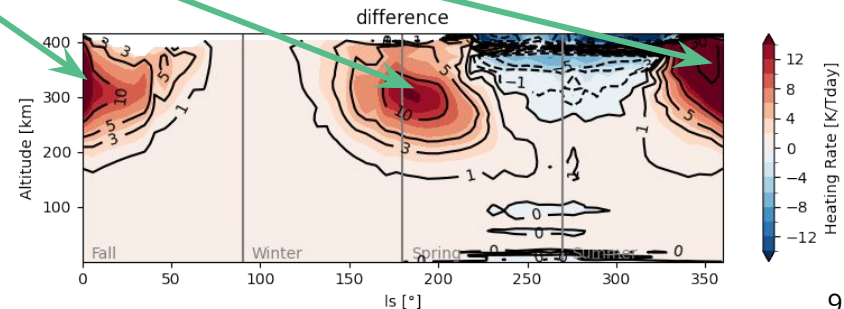


Main differences at equinox

htrdr
3D



differences



Effects on the polar regions

Heating rates in the polar region during winter

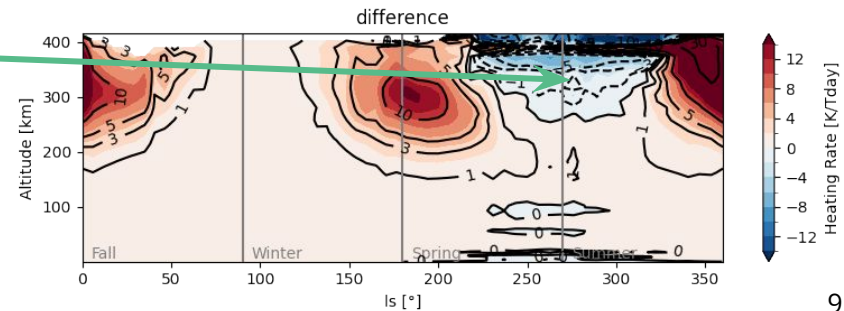
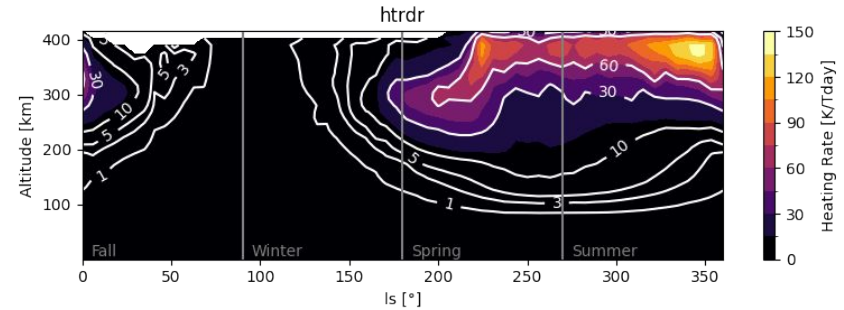
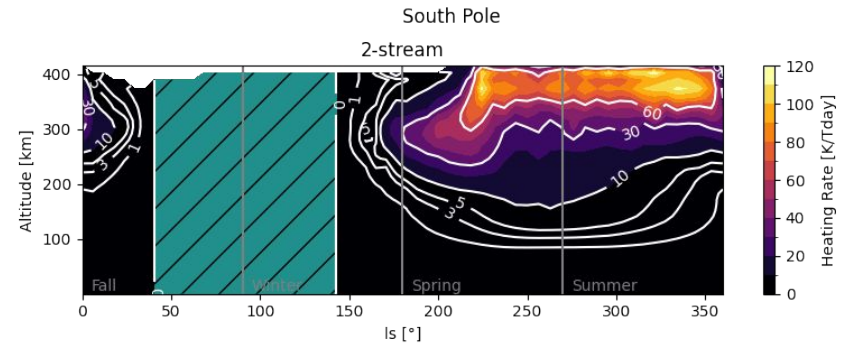
Main differences at equinox

Lower heating rates during spring/summer

plane-parallel

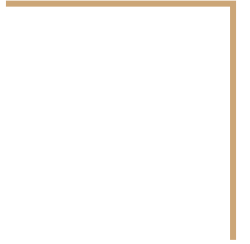
htrdr
3D

differences





Coupling with Titan LMDZ PCM



Preliminary results: circulation

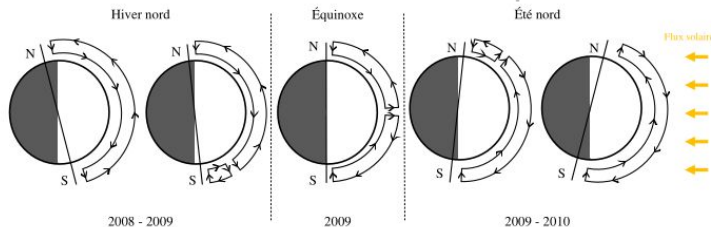
Color Scale: zonal wind

White lines: meridional circulation

Dotted: anti-clockwise rotation

Solid: clockwise rotation

© Bruno de Batz de Trenquelléon



Northern spring

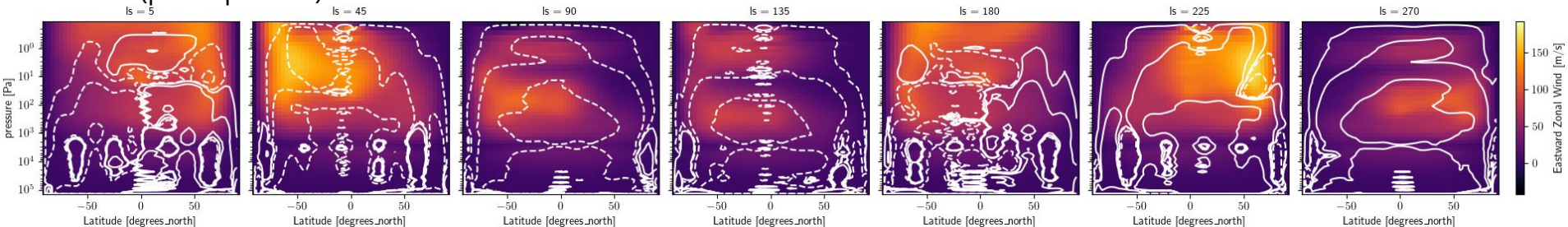
Northern summer

Northern fall

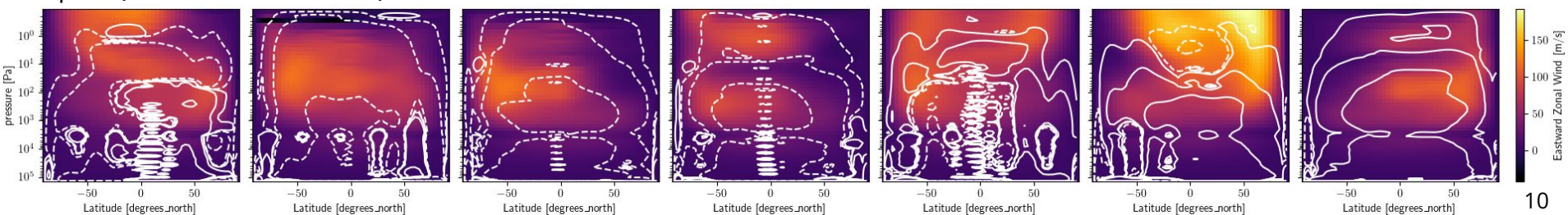
Northern winter



Reference (plane-parallel)



Coupled (3D radiative transfer)



Preliminary results: circulation

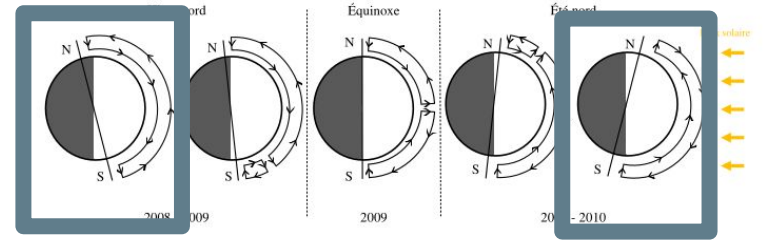
© Bruno de Batz de Trenquelléon

Color Scale: zonal wind

White lines: meridional circulation

Dotted: anti-clockwise rotation

Solid: clockwise rotation



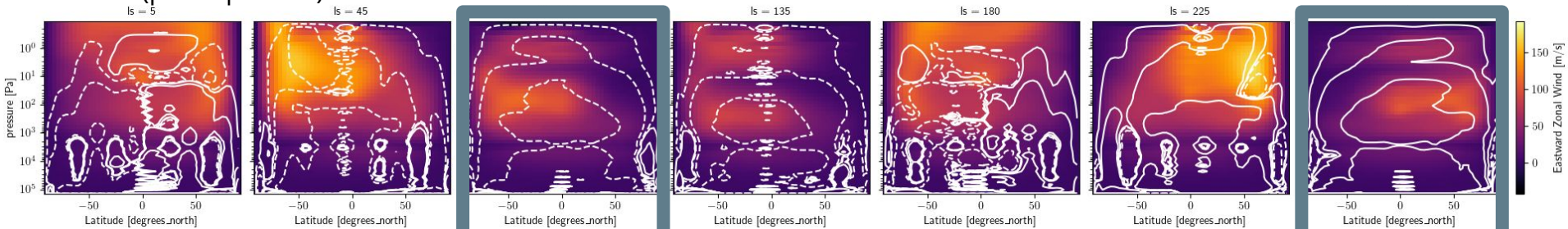
Northern spring

Northern summer

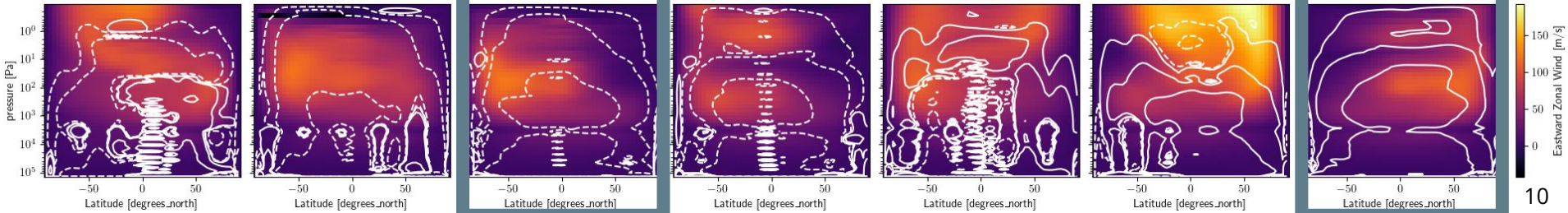
Northern fall

Northern winter

Reference (plane-parallel)



Coupled (3D radiative transfer)



Preliminary results: circulation

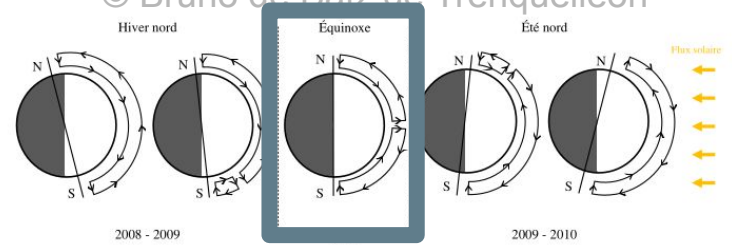
© Bruno de Batz de Trenquelléon

Color Scale: zonal wind

White lines: meridional circulation

Dotted: anti-clockwise rotation

Solid: clockwise rotation



Northern spring

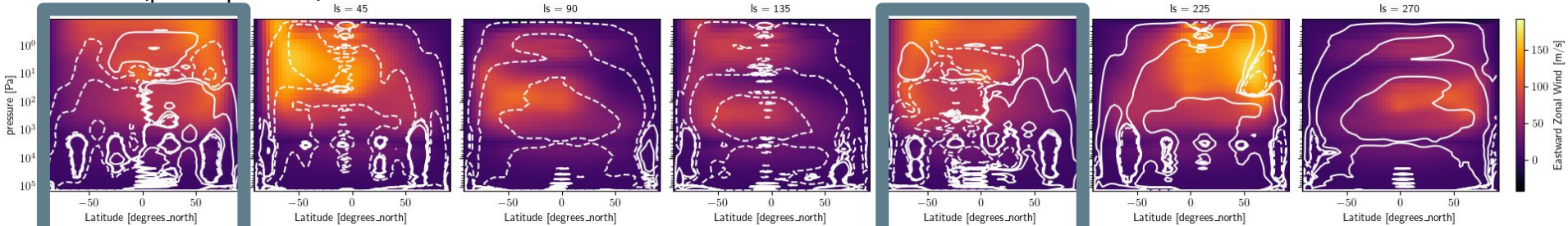
Northern summer

Northern fall

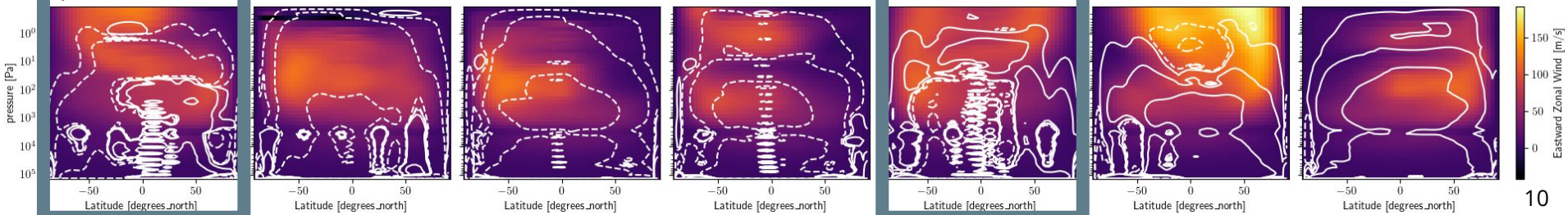
Northern winter



Reference (plane-parallel)



Coupled (3D radiative transfer)



Preliminary results: circulation

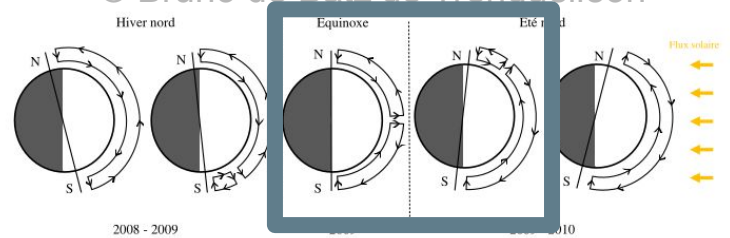
Color Scale: zonal wind

White lines: meridional circulation

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Solid: clockwise rotation

© Bruno de Batz de Trenquelléon



2008 - 2009

2010

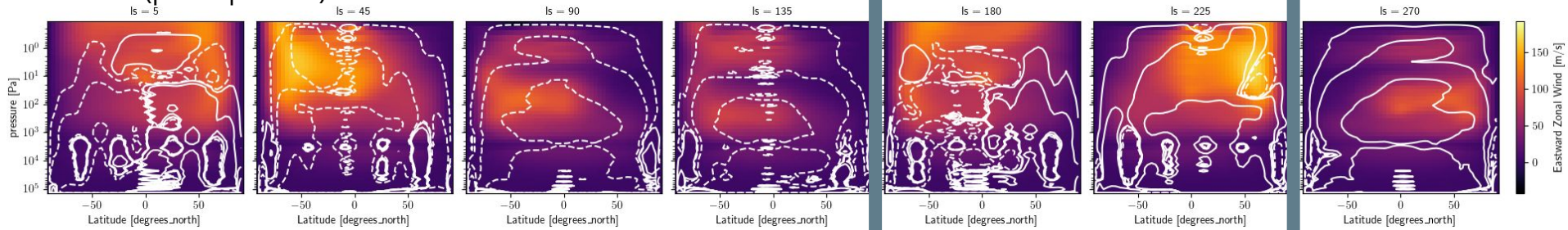
Northern spring

Northern summer

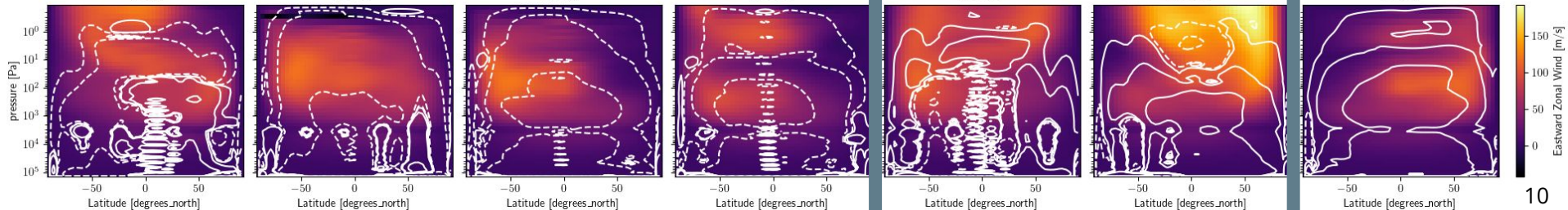
Northern fall

Northern winter

Reference (plane-parallel)

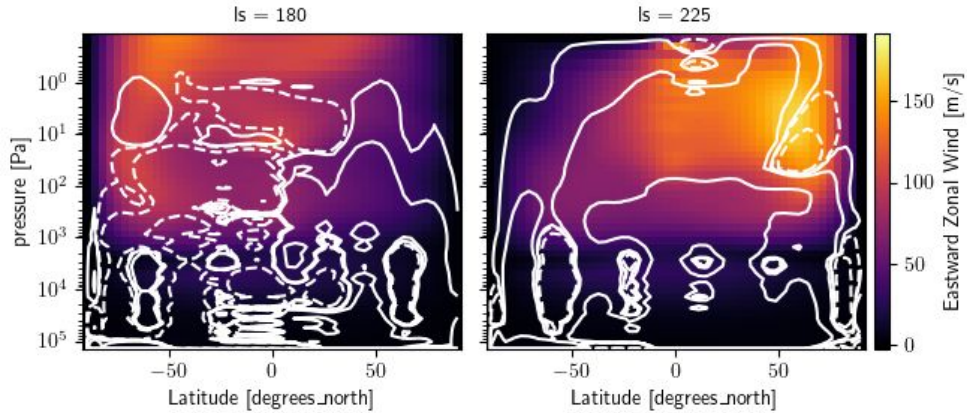


Coupled (3D radiative transfer)

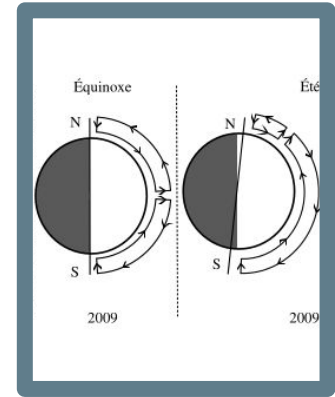
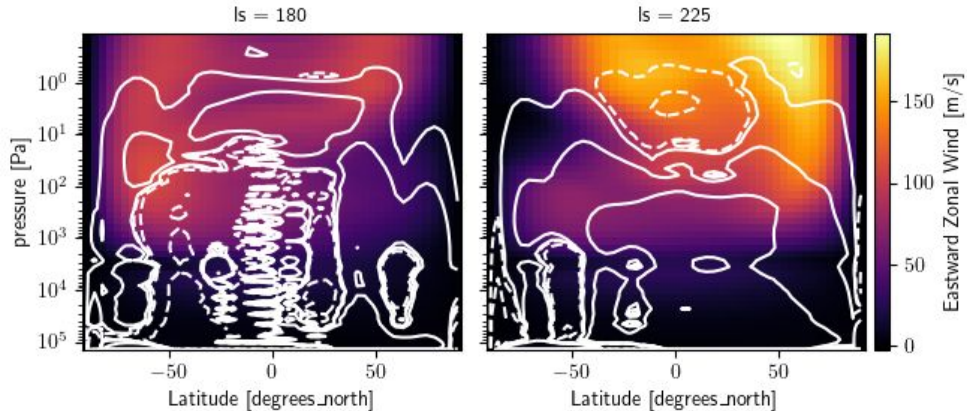


Preliminary results: circulation

Reference
(plane-parallel)

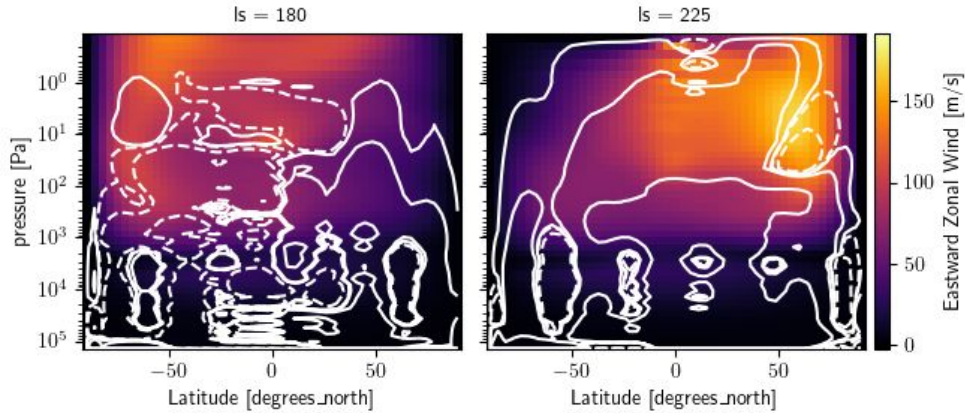


Coupled
(3D
radiative
transfer)

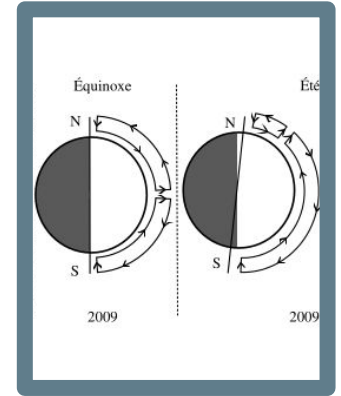
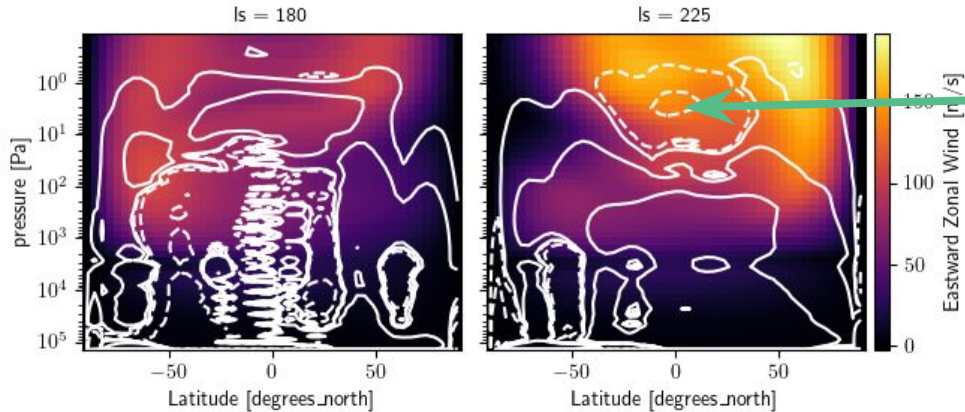


Preliminary results: circulation

Reference
(plane-parallel)

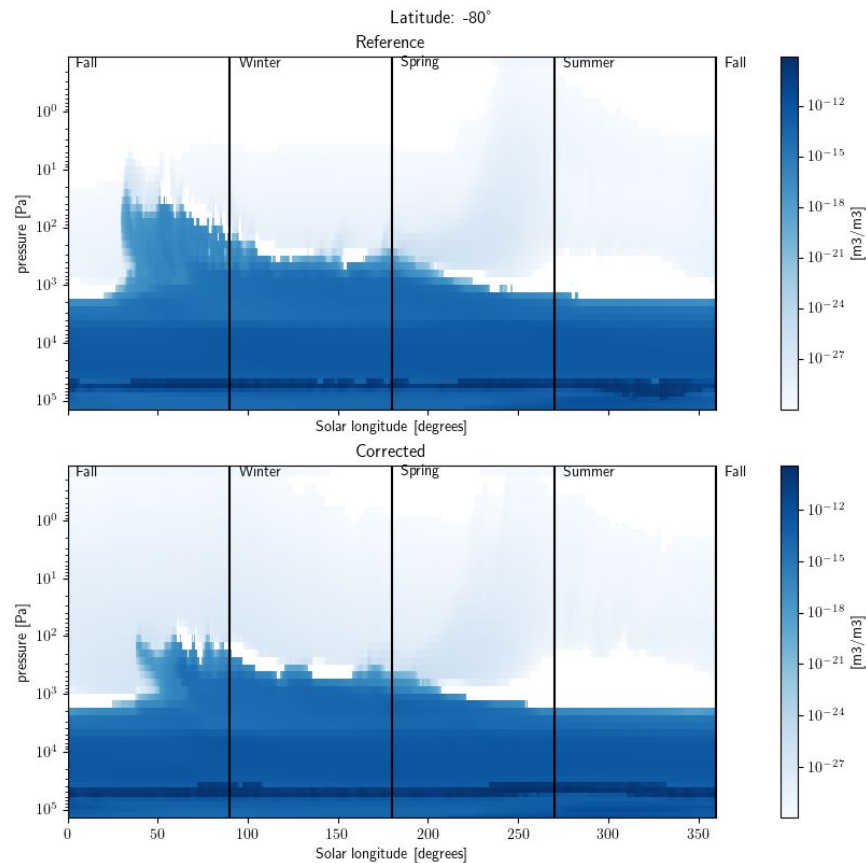


Coupled
(3D
radiative
transfer)



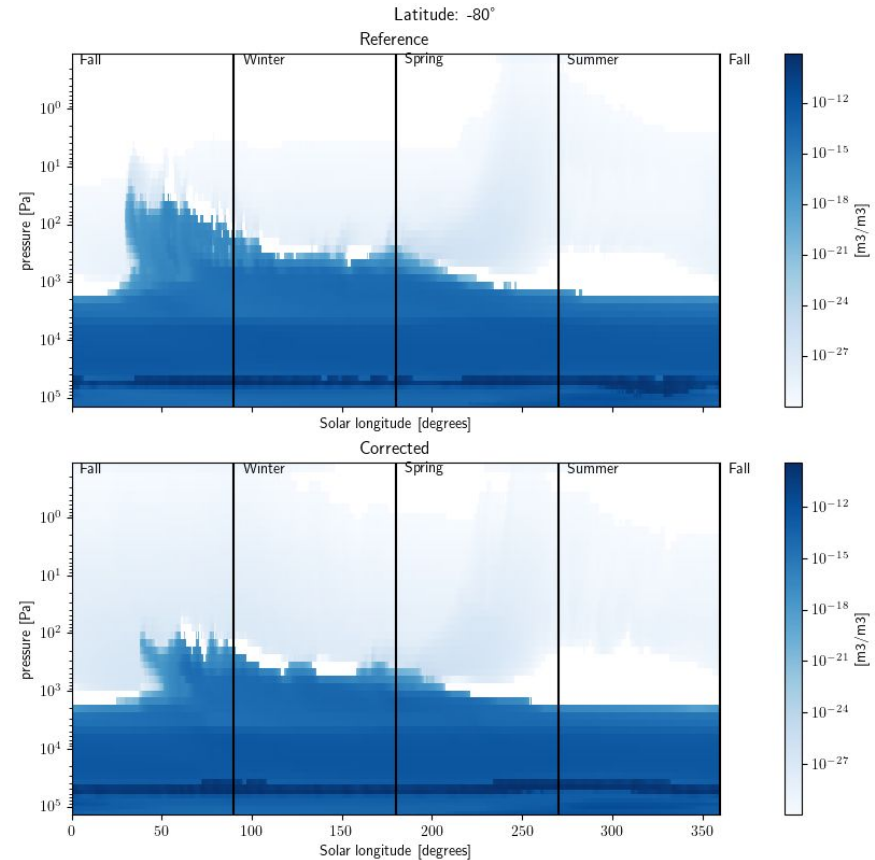
Additional cells in
upper stratosphere
after the equinox

Preliminary results: polar cloud



Preliminary results: polar cloud

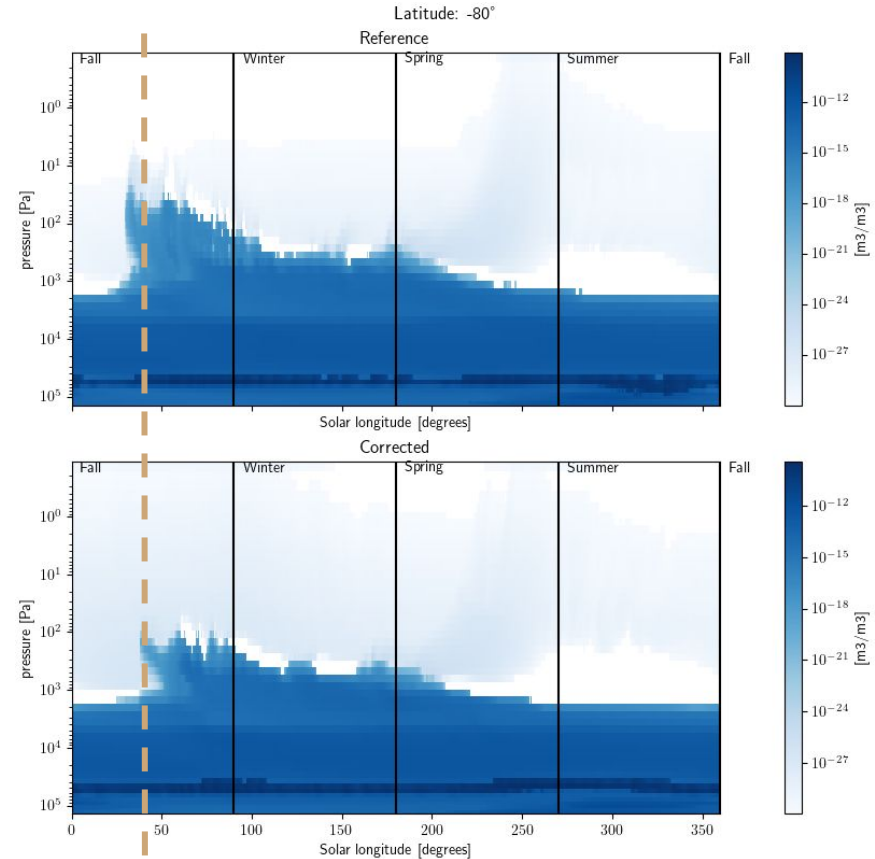
Heating of the upper layer of the polar region during winter:



Preliminary results: polar cloud

Heating of the upper layer of the polar region during winter:

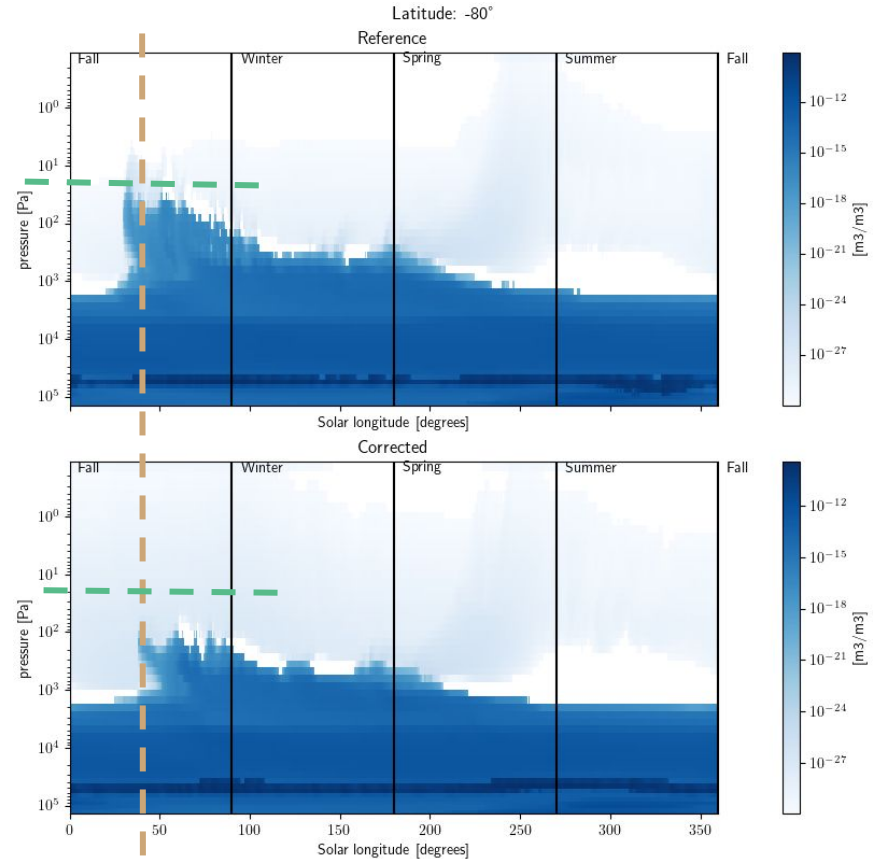
- Polar cloud forms later



Preliminary results: polar cloud

Heating of the upper layer of the polar region during winter:

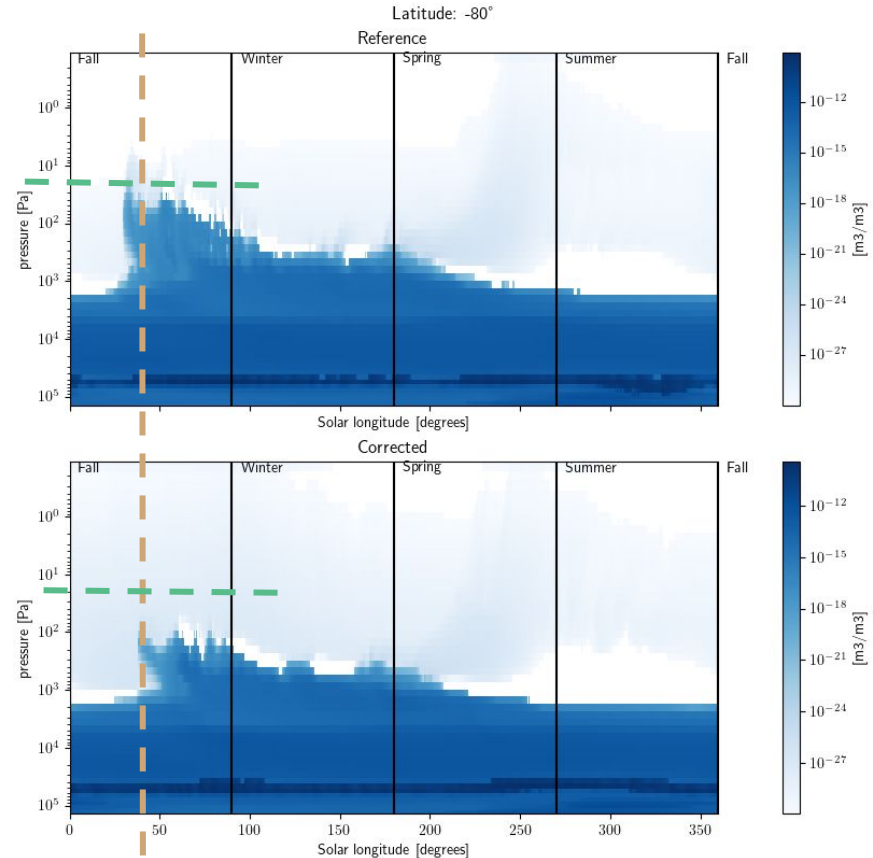
- Polar cloud forms later
- Starts at lower altitudes



Preliminary results: polar cloud

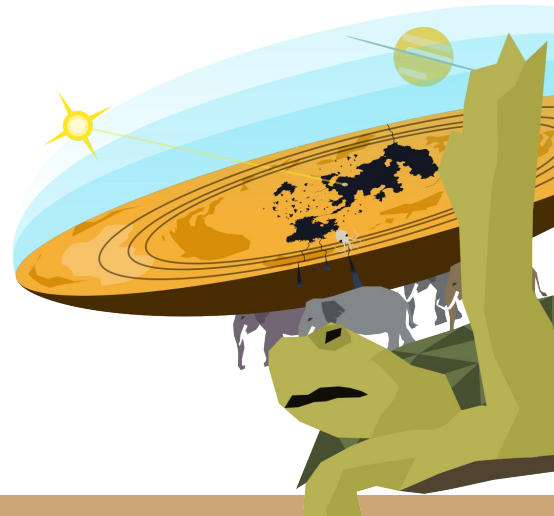
Heating of the upper layer of the polar region during winter:

- Polar cloud forms later
- Starts at lower altitudes
- Does not match observations anymore...



Takeaways:

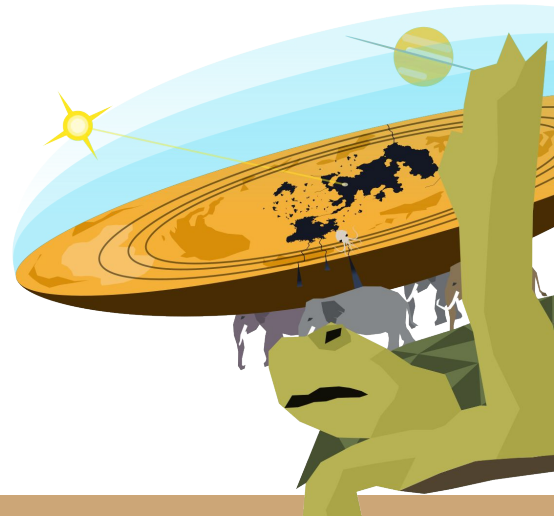
I am looking for a postdoc



Takeaways:

I am looking for a postdoc

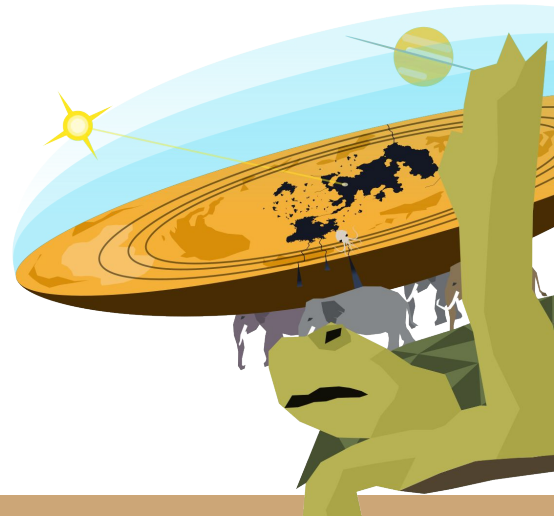
- Major effects along terminator and polar regions



Takeaways:

I am looking for a postdoc

- Major effects along terminator and polar regions
- Modify/delay dynamics



Takeaways:

I am looking for a postdoc

- Major effects along terminator and polar regions
- Modify/delay dynamics
- Impede the formation of the polar cloud



Takeaways:

I am looking for a postdoc

- Major effects along terminator and polar regions
 - Modify/delay dynamics
 - Impede the formation of the polar cloud
- ⇒ Is Titan flat after all?



Takeaways:

I am looking for a postdoc

- Major effects along terminator and polar regions
- Modify/delay dynamics
- Impede the formation of the polar cloud
⇒ Is Titan flat after all?

No! Just need some tuning...





Thank you!

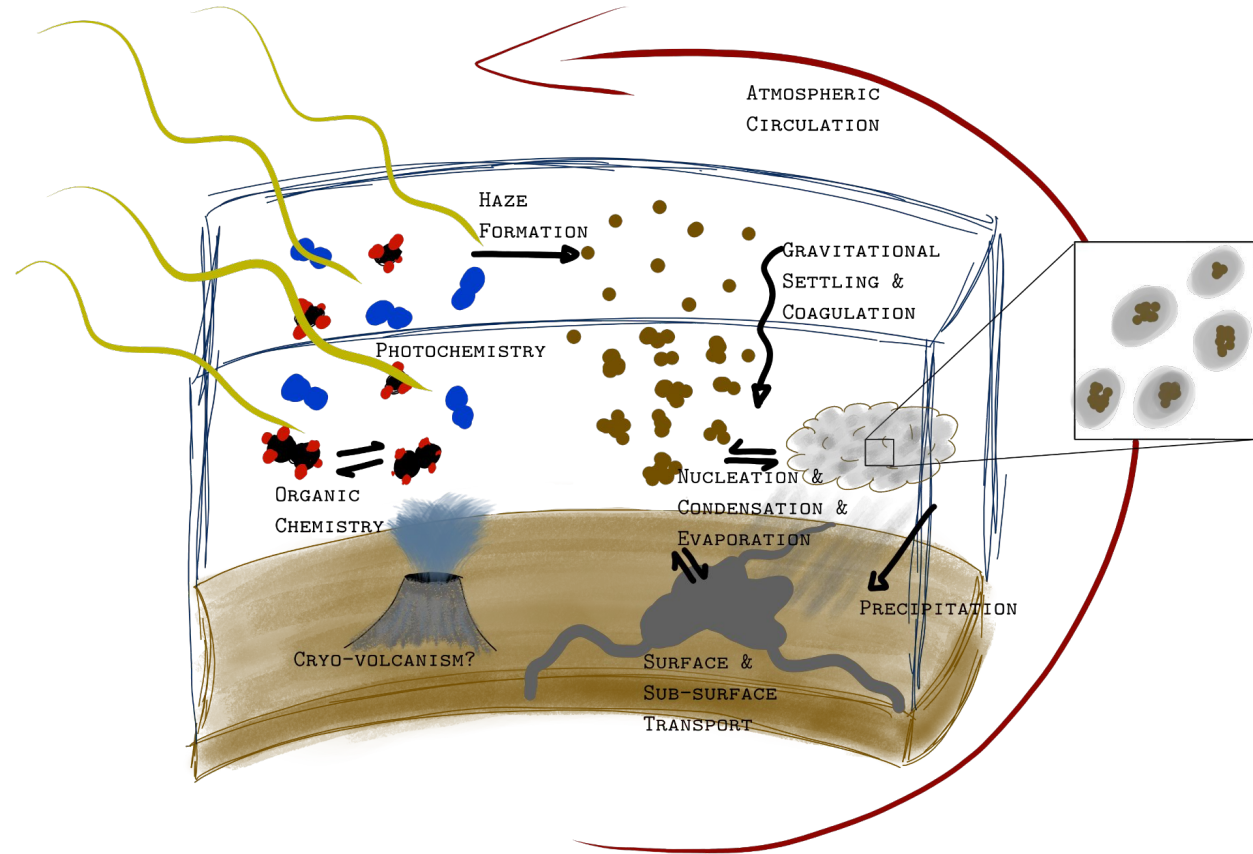
p.s. I am still looking for a postdoc ...



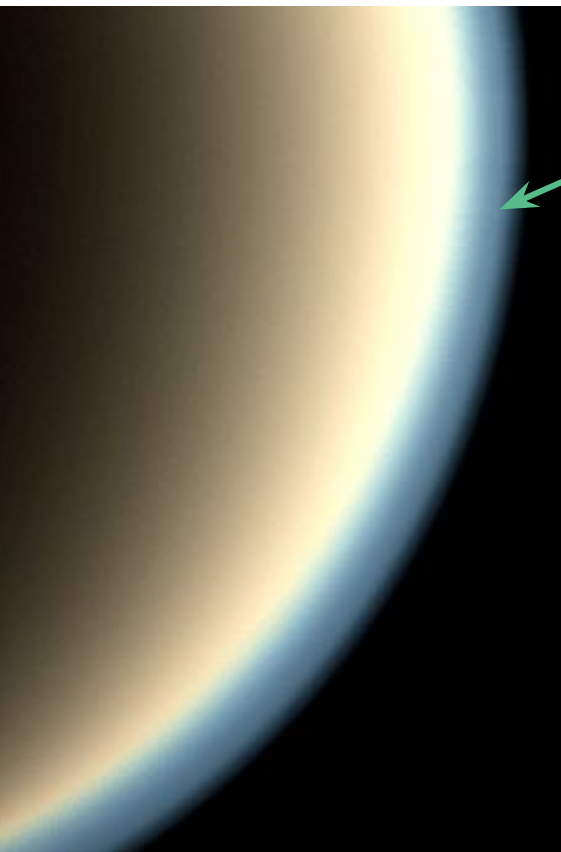
This work has been funded by the ~~Flat Titan Society~~ ANR RaD³-net

Titan in a nutshell

- Organic chemistry
- Photochemistry
- Haze & Cloud microphysics
- HCN below the tropopause
- CH₄ clouds deeper
- Surface / Atmosphere exchange



Titan main features



Detached haze layer

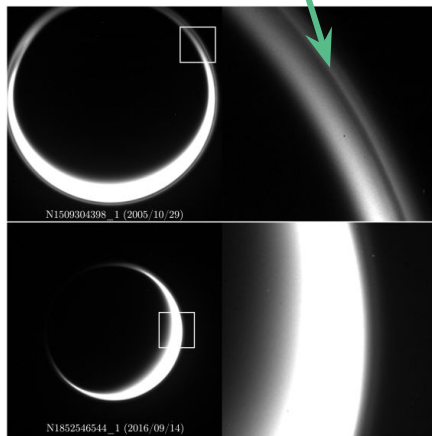
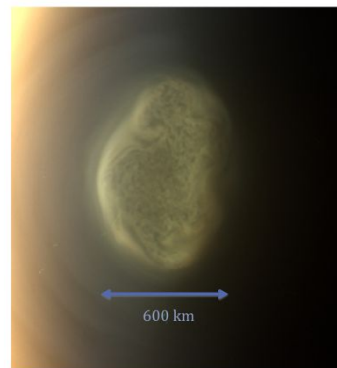


Figure S3: High phase angle images of the main and detached hazes of Titan before the collapse of the detached haze (top) and at the time of the reappearance (bottom). The observations are quite similar in terms of geometry with 155° and 142° phase angle for the 2005 and 2016 images. Images on the right half of the figure show the equatorial region corresponding to the squares in the left half.

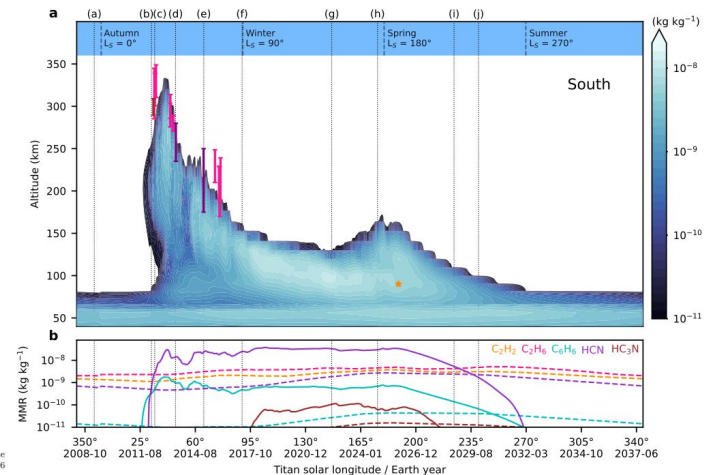
West et al. (2018)

West et al. (2016)



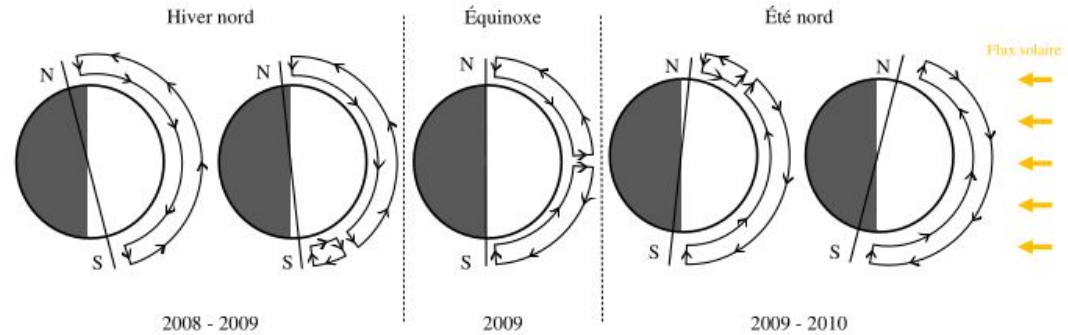
Polar cloud

Fig. 6. This image taken on 2012-178 (26 June) shows a previously unseen cloud patch near Titan's south pole. Solar illumination is from the left. This color image was constructed from images N1719446170 (red), N1719446086 (green) and N1719446000 (blue).



de Batz de Trenquelléon et al. (2026)

Titan dynamics



(Bruno De Batz de Tranquelléon 2024)

Super-rotation of the stratosphere

At solstices: A single Hadley cell from

Upwelling at summer pole, Downwelling at winter pole

During winter / summer: inversion of the circulation (split in 2 cells)

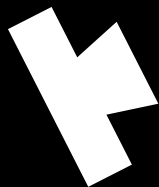
At equinoxes: 2 cells

Upwelling at equator, Downwelling at the poles

Source

Ray tracing procedure

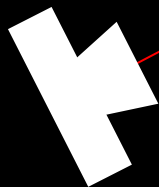
$w = 0$



Source

Ray tracing procedure

$w = 0$



Sampling of a direction
and length

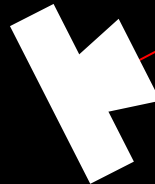
Ray tracing procedure

On Event:

- scattering ($p = k_s / k_{\text{ext}}$)

- absorption ($p = k_a / k_{\text{ext}}$)

$w = 0$



Sampling of a direction
and length

Source

Ray tracing procedure

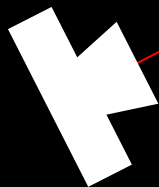
On Event:

- scattering ($p = k_s / k_{ext}$) \Rightarrow
 - sample ray in solar cone:
get solar contribution

- absorption ($p = k_a / k_{ext}$)

$$w = w + L^*$$

$$w = 0$$



Sampling of a direction
and length

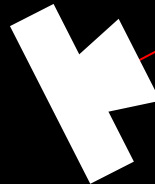
Source

Ray tracing procedure

On Event:

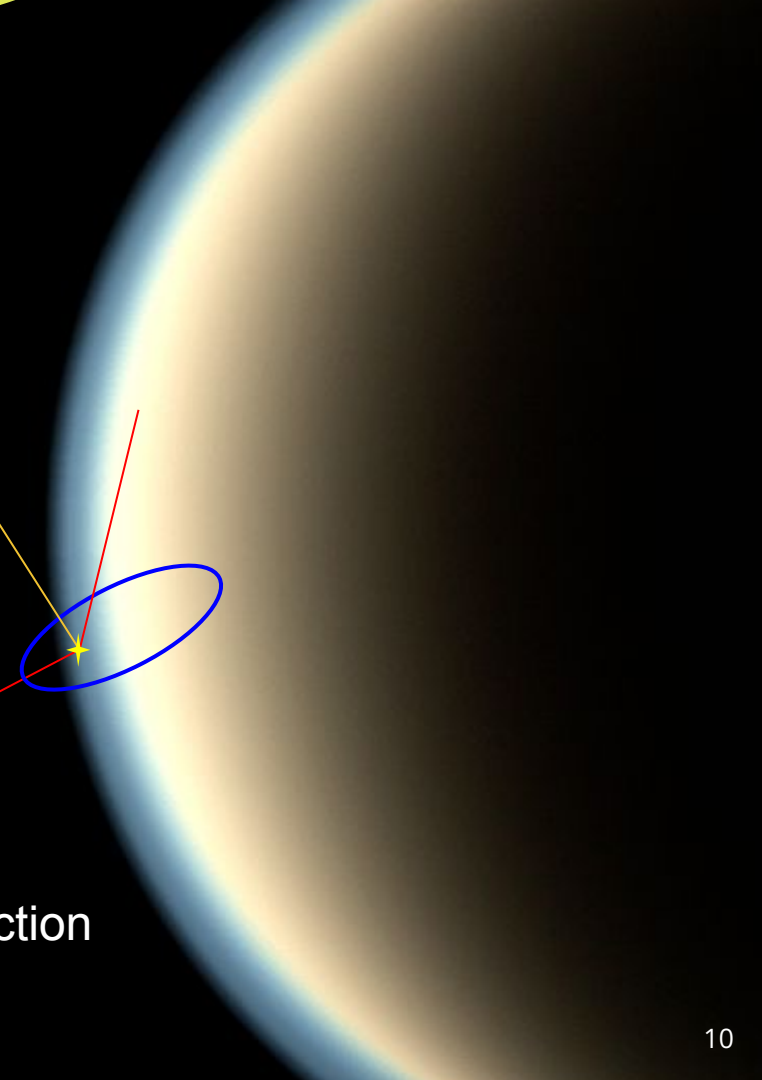
- scattering ($p = k_s / k_{ext}$) \Rightarrow
 - sample ray in solar cone:
get solar contribution
 - ray tracing resumes: new
direction sampled on
phase function
- absorption ($p = k_a / k_{ext}$)

$w = 0$



Sampling of a direction
and length

$$w = w + L^*$$



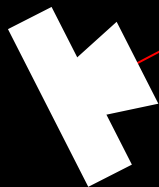
Source

Ray tracing procedure

On Event:

- scattering ($p = k_s / k_{ext}$) \Rightarrow
 - sample ray in solar cone:
get solar contribution
 - ray tracing resumes: new
direction sampled on
phase function
- absorption ($p = k_a / k_{ext}$)

$w = 0$



Sampling of a direction
and length

$w = w + L^*$

$w = w + L^*$

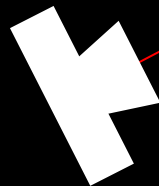
Source

Ray tracing procedure

On Event:

- scattering ($p = k_s / k_{ext}$) \Rightarrow
 - sample ray in solar cone:
get solar contribution
 - ray tracing resumes: new
direction sampled on
phase function
- absorption ($p = k_a / k_{ext}$) \Rightarrow ray is
terminated

$w = 0$



Sampling of a direction
and length

$w = w + L^*$

$w = w + L^*$

$w = w + L^*$

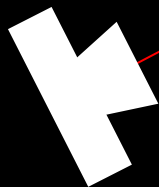
Source

Ray tracing procedure

On Event:

- scattering ($p = k_s / k_{ext}$) \Rightarrow
 - sample ray in solar cone:
get solar contribution
 - ray tracing resumes: new
direction sampled on
phase function
- absorption ($p = k_a / k_{ext}$) \Rightarrow ray is
terminated

$w = 0$



Sampling of a direction
and length

$w = w + L^*$

$w = w + L^*$

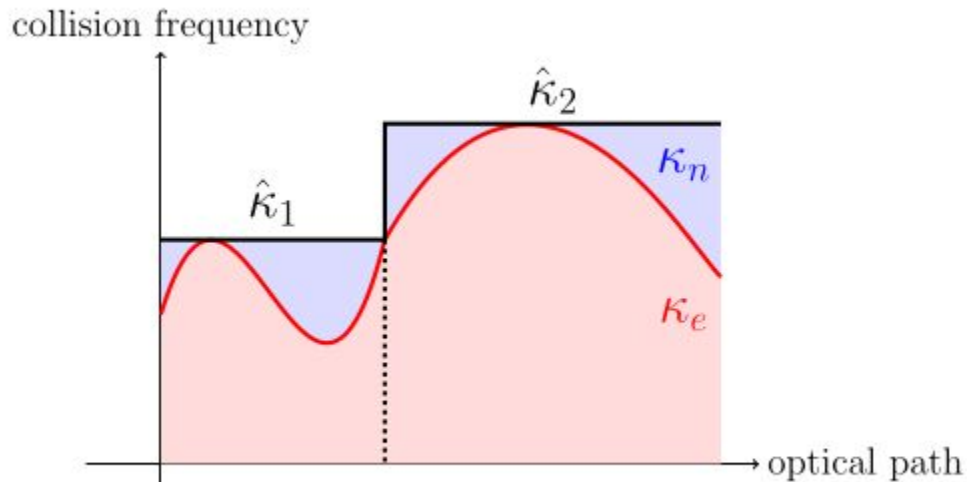
$w = w + L^*$

$$L = 1/N \sum w_i$$

Nul-collision algorithm

On Event:

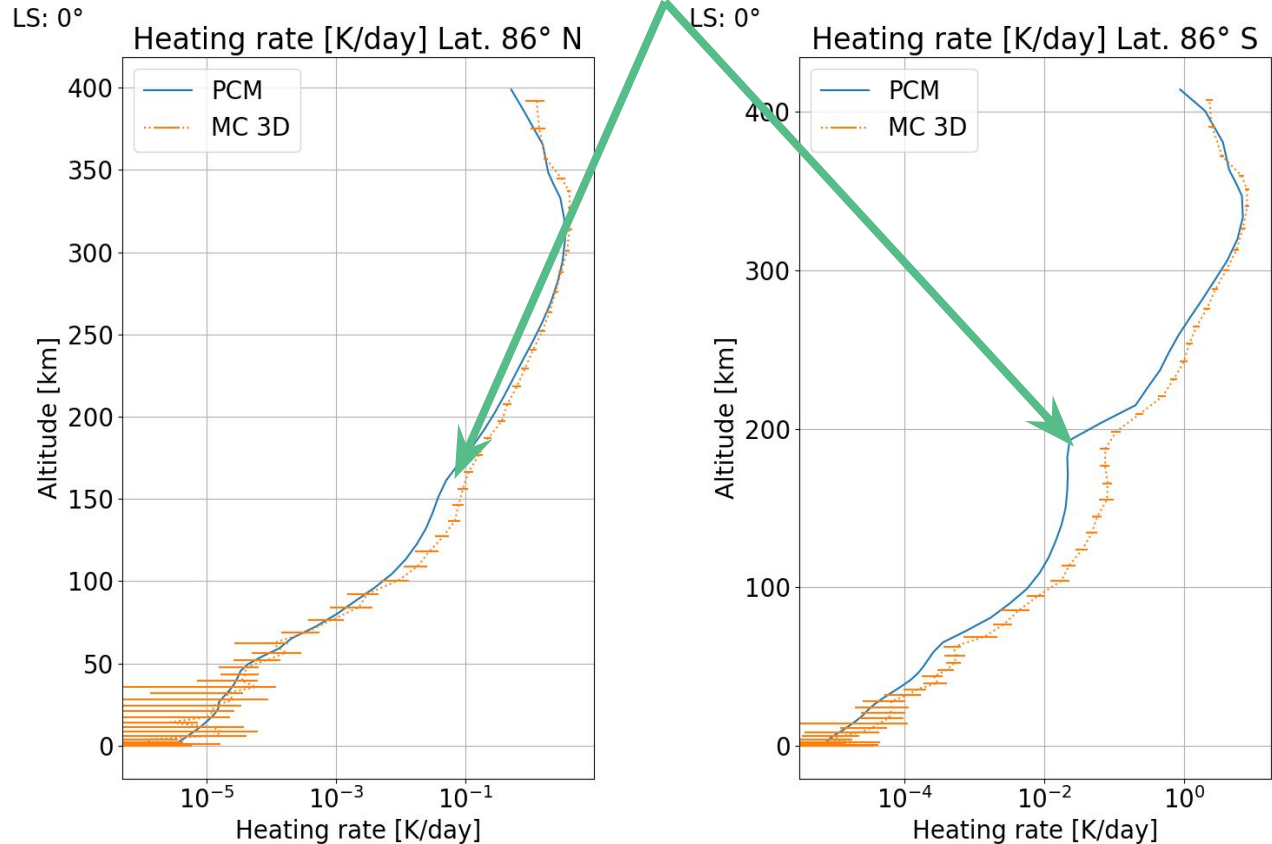
- scattering $(p=k_s/k_{\max}) \Rightarrow$
 - sample ray in solar cone:
get solar contribution
 - ray tracing resumes: new
direction sampled on
phase function
- absorption $(p=k_a/k_{\max}) \Rightarrow$ ray is
terminated
- nul-collision $(p=1-(k_a+k_s)/k_{\max}) \Rightarrow$ ray tracing resumes in same
direction



He et al. (2026)

Heterogeneity effects

North / South
asymmetry
below 200 km



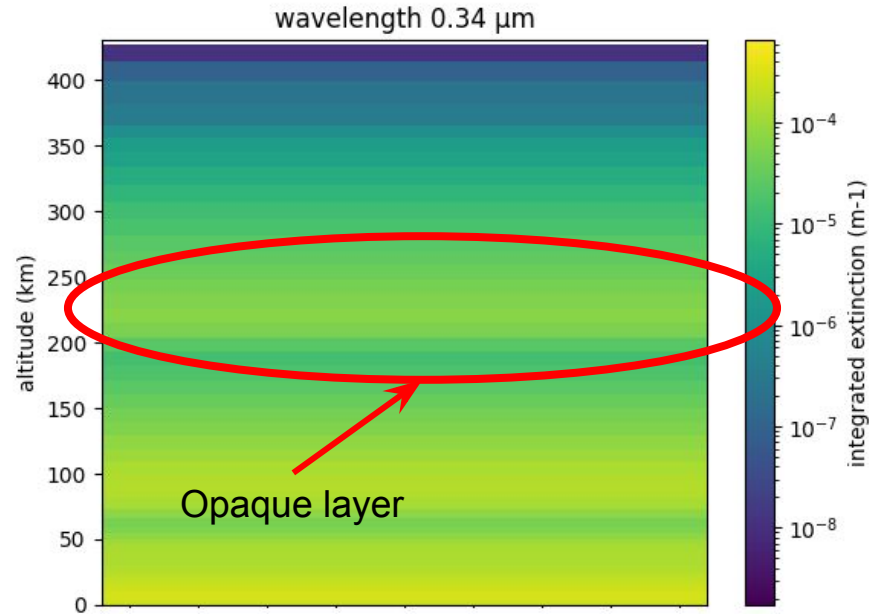
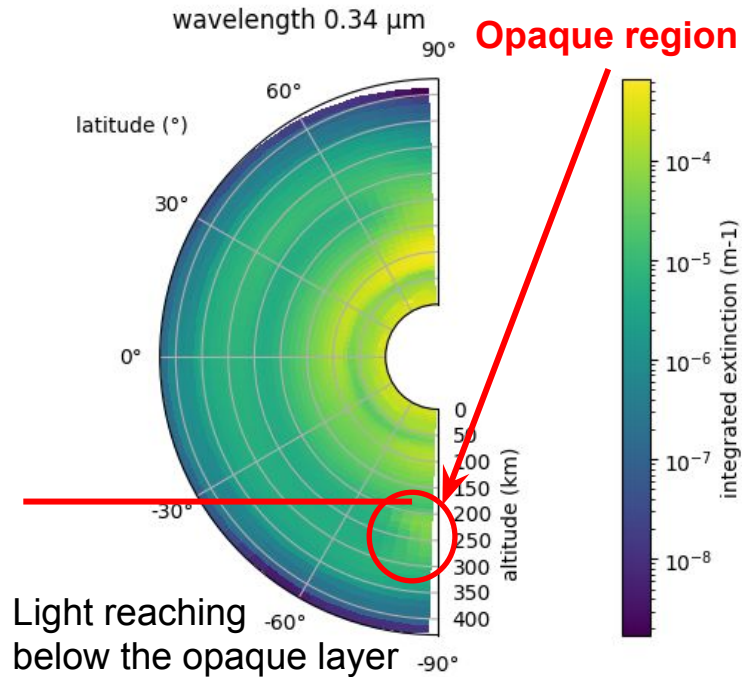
Zonal average

Northern spring
equinox

Comparing North
and South poles

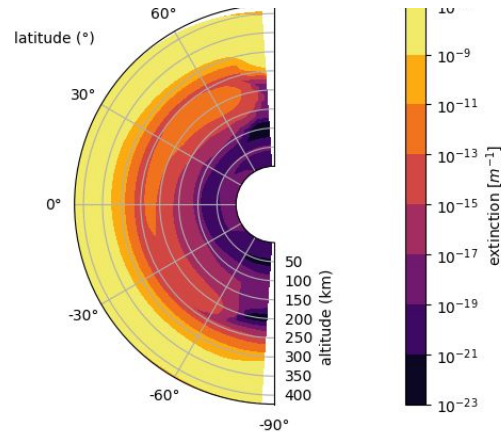
Heterogeneity effects

What a plane-parallel would see:



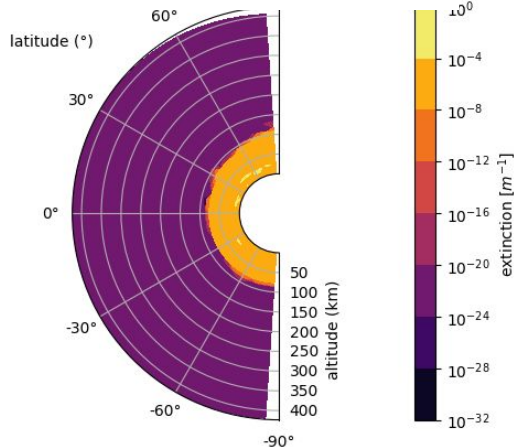
Heterogeneity effects

Small particles extinction

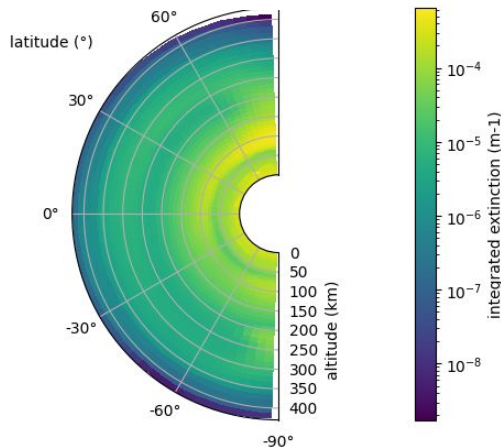


“Patch” of haze detached by circulation

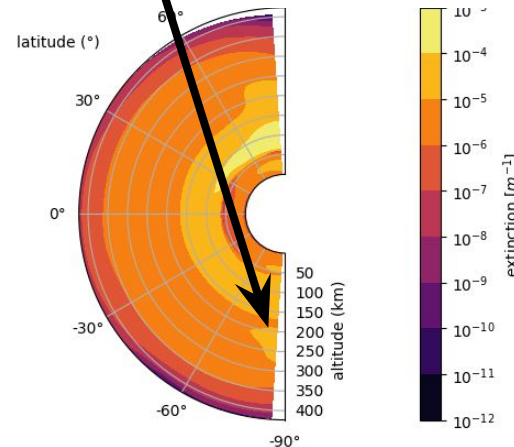
Cloud particles extinction



Total extinction

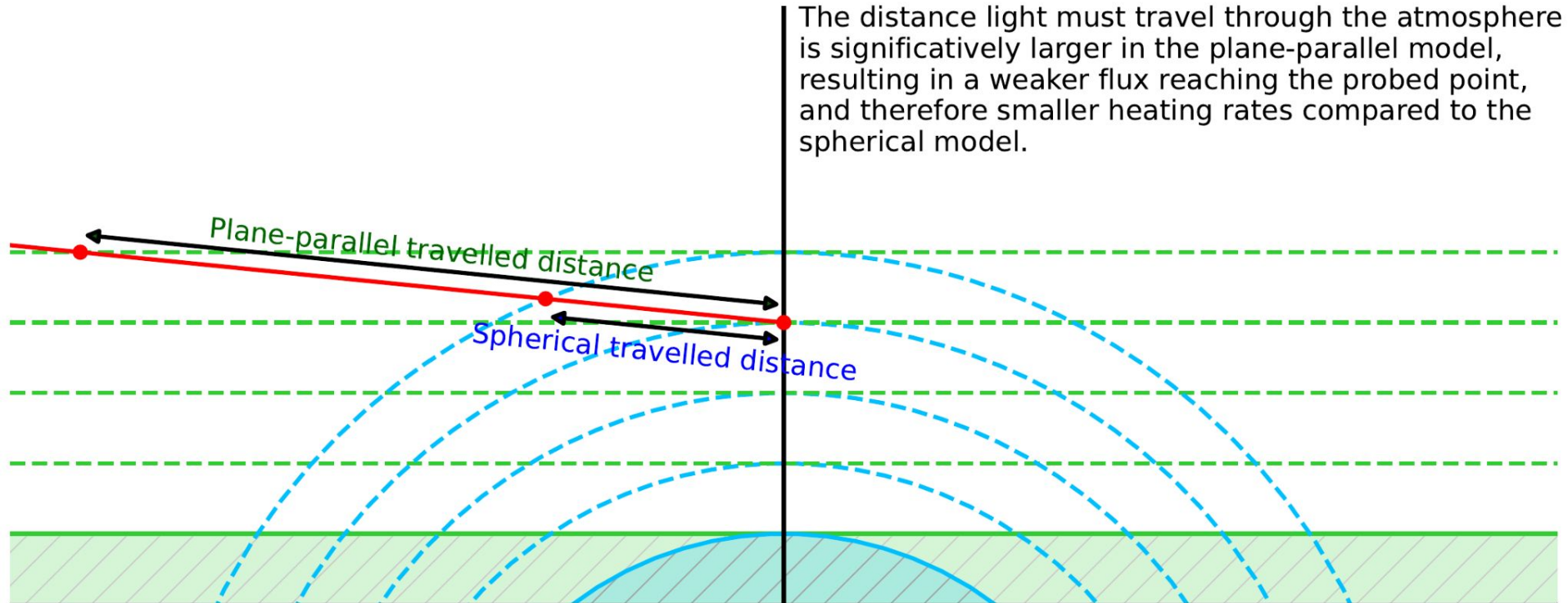


Fractal aggregates extinction



Terminator effects

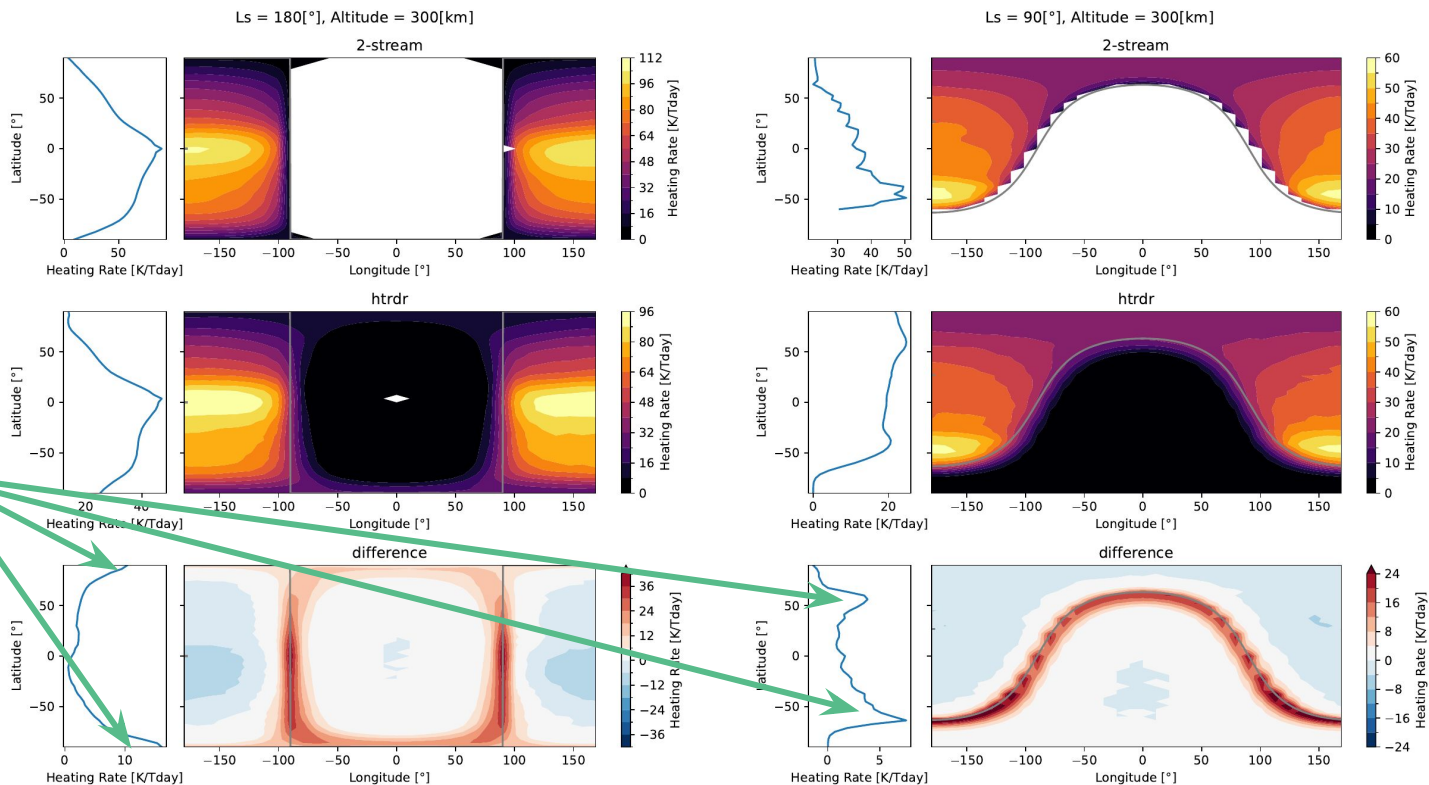
Higher heating rates along the terminator



Terminator effects

Differences
along the
terminator

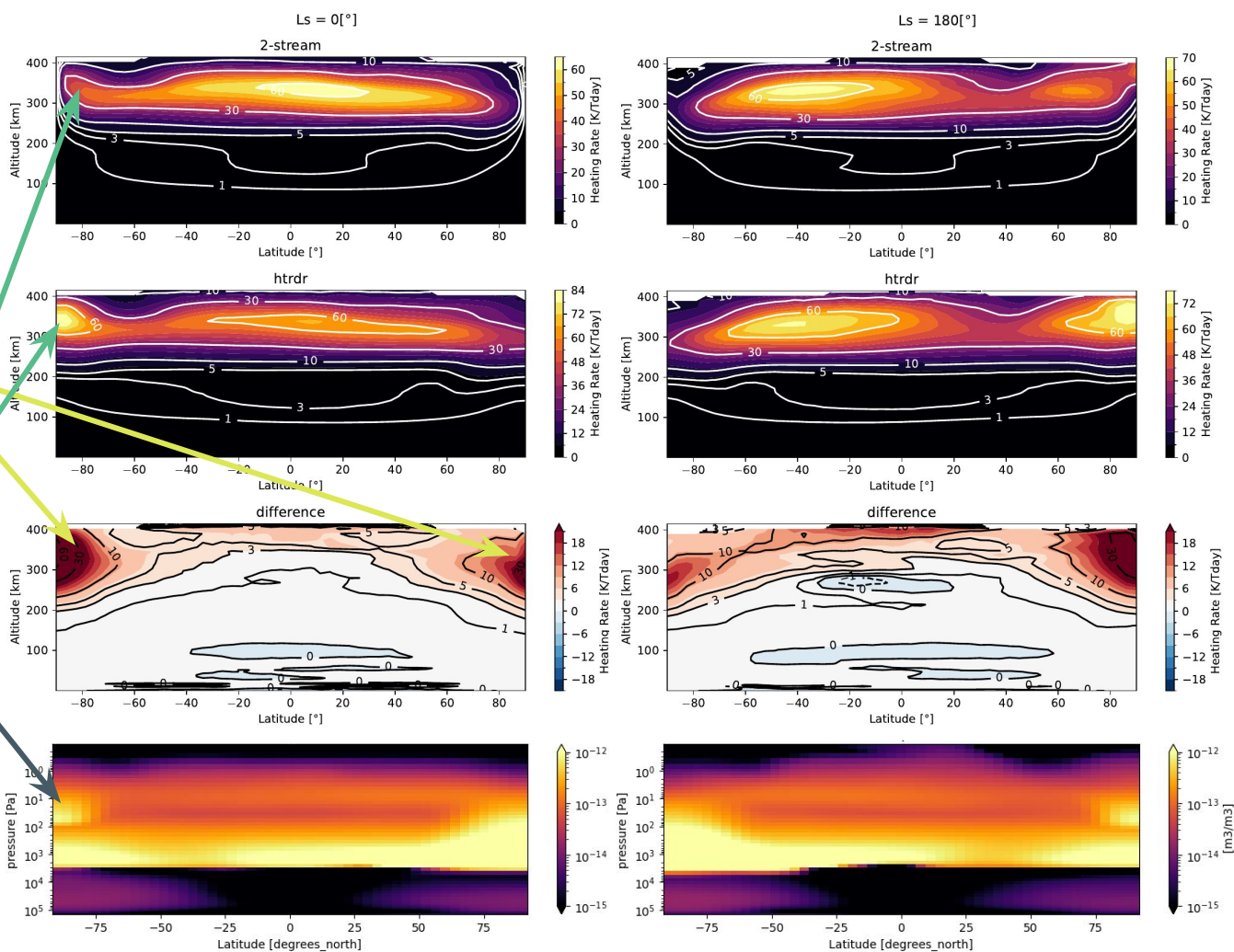
Stronger
effects near
limits of polar
night/day



Equinoxes

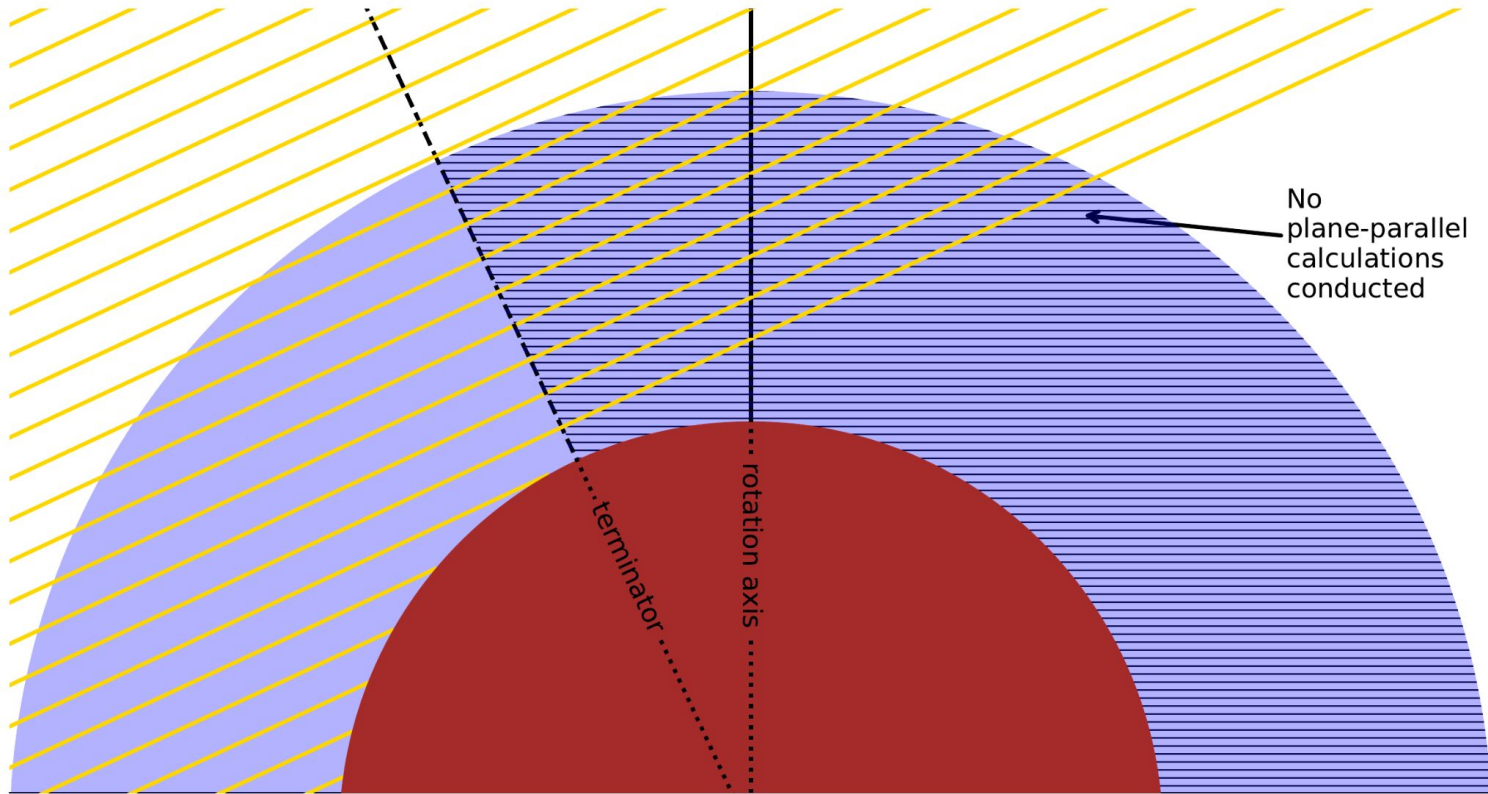
Main differences at the poles

Higher heating rates above former summer poles:
detached haze layer at the limb

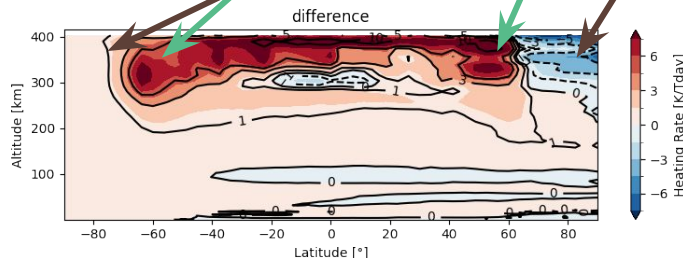
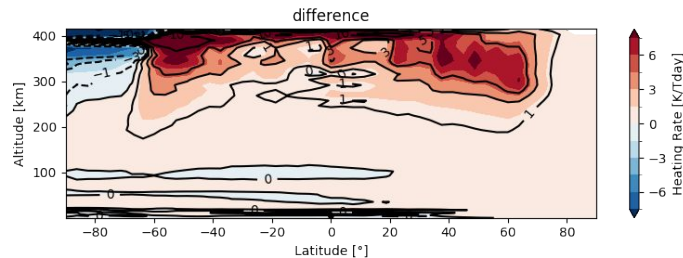
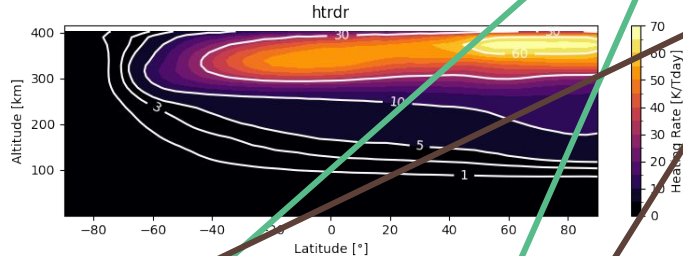
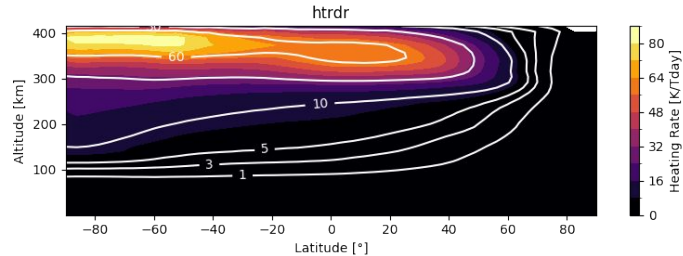
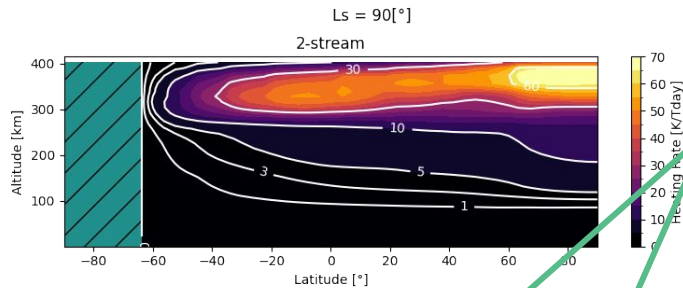
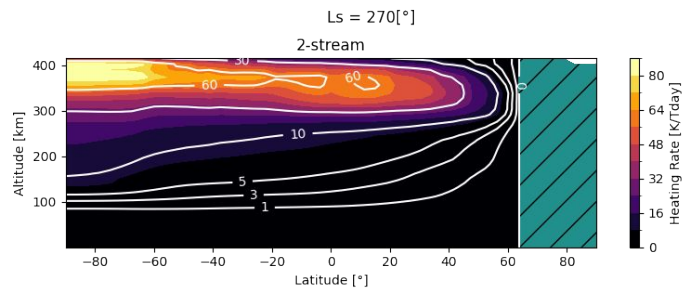


Effects on the polar regions

Higher heating rates in the polar regions



Solstices

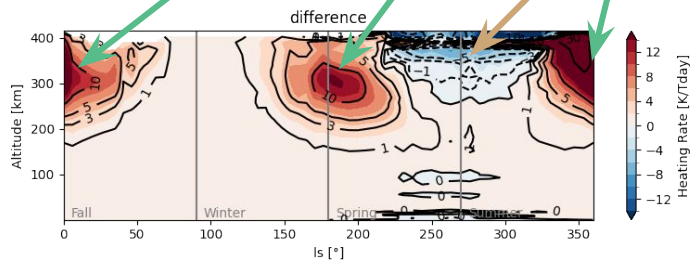
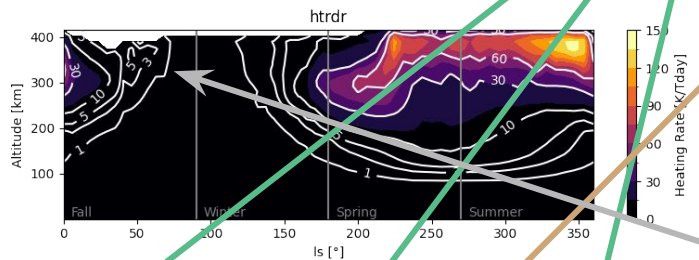
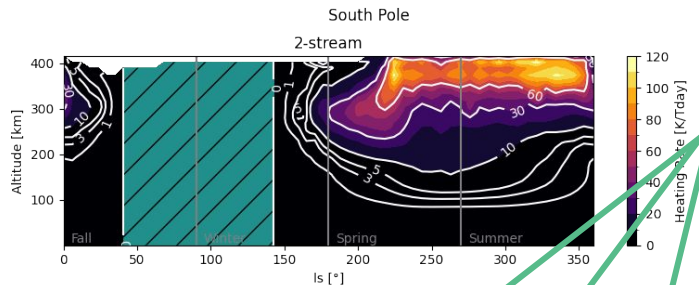
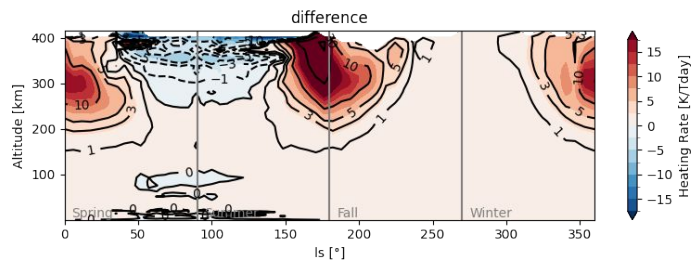
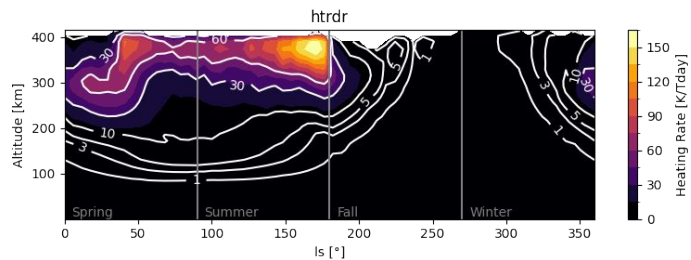
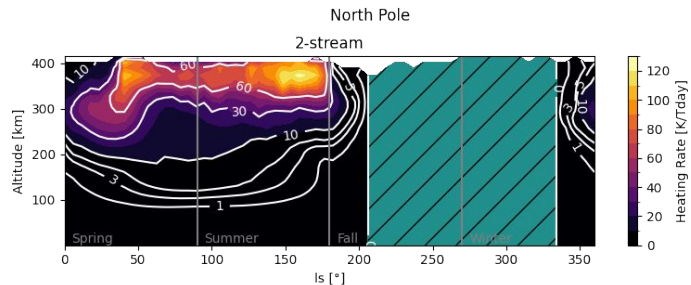


Main differences near 60° latitude: effect of terminator

Hotter polar night, cooler polar day

Higher heating rates at northern winter solstice: closer to periastron

Effects on the polar regions



Main differences at equinoxes

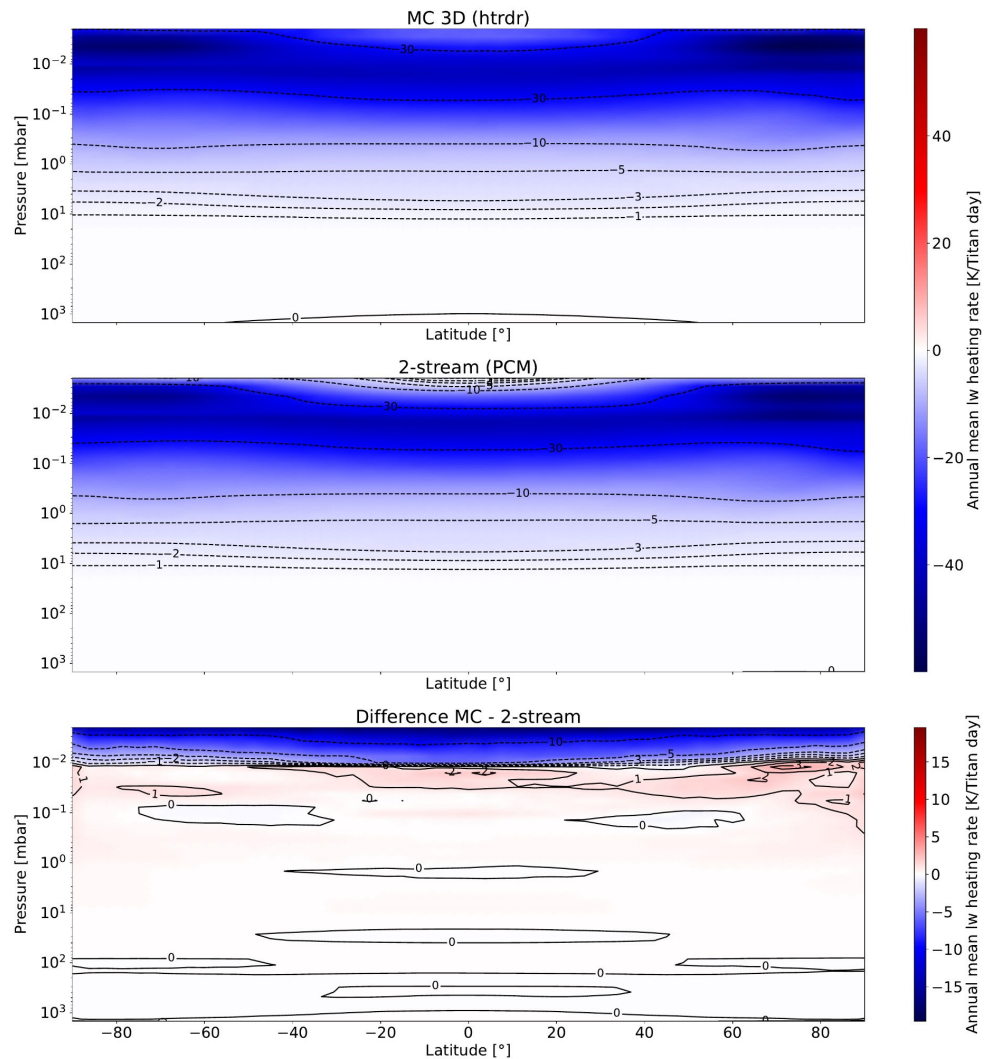
Lower heating rates during spring/summer

Non-zero heating rates in the polar region during winter

Longwave (LW)

Main differences very high up: boundary conditions of the PCM?

Small differences in main part of the atmosphere.



Coupling

1 calculation every Titan week (= 11 Titan days)

Calculation conducted on the sphere: Longitude \Leftrightarrow Local time

~50'000 samples per PCM cell

20-30 min calculation time on Roméo (+50% time on PCM)

Interpolation in Ls with previous year calculation

14 Titan years have run

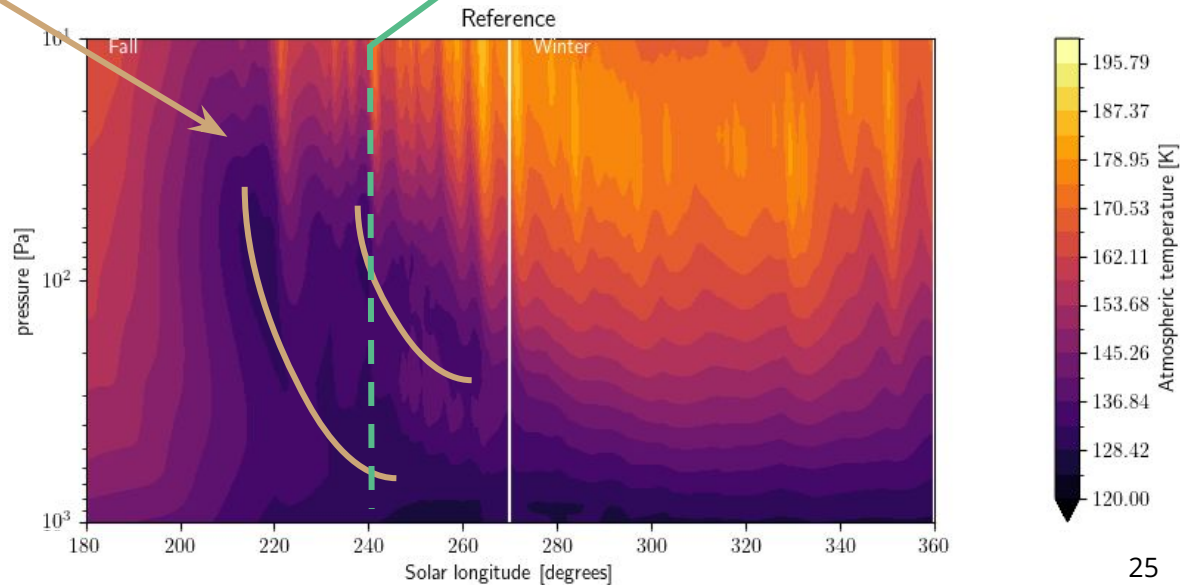
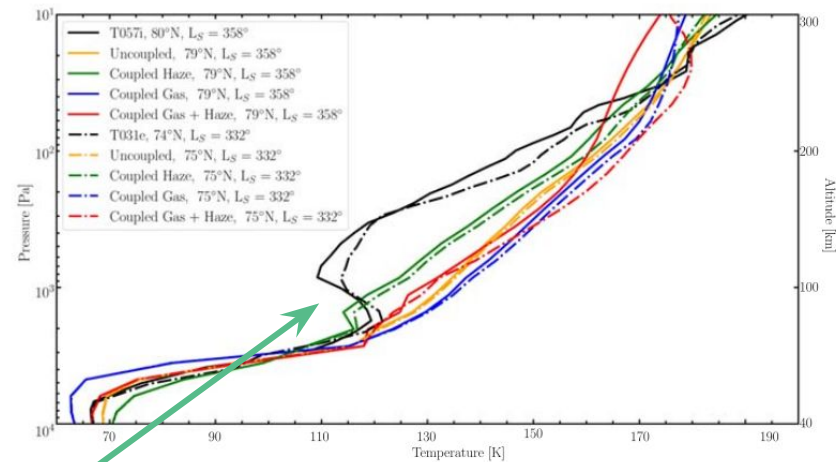
Reference case: polar shoulder

First cooling in upper atmosphere

Descent to deeper layers

Second descent of cooler air

Hotter atmosphere trapped between cold layers



Reference case: polar shoulder

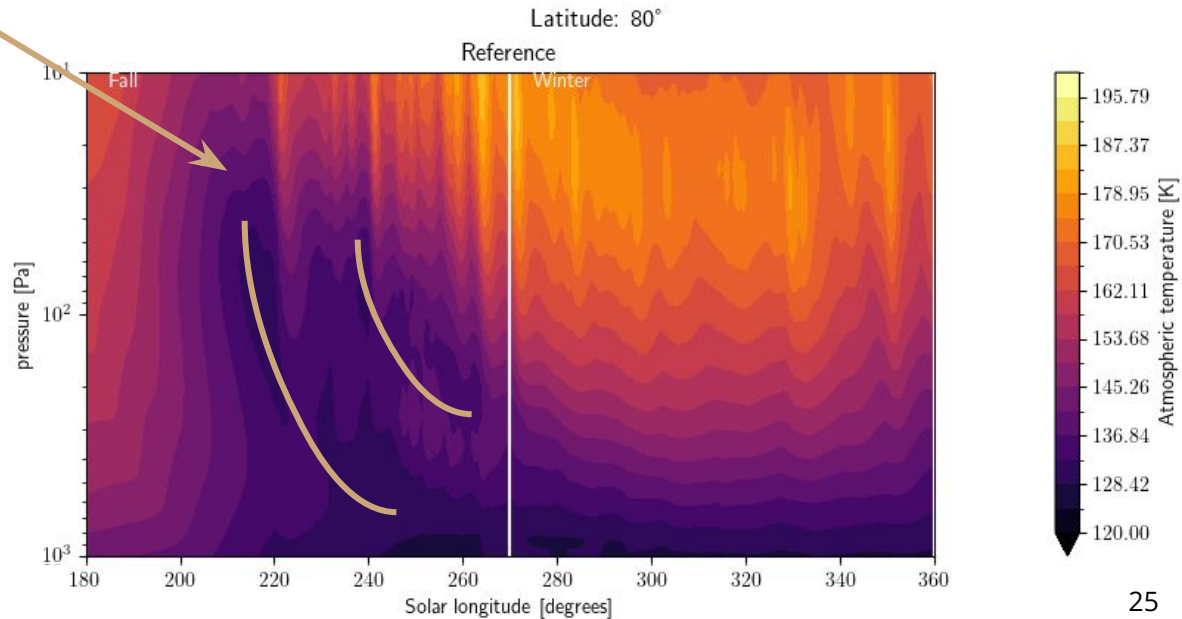
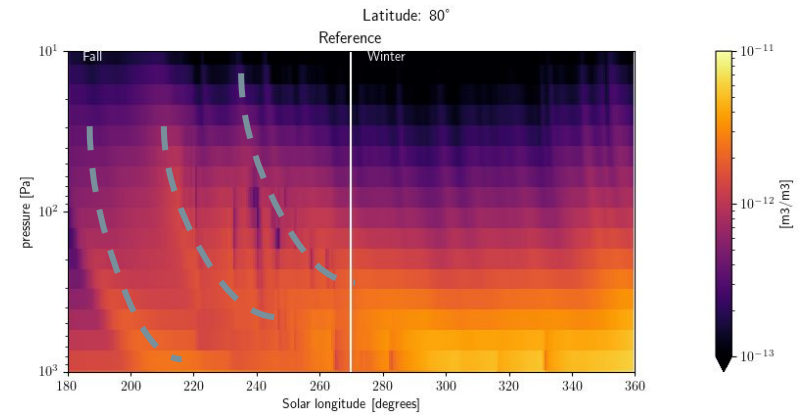
First cooling in upper atmosphere

Descent to deeper layers

Second descent of cooler air

Hotter atmosphere trapped between cold layers

Due to successive descent of haze particles



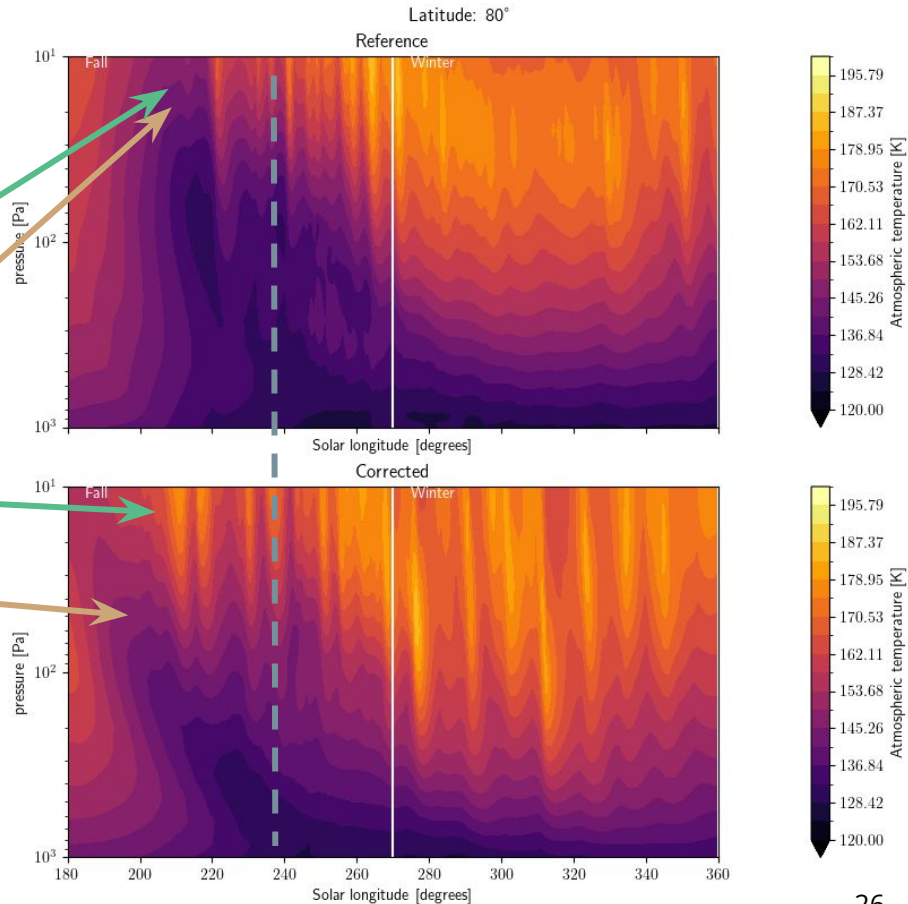
Preliminary results: temperature

Heating within the polar night

Hotter atmosphere

Cooling happens deeper

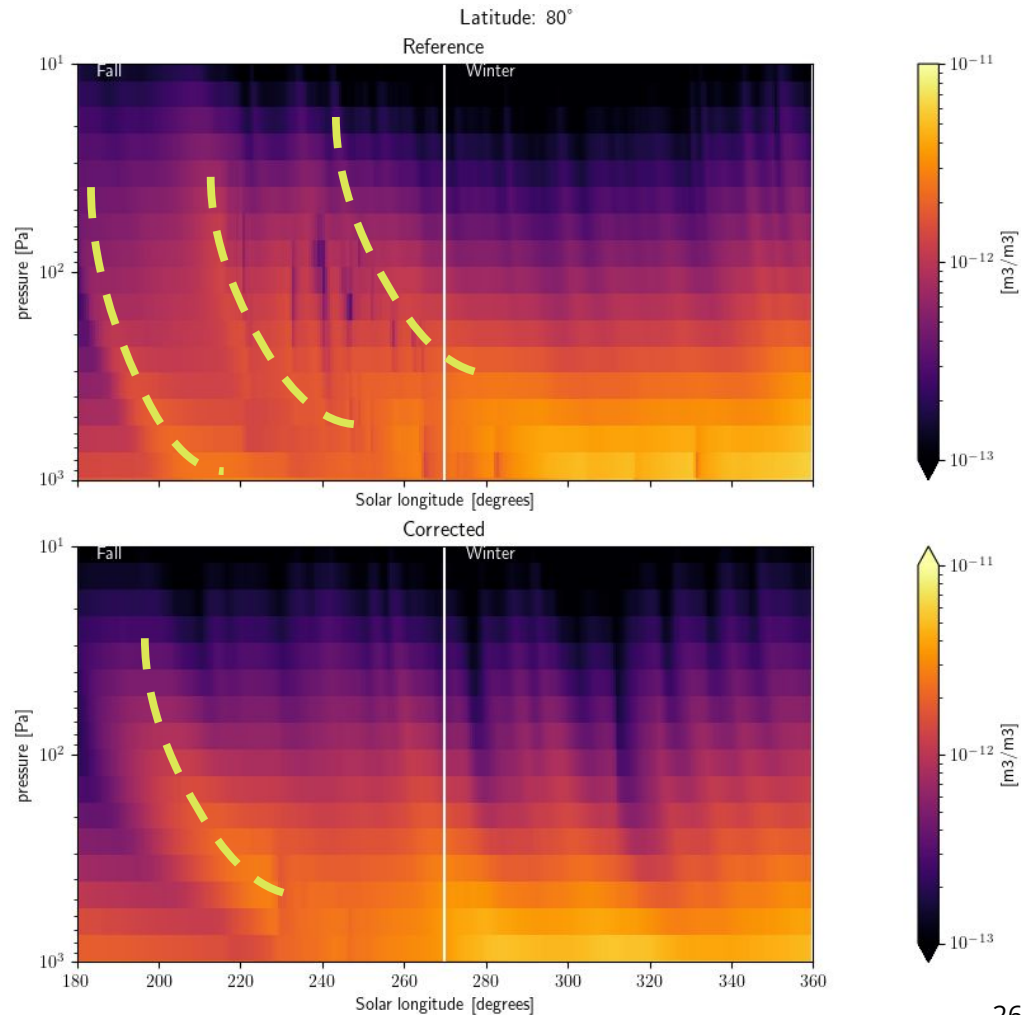
Disapparition of the polar shoulder ...



Preliminary results: haze

With correction: single descent

⇒ Effects on circulation ?



Preliminary results: circulation

Color Scale: zonal wind

White lines: meridional circulation

Dotted: anti-clockwise rotation

Solid: clockwise rotation

© Bruno de Batz de Trenquelléon



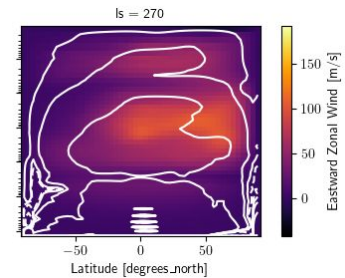
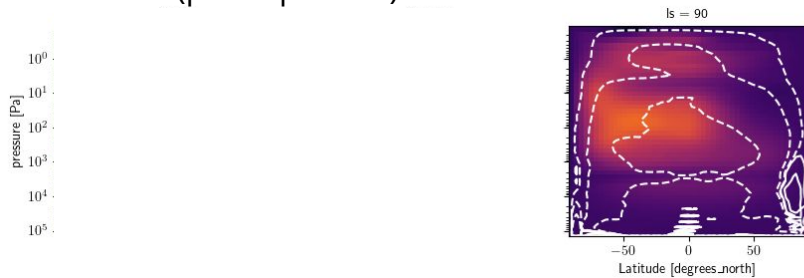
Northern spring

Northern summer

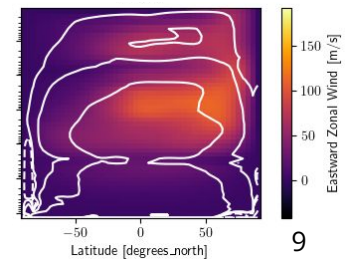
Northern fall

Northern winter

Reference (plane-parallel)



Coupled (3D radiative transfer)



Preliminary results: circulation

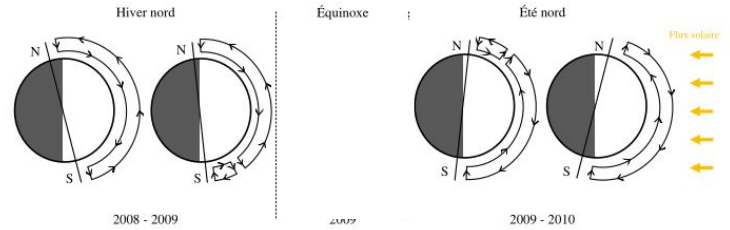
Color Scale: zonal wind

White lines: meridional circulation

Dotted: anti-clockwise rotation

Solid: clockwise rotation

© Bruno de Batz de Trenquelléon



Northern spring

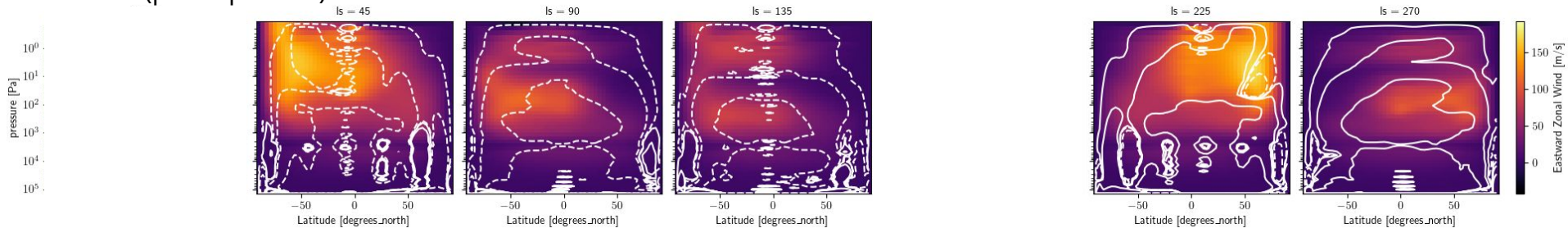
Northern summer

Northern fall

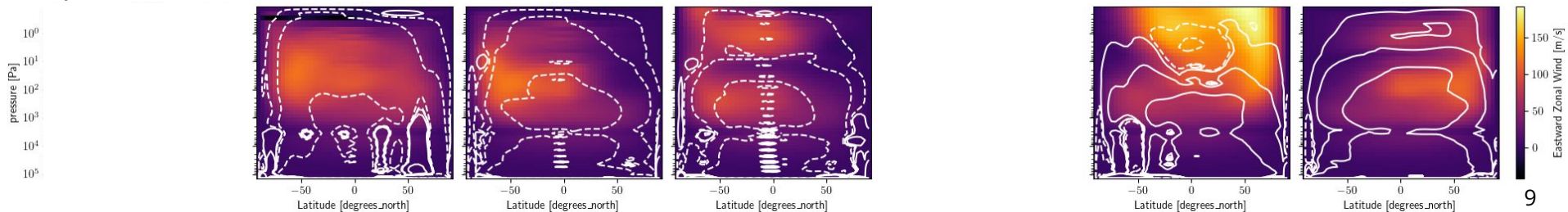
Northern winter

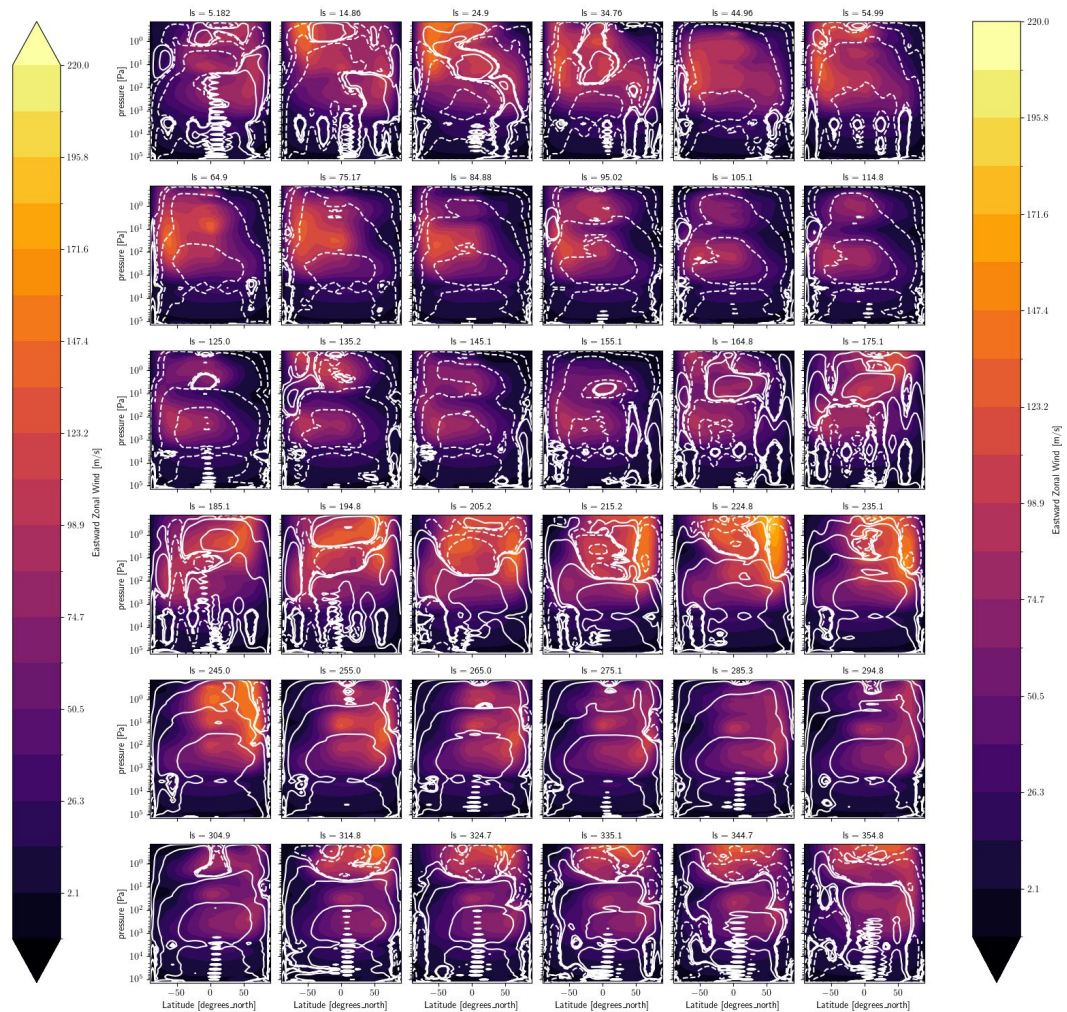
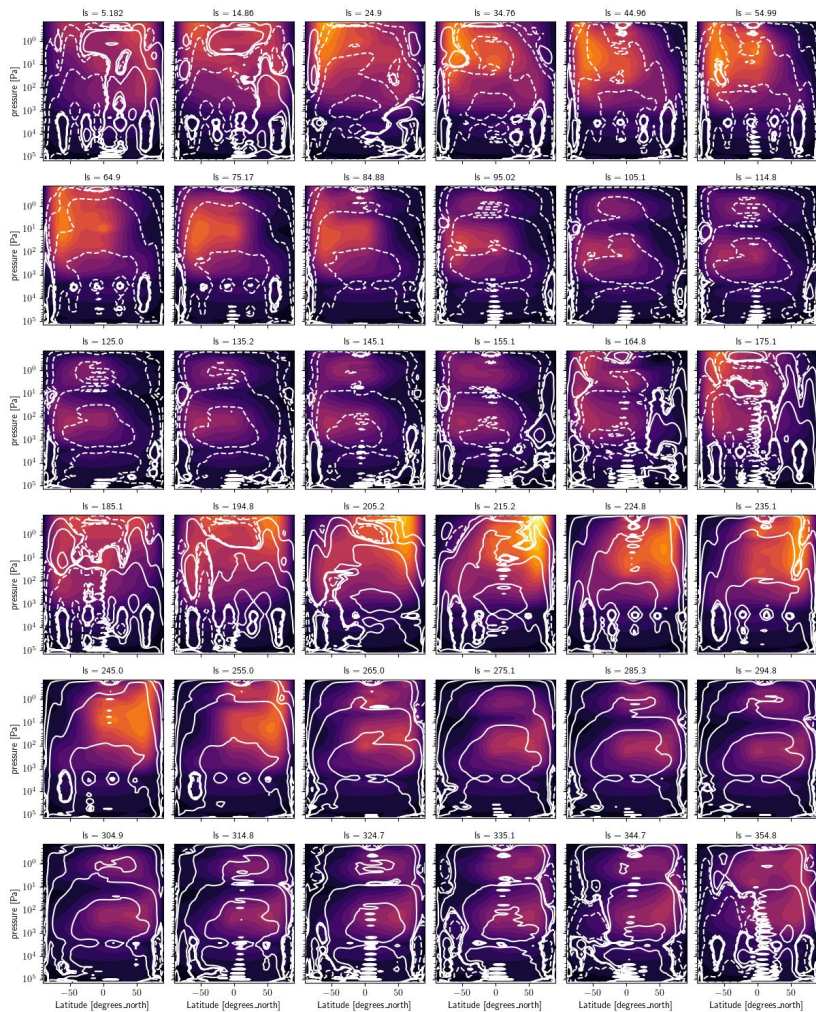


Reference (plane-parallel)



Coupled (3D radiative transfer)



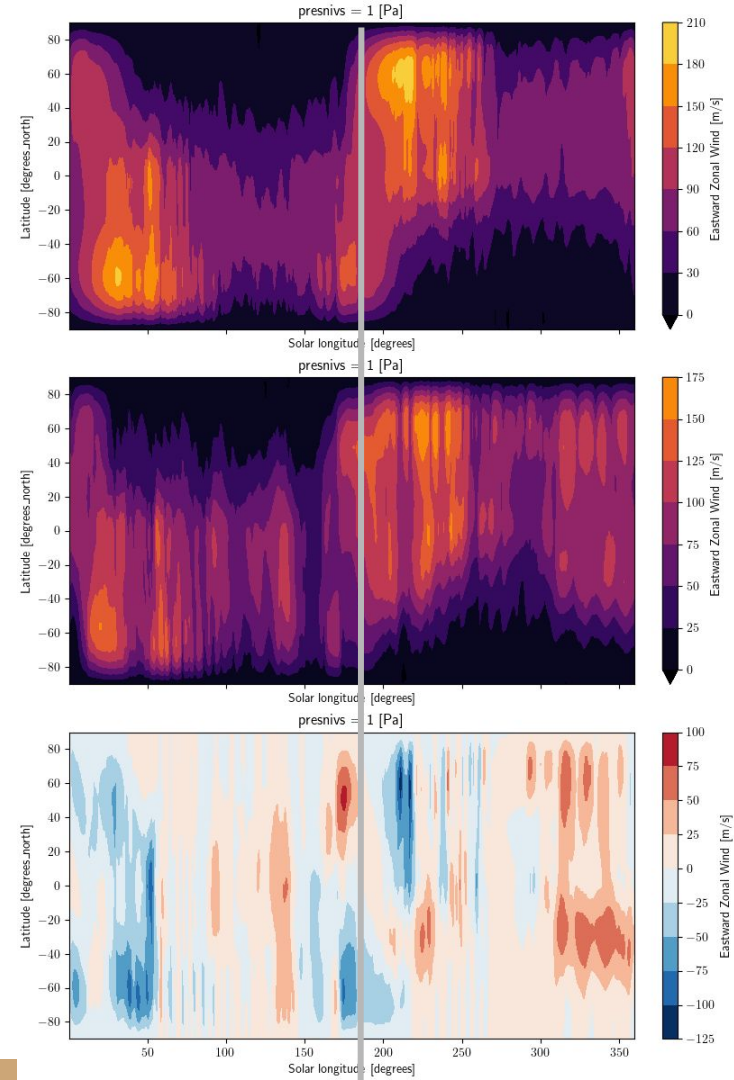


Preliminary results: circulation

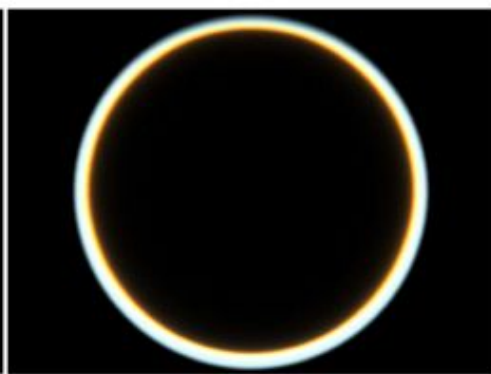
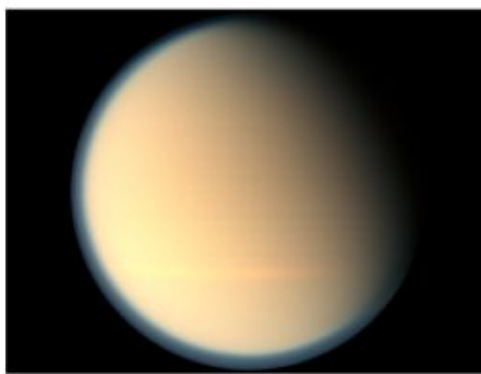
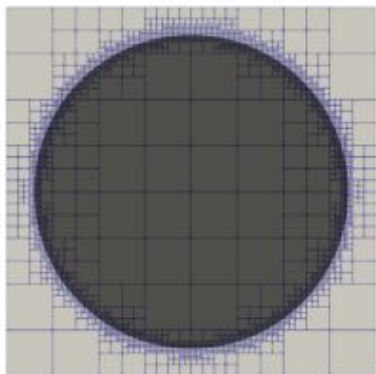
Changes in the structure of the zonal wind

More variability in strength

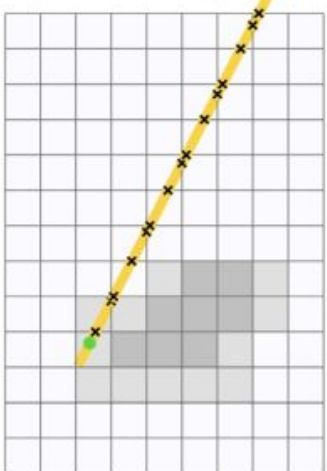
Inversion from south to north earlier



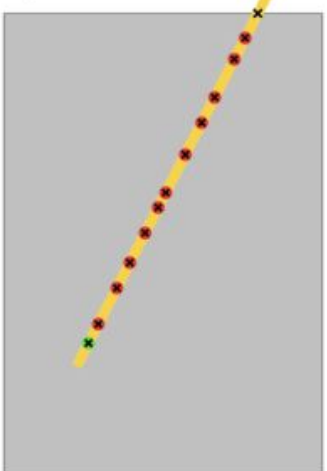
Nul-collision



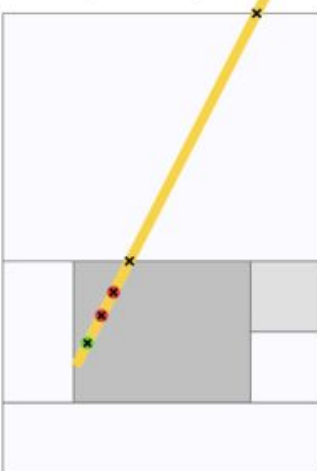
a) Path tracking



b) Null-collision



c) Null-collision + adaptative grid



× Access data ● True collision ● Null collision

He et al. (2025)

A control variates technique for the null-collision Monte Carlo algorithms

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^b*GSMA, UMR CNRS 6089, Université de Reims Champagne-Ardenne, Reims, France*

^c*Instituto de Astrofísica de Andalucía (IAA-CSIC), Glorieta de la Astronomía s/n, 18008 Granada, Spain*

^d*Laboratoire d'Optique Atmosphérique, CNRS, UMR 8518, Université de Lille, Lille, France*

Abstract

We propose a control variates technique to reduce the variance of null-collision Monte Carlo algorithms used for solving the Radiative Transfer Equation (RTE) in highly heterogeneous media. The method complements the classical spatially partitioned overestimate approach by additionally recording the minimum absorption coefficient within each voxel during preprocessing. During path tracing, the attenuation due to this minimum absorption is evaluated analytically, while the residual part is handled by path-samplings. This analytical treatment significantly improves convergence particularly in strongly absorbing media such as the planetary atmospheres in infrared absorbing band. The mathematical equivalence between the original and control-variates estimators is demonstrated, and numerical applications for Earth's and Titan's atmospheres confirm the expected variance reduction.

Keywords: Radiative Transfer, Monte Carlo, Control Variates, Null-collision

Traiter une partie du problème en analytique pour réduire la variance
Nouvel algo: une partie de l'absorption en collision-nulle est traité analytiquement

Rappel: collision nulle

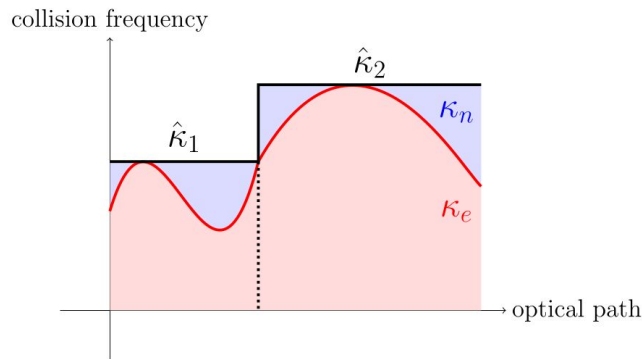
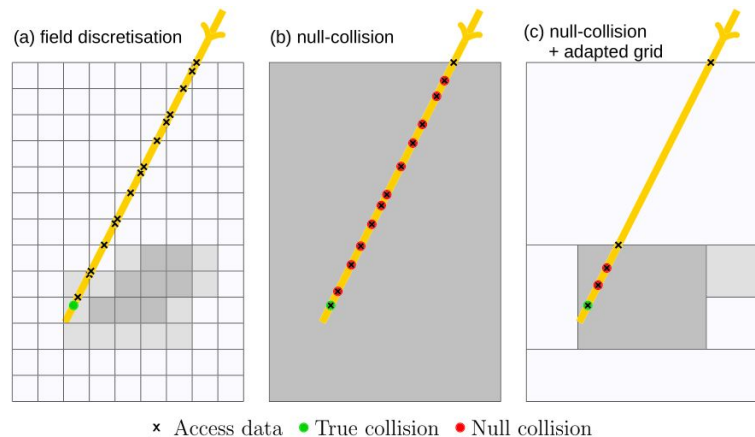


Figure 2: Example of an optical path traversing two voxels in the adapted grid. Fictitious colliders with collision frequency κ_n are added to the real, spatially varying extinction field κ_e to yield a constant overestimated extinction $\hat{\kappa}_1$ and $\hat{\kappa}_2$ within each voxels and stored during pre-processing.



Voxels homogènes: “cachent la variabilité du milieu”

Utilisation de majorant

Introduction de collision-nul

Control variates

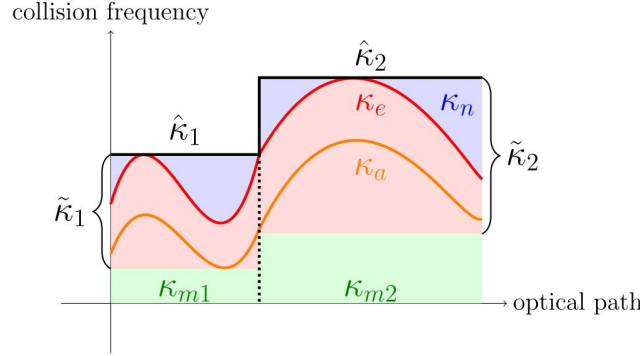


Figure 4: In addition to the maximum extinction coefficients $\hat{\kappa}_1$, $\hat{\kappa}_2$, the minimum absorption coefficients κ_{m1} and κ_{m2} are also stored during preprocessing. The reduced overestimates $\tilde{\kappa}_1 = \hat{\kappa}_1 - \kappa_{m1}$ and $\tilde{\kappa}_2 = \hat{\kappa}_2 - \kappa_{m2}$ are used to sample collision distances during path tracing.

$$w' = T_m * w + S_m$$

sum of expected values of Q_i , which are sampled within each traversed voxels (as illustrated in Fig. 6):

$$S_m(\vec{x}, \vec{x}') = \sum_{i=1}^n \mathbb{E}[Q_i], \quad (9)$$

with each contribution defined as:

$$\mathbb{E}[Q_i] = \prod_{j=1}^{i-1} \exp(-\kappa_{m_j} \sigma_j) \int_0^{\sigma_i} p_{\Sigma'}(\sigma') d\sigma' \cdot (1 - \exp(-\kappa_{m_i} \sigma_i)) \cdot B(\vec{x} - \vec{u}\sigma'), \quad (10)$$

On introduit un minorant de l'absorption

Cette partie (en verte) est traitée de façon analytique

Au cours du tracé, on conserve:

- La transmission le long du chemin lié au minorant

- L'émission atmosphérique le long du chemin lié au minorant

Une fois le rayon terminé, le poids est atténué par la transmission et augmenté de l'émission, liés au minorant

Here, $T_m(\vec{x}, \vec{x}')$ is the transmissivity due to κ_m between two positions, calculated analytically as:

$$T_m(\vec{x}, \vec{x}') = \prod_{i=1}^n \exp(-\kappa_{m_i} \sigma_i), \quad (7)$$

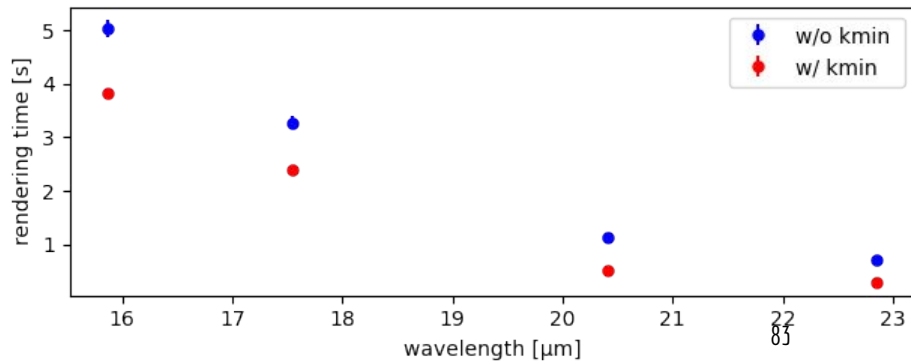
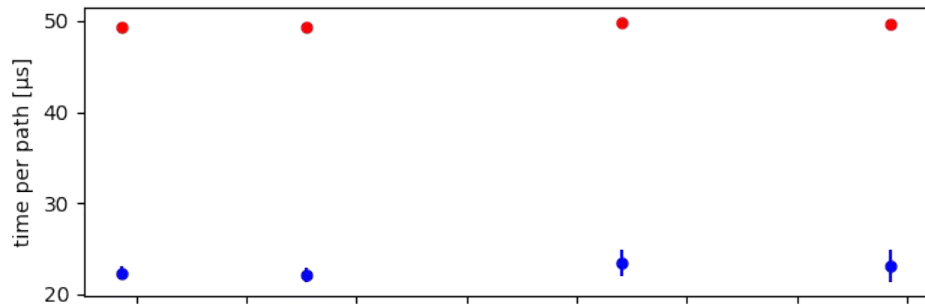
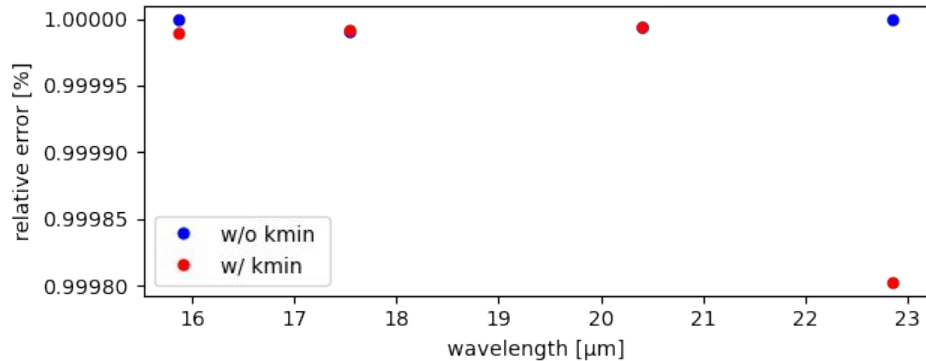
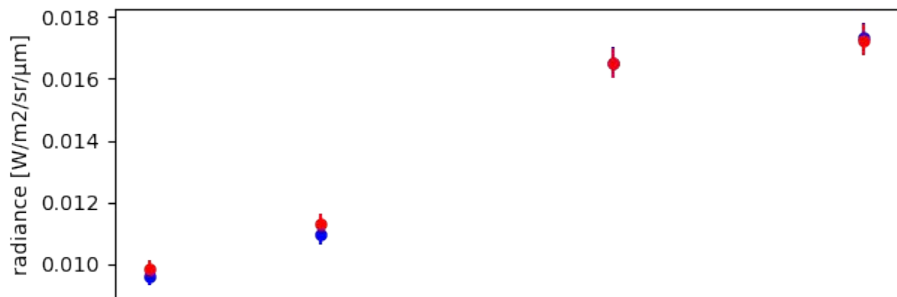
Résultats

Table B.1: Results of 1D-atmospheric slab configuration in $[W/m^2 \cdot sr \cdot \mu m]$. The number of realisation is 100000.

Solution	Earth 650cm ⁻¹	Titan 650cm ⁻¹
$\kappa_m = 0$	1.94129	7.441×10^{-3}
$\kappa_m = 0.25\kappa_a$	$1.94148 \pm 6.06 \times 10^{-4}$	$7.564 \times 10^{-3} \pm 9.17 \times 10^{-5}$
$\kappa_m = 0.5\kappa_a$	$1.94088 \pm 4.96 \times 10^{-4}$	$7.418 \times 10^{-3} \pm 7.07 \times 10^{-5}$
$\kappa_m = 0.75\kappa_a$	$1.94116 \pm 3.88 \times 10^{-4}$	$7.533 \times 10^{-3} \pm 5.95 \times 10^{-5}$
$\kappa_m = \kappa_a$	$1.94111 \pm 2.71 \times 10^{-4}$	$7.453 \times 10^{-3} \pm 3.61 \times 10^{-5}$
	1.94129 ± 0	$7.441 \times 10^{-3} \pm 0$

Table B.2: Results of the full-complexity configuration in $[W/m^2 \cdot sr \cdot \mu m]$. The number of realisation is 100000

	Titan 650cm ⁻¹	Titan 579cm ⁻¹
reference	$9.87 \times 10^{-3} \pm 2.96 \times 10^{-4}$	$11.04 \times 10^{-3} \pm 2.67 \times 10^{-4}$
our approach	$9.66 \times 10^{-3} \pm 1.59 \times 10^{-4}$	$11.38 \times 10^{-3} \pm 1.61 \times 10^{-4}$
	Titan 490cm ⁻¹	Titan 437cm ⁻¹
reference	$16.54 \times 10^{-3} \pm 2.24 \times 10^{-4}$	$17.43 \times 10^{-3} \pm 1.96 \times 10^{-4}$
our approach	$16.71 \times 10^{-3} \pm 1.32 \times 10^{-4}$	$17.52 \times 10^{-3} \pm 1.13 \times 10^{-4}$



Limitations

Calcul intermédiaire coûteux

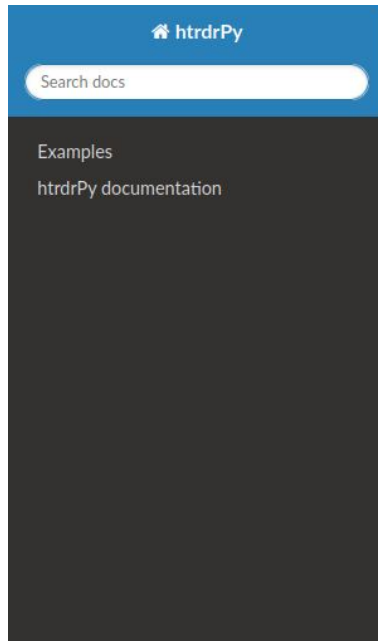
==> nécessite un cas optimal:

- Forte absorption, peu de diffusion
- Faible variabilité au sein du voxel

==> problème avec les grilles actuelles de htrdr-planets

Amélioration avec les futures grilles ?

htrdrPy



htrdrPy

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htrdrPy

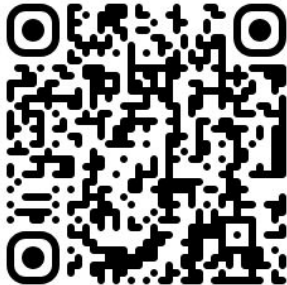
Introduction

htrdrPy is a wrapper for htrdr-planets (<https://www.meso-star.com/projects/htrdr/htrdr.html>). htrdr-planets is not included in the wrapper, it must be installed separately. htrdrPy is meant to simplify the use, by managing the input data, scripts, observation geometry and treatment of the results.

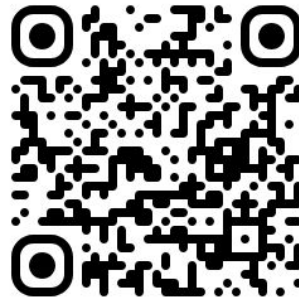
Installation

htrdrPy can be installed from the Python Package Index with the following command:

```
$ pip install htrdrPy
```



Documentation



GitLab



PyPi