

# What would be the magnetic field of the Earth if it spun as slowly as Venus?

Henri-Claude Nataf<sup>1</sup>, Nathanaël Schaeffer<sup>1</sup>, Quentin Noraz<sup>2</sup>,  
Allan Sacha Brun<sup>3</sup> & Antoine Strugarek<sup>3</sup>

<sup>1</sup>ISTerre, Univ. Grenoble Alpes, CNRS, Grenoble

<sup>2</sup>KU Leuven, Leuven, Belgique

<sup>3</sup>AIM, CEA, Univ. Paris-Saclay

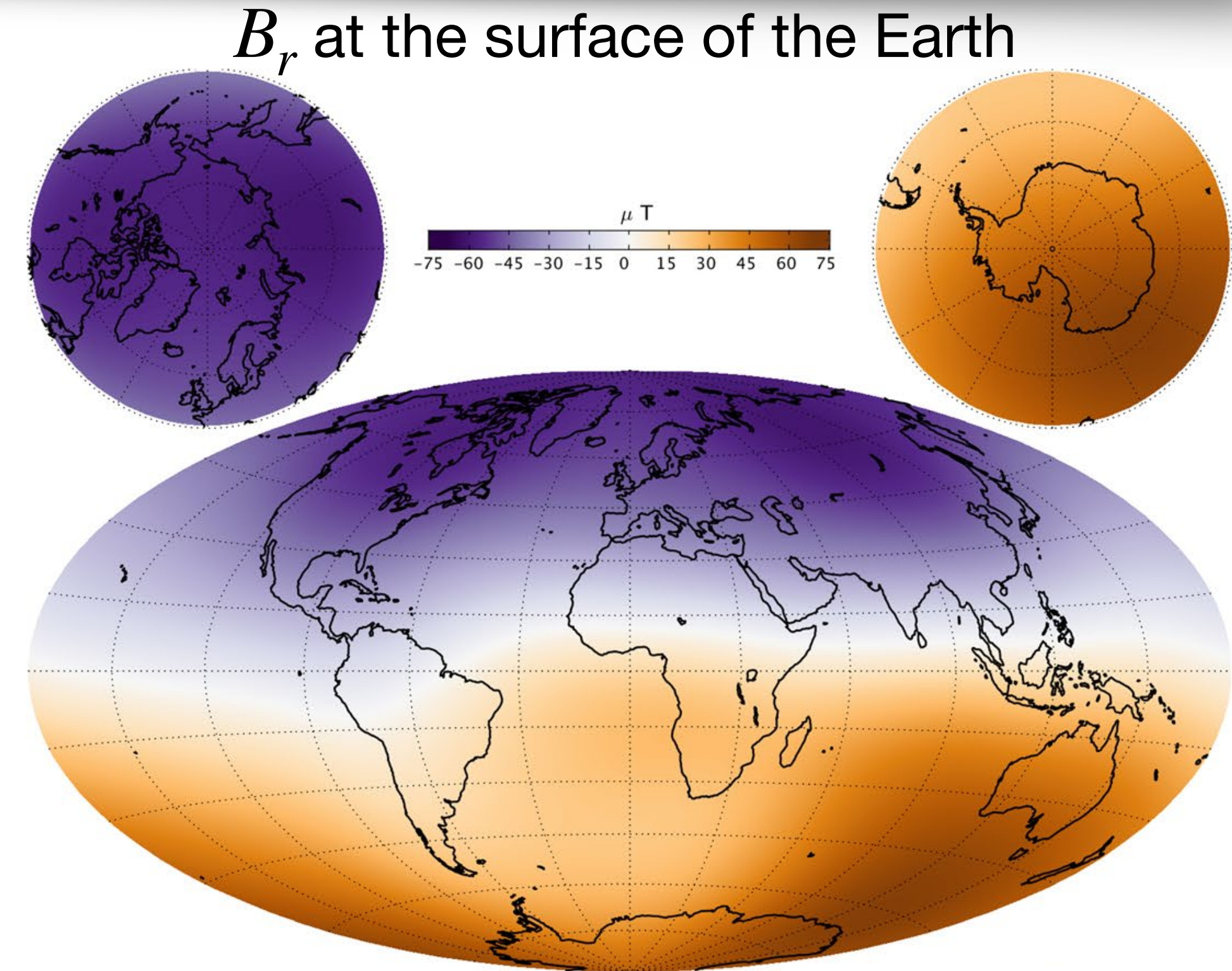
Journées SF2A, June 23<sup>rd</sup> 2026



# The magnetic field of the Earth

# The magnetic field of the Earth

- $B \sim 0.5$  Gauss = 0.05mT at the Earth surface, 10 times larger at the top of the core, dominated by a **dipole almost aligned with Earth's spin axis.**

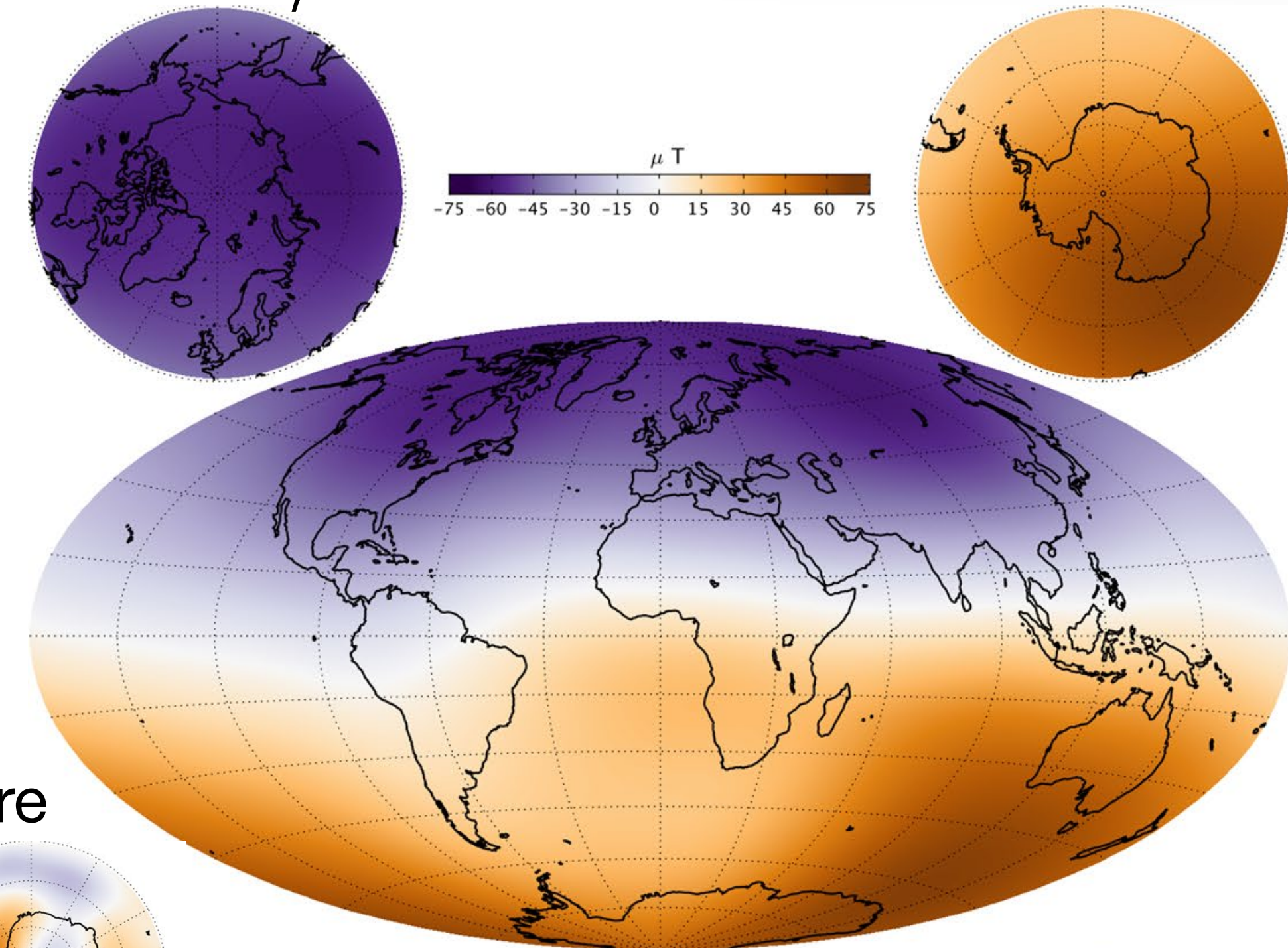


Finlay+ 2020

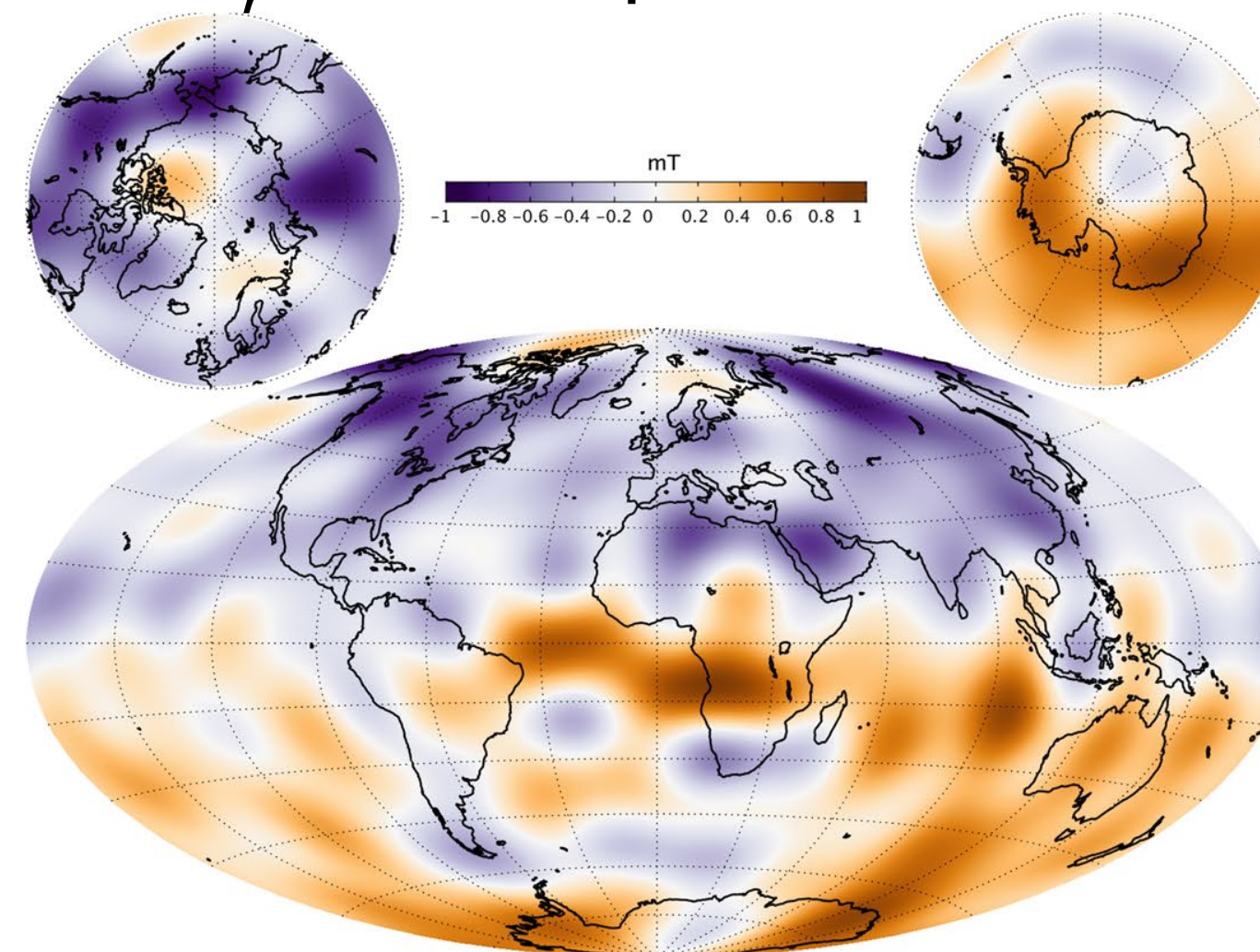
# The magnetic field of the Earth

- $B \sim 0.5$  Gauss = 0.05mT at the Earth surface, 10 times larger at the top of the core, dominated by a **dipole almost aligned with Earth's spin axis.**

$B_r$  at the surface of the Earth



$B_r$  at the top of the core

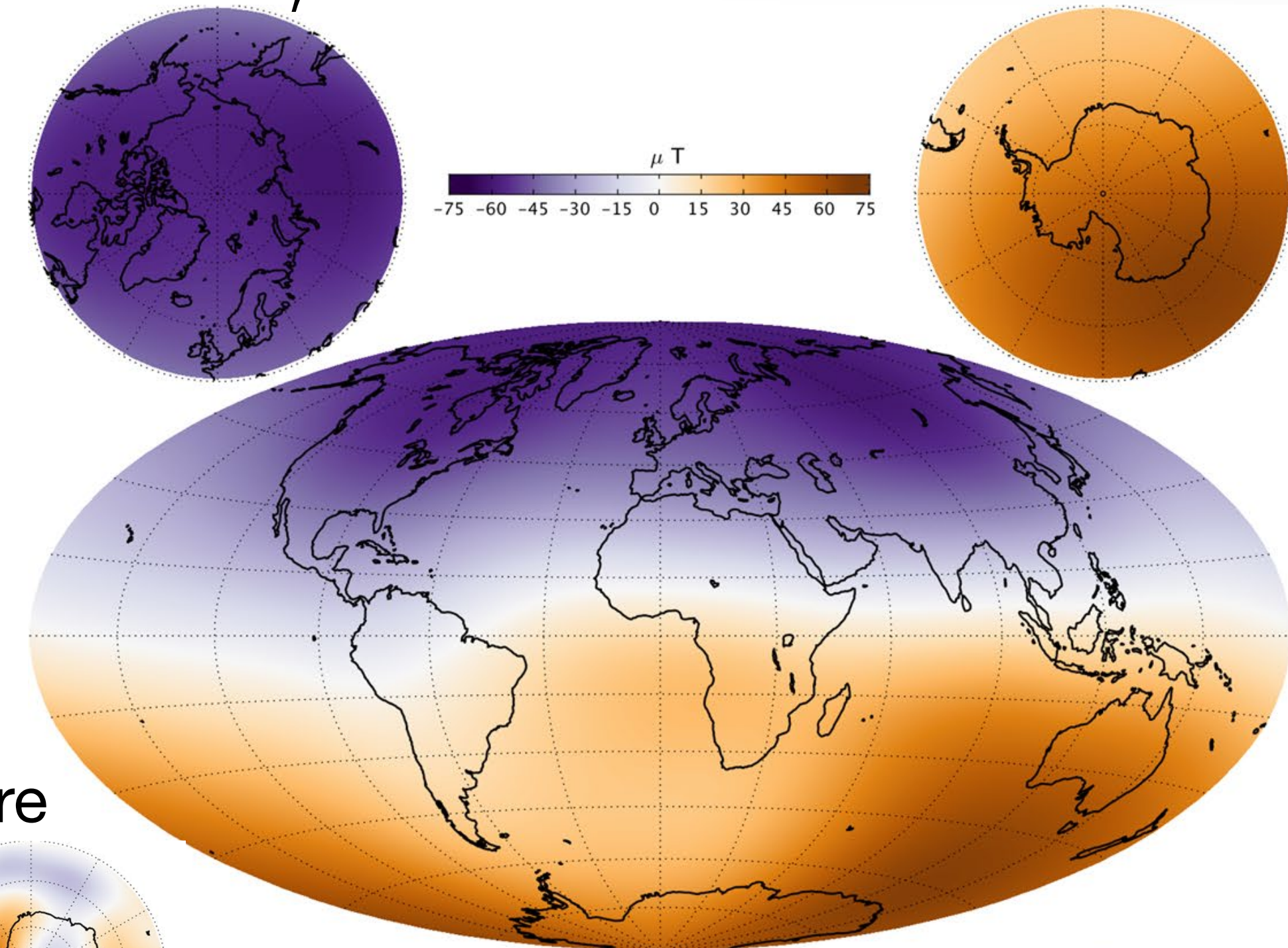


Finlay+ 2020

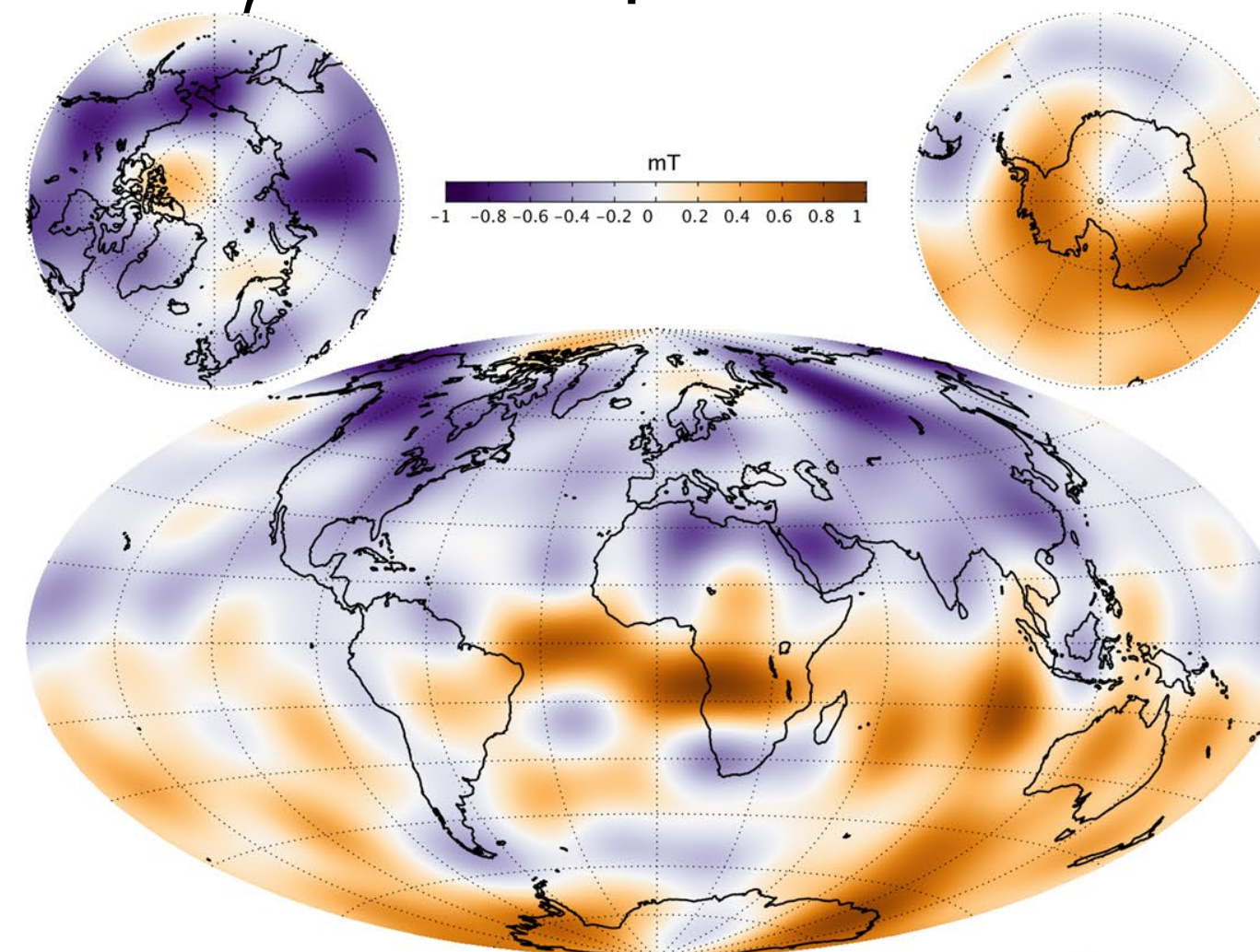
# The magnetic field of the Earth

- $B \sim 0.5$  Gauss = 0.05mT at the Earth surface, 10 times larger at the top of the core, dominated by a **dipole almost aligned with Earth's spin axis.**
- Produced by a dynamo powered by thermo-compositional convection  $\mathcal{P} \simeq 3$  TW.

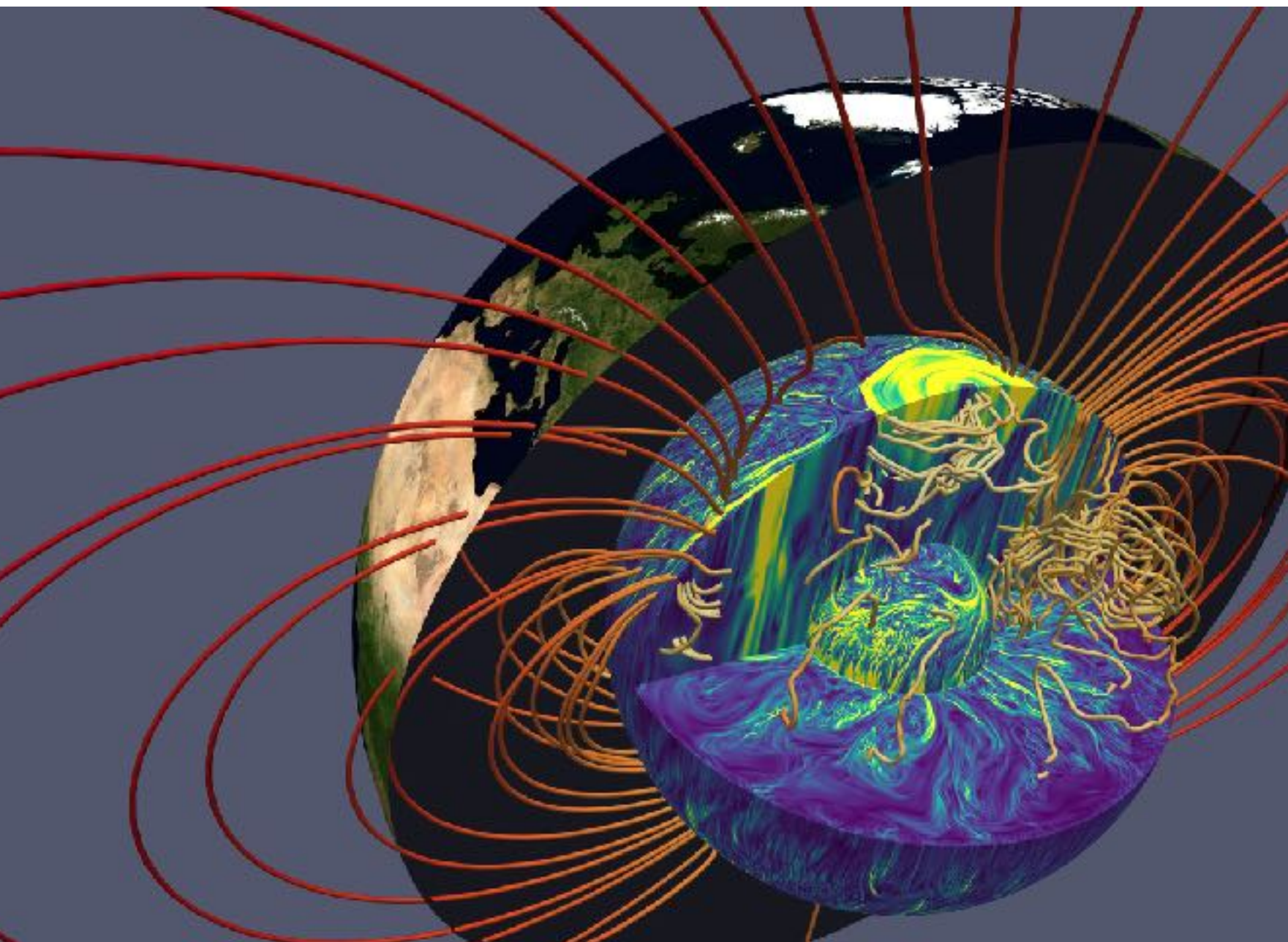
$B_r$  at the surface of the Earth



$B_r$  at the top of the core



Finlay+ 2020



Schaeffer+ 2017

$$t_{\Omega} = 1 \text{ day}$$

$$\mathcal{P} \simeq 3 \text{ TW}$$

$$B \simeq 3 \text{ mT}$$

$$U \simeq 0.5 \text{ mm/s}$$

~flat spectra for  $\ell \leq 10$

Roberts+ 2013; Gillet+2015

$$t_{\Omega} = 1 \text{ day}$$

$$\mathcal{P} \simeq 3 \text{ TW}$$

$$\left. \begin{array}{l} B \simeq 3 \text{ mT} \\ U \simeq 0.5 \text{ mm/s} \end{array} \right\} \Longrightarrow \frac{E_B}{E_U} \sim 5\,000$$

~flat spectra for  $\ell \leq 10$

Roberts+ 2013; Gillet+2015

$$t_{\Omega} = 1 \text{ day}$$

$$\mathcal{P} \simeq 3 \text{ TW}$$

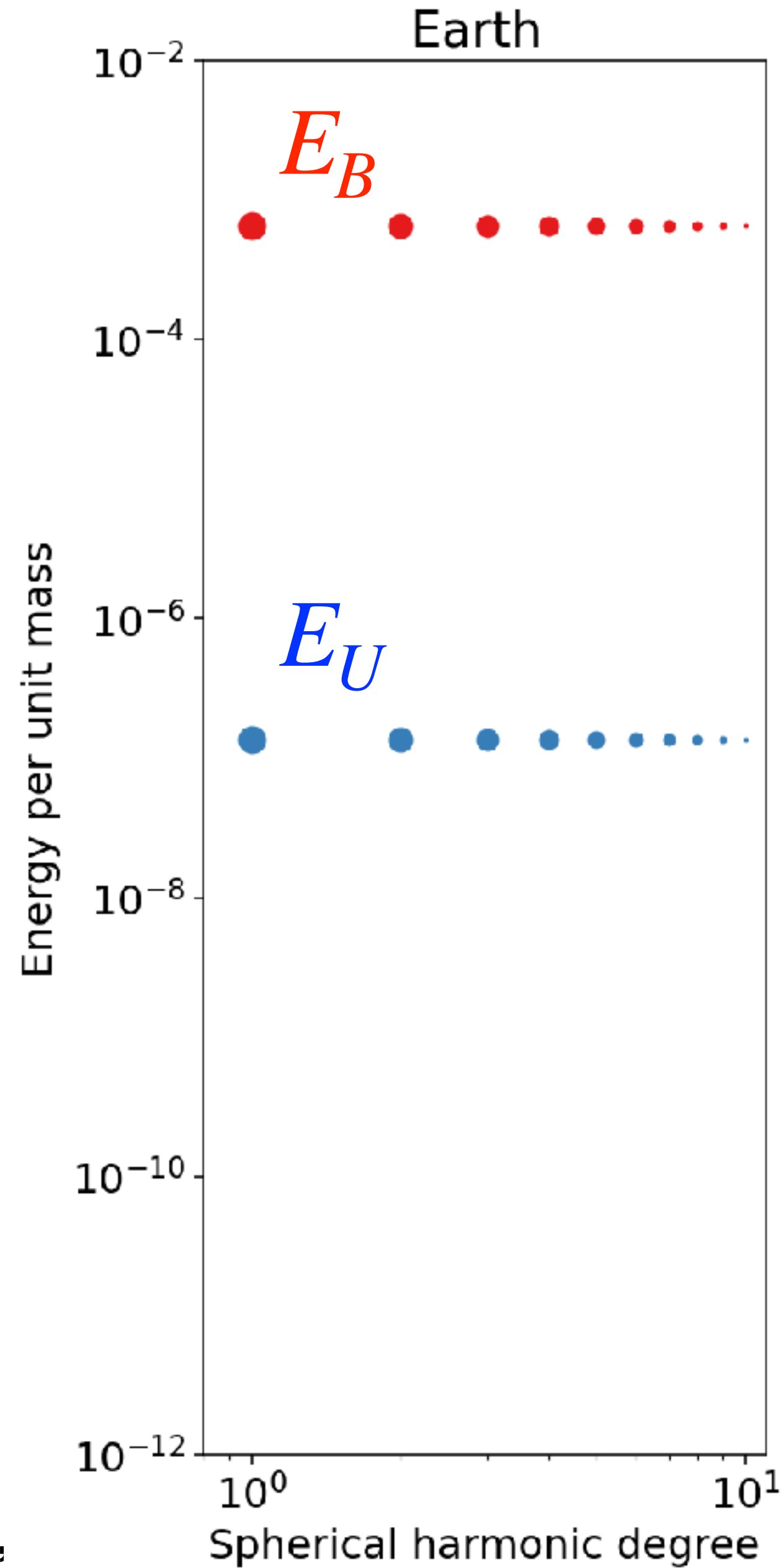
$$B \simeq 3 \text{ mT}$$

$$U \simeq 0.5 \text{ mm/s}$$

$$\left. \begin{array}{l} B \simeq 3 \text{ mT} \\ U \simeq 0.5 \text{ mm/s} \end{array} \right\} \Rightarrow \frac{E_B}{E_U} \sim 5000$$

~flat spectra for  $\ell \leq 10$

Roberts+ 2013; Gillet+2015



$$t_{\Omega} = 1 \text{ day}$$

$$\mathcal{P} \simeq 3 \text{ TW}$$

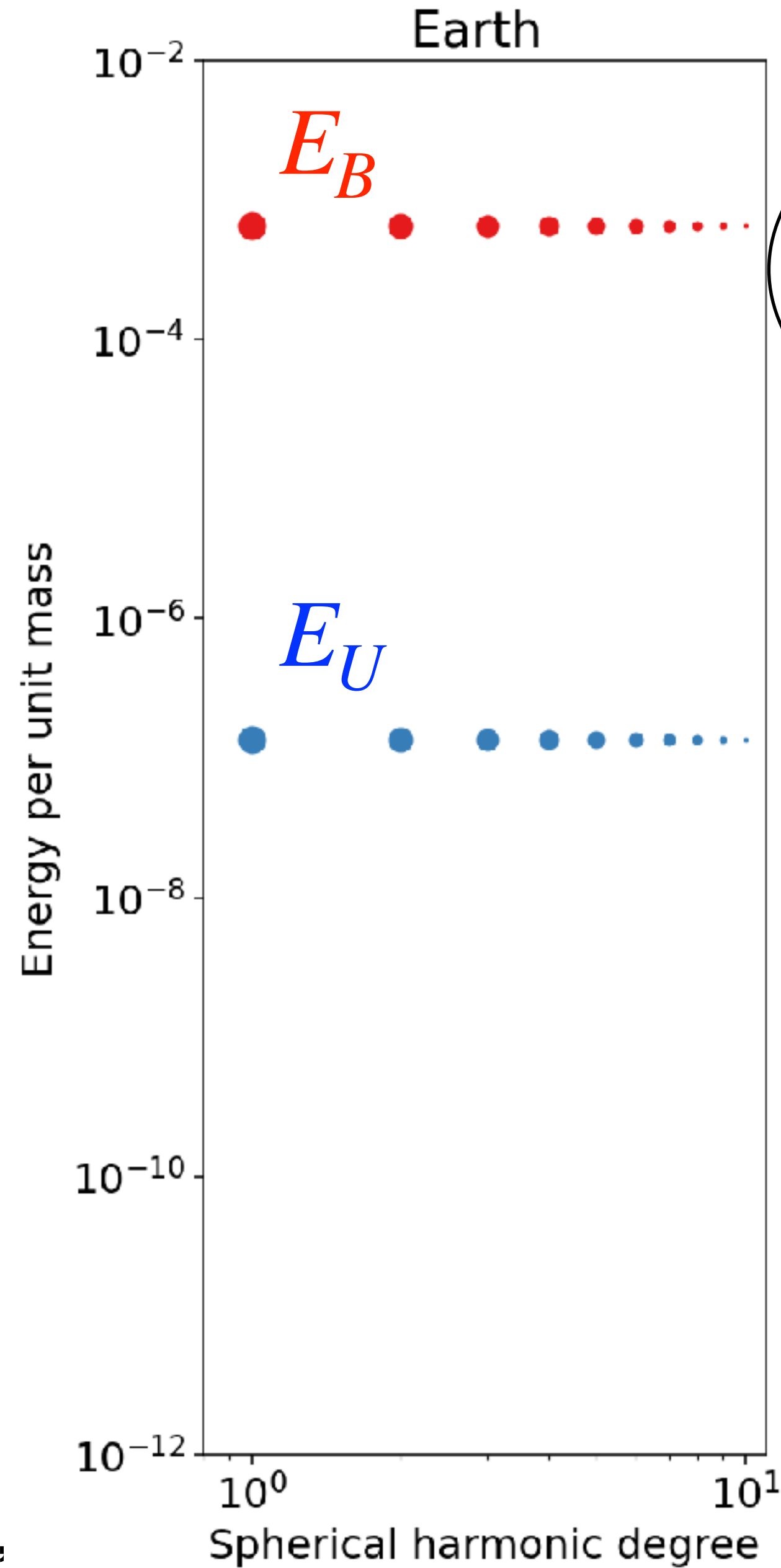
$$B \simeq 3 \text{ mT}$$

$$U \simeq 0.5 \text{ mm/s}$$

$$\left. \begin{array}{l} B \simeq 3 \text{ mT} \\ U \simeq 0.5 \text{ mm/s} \end{array} \right\} \Rightarrow \frac{E_B}{E_U} \sim 5000$$

~flat spectra for  $\ell \leq 10$

Roberts+ 2013; Gillet+2015



How would they change if the Earth spun as slowly as Venus?

$t_{\Omega} = 243 \text{ days}$

# The classical answer is...

**Not much!**

Both are « rapidly rotating dynamos » since their **Ekman** number is very small

$$Ek < 10^{-12}$$

$$Ek = \frac{\nu}{\Omega R^2}$$

$\nu = 10^{-6} \text{ m}^2/\text{s}$	kinematic viscosity
$R = 3\,500 \text{ km}$	radius of the core

# Scaling laws for planetary dynamos

Indeed, the classical scaling law for **magnetic** intensity is *independent* of  $t_{\Omega}$

$$\frac{B}{\sqrt{\rho\mu}} \simeq R \left( \frac{\mathcal{P}}{MR^2} \right)^{1/3} \quad \text{Christensen 2010}$$

$\nu = 10^{-6} \text{ m}^2/\text{s}$	kinematic viscosity
$R = 3500 \text{ km}$	radius of the core
$M = 1.8 \times 10^{24} \text{ kg}$	mass of the core
$\mathcal{P} = 3 \text{ TW}$	convective power

# Scaling laws for planetary dynamos

Indeed, the classical scaling law for **magnetic** intensity is *independent* of  $t_{\Omega}$

$$\frac{B}{\sqrt{\rho\mu}} \simeq R \left( \frac{\mathcal{P}}{MR^2} \right)^{1/3} \quad \text{Christensen 2010}$$

However, the scaling law for **velocity** predicts a **~6 fold increase**

$$U \simeq \frac{R}{\Omega^{1/3}} \left( \frac{\mathcal{P}}{MR^2} \right)^{4/9} \quad \text{Davidson 2013}$$

which implies that the **Rossby** number would be **~1500 times larger**.

$$Ro = \frac{U}{\Omega R}$$

$\nu = 10^{-6} \text{ m}^2/\text{s}$	kinematic viscosity
$R = 3500 \text{ km}$	radius of the core
$M = 1.8 \times 10^{24} \text{ kg}$	mass of the core
$\mathcal{P} = 3 \text{ TW}$	convective power

# Scaling laws for planetary dynamos

Indeed, the classical scaling law for **magnetic** intensity is *independent* of  $t_{\Omega}$

$$\frac{B}{\sqrt{\rho\mu}} \simeq R \left( \frac{\mathcal{P}}{MR^2} \right)^{1/3} \quad \text{Christensen 2010}$$

However, the scaling law for **velocity** predicts a **~6 fold increase**

$$U \simeq \frac{R}{\Omega^{1/3}} \left( \frac{\mathcal{P}}{MR^2} \right)^{4/9} \quad \text{Davidson 2013}$$

which implies that the **Rossby** number would be **~1500 times larger**.

$$Ro = \frac{U}{\Omega R}$$

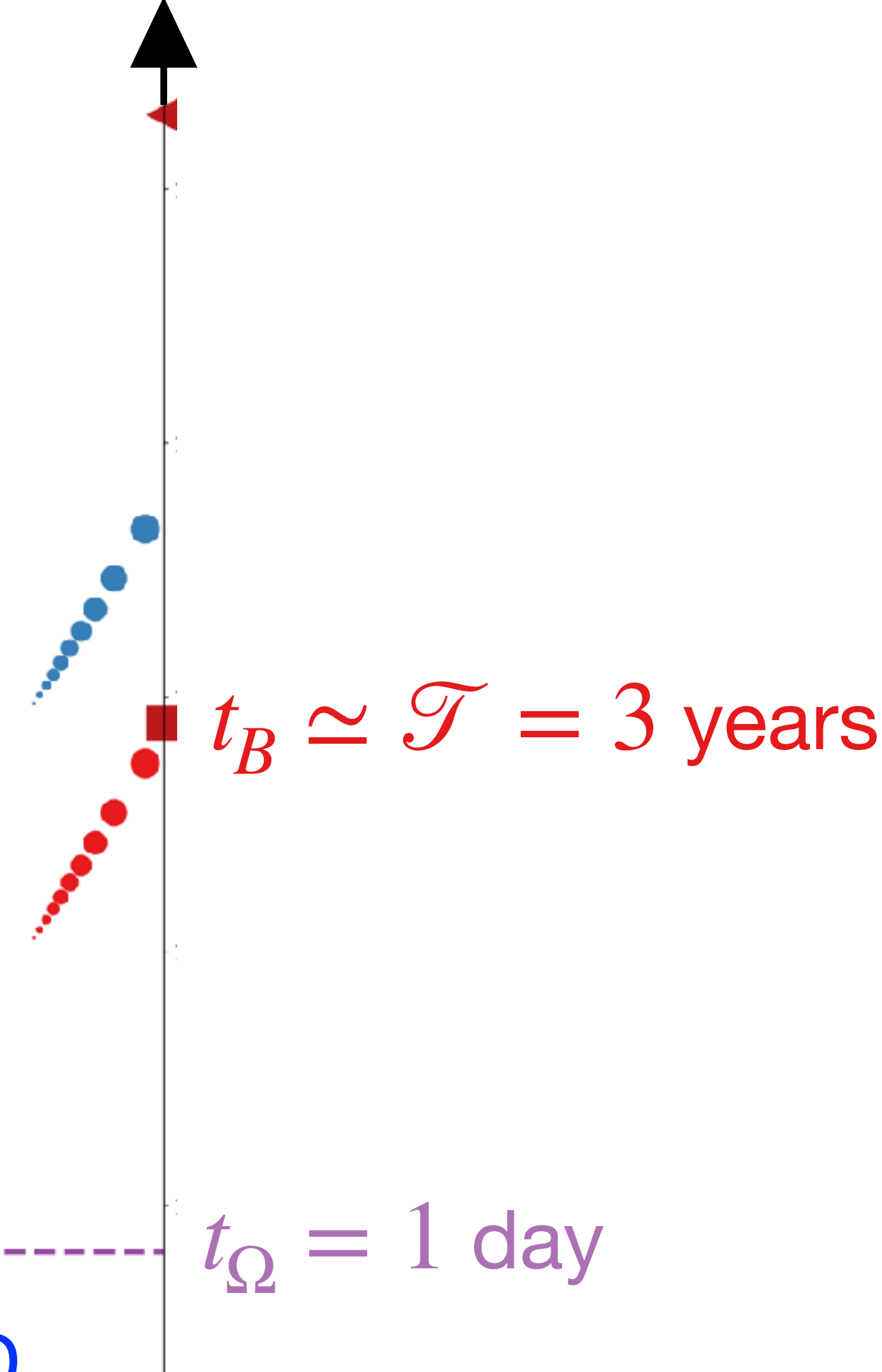
## force balance

These laws apply to dynamos in the QG-MAC regime (Quasi-Geostrophic Magneto-Archimedean-Coriolis)

$\nu = 10^{-6} \text{ m}^2/\text{s}$	kinematic viscosity
$R = 3500 \text{ km}$	radius of the core
$M = 1.8 \times 10^{24} \text{ kg}$	mass of the core
$\mathcal{P} = 3 \text{ TW}$	convective power

# QG-MAC characteristic times for the Earth

time (log scale)



$$t_B = \frac{R\sqrt{\rho\mu}}{B}$$

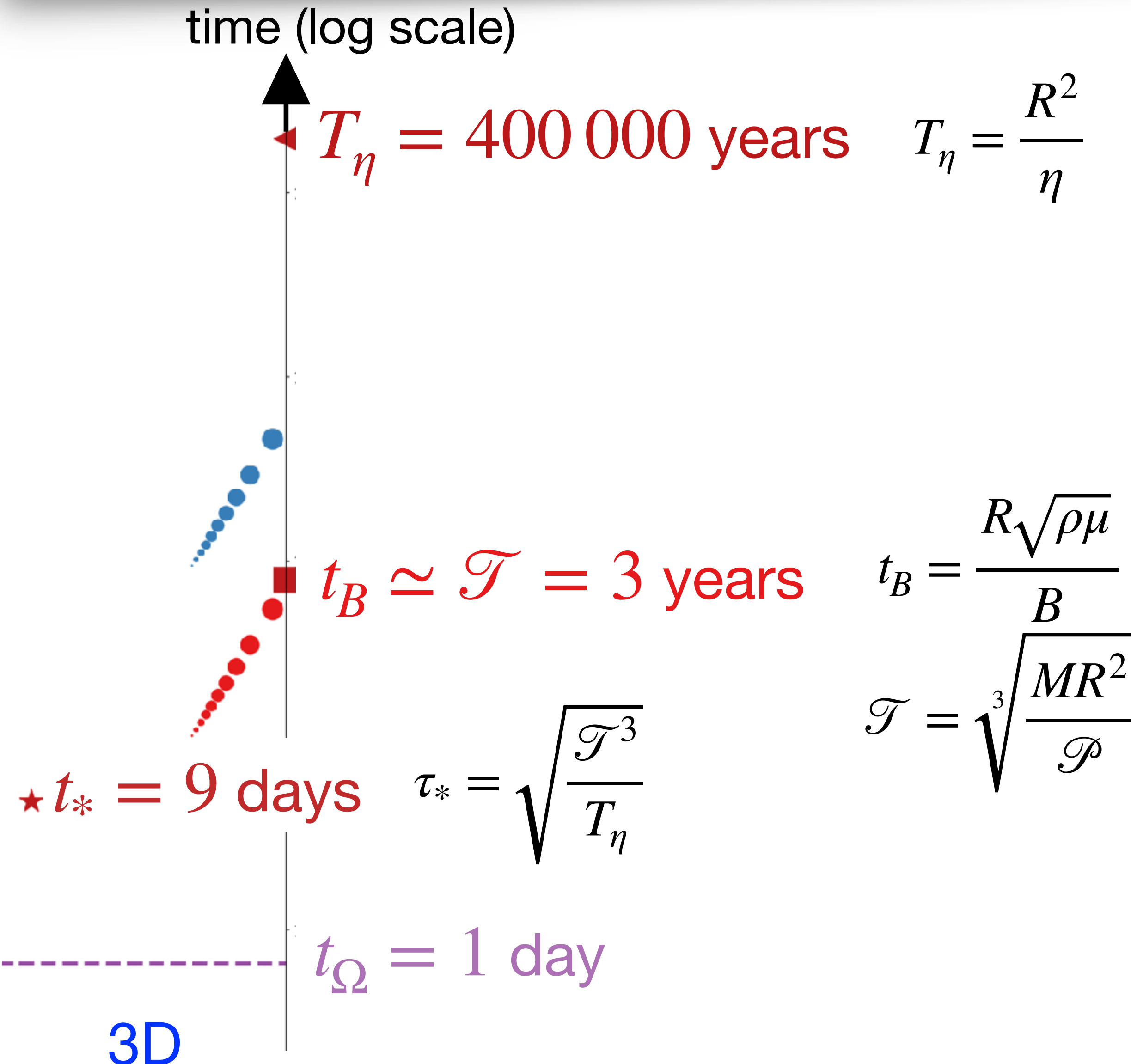
$$\mathcal{T} = \sqrt[3]{\frac{MR^2}{\mathcal{P}}}$$

A ‘**power**’ time  $\mathcal{T}$  is obtained from the power  $\mathcal{P}$  and we recover **magnetic time**  $t_B \simeq \mathcal{T}$  in this QG-MAC scenario.

$\nu = 10^{-6} \text{ m}^2/\text{s}$	kinematic viscosity
$R = 3500 \text{ km}$	radius of the core
$M = 1.8 \times 10^{24} \text{ kg}$	mass of the core
$\mathcal{P} = 3 \text{ TW}$	convective power
$\eta = 1 \text{ m}^2/\text{s}$	magnetic diffusivity

3D

# QG-MAC characteristic times for the Earth

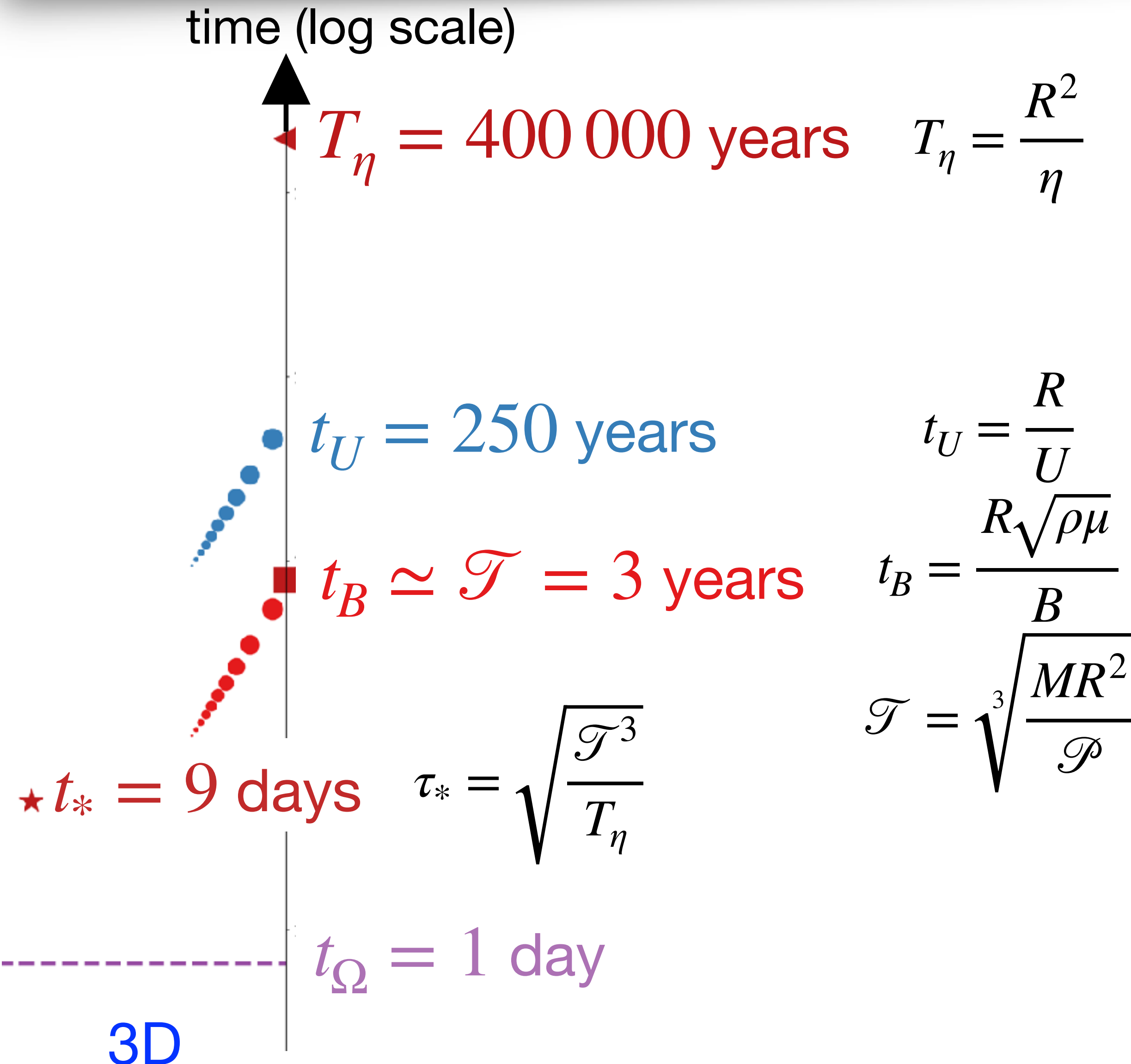


A ‘**power**’ time  $\mathcal{T}$  is obtained from the power  $\mathcal{P}$  and we recover **magnetic time**  $t_B \simeq \mathcal{T}$  in this QG-MAC scenario.

However, Ohmic dissipation occurs at small scales yielding **dissipation time**  $t_*$ .

$\nu = 10^{-6}$ m <sup>2</sup> /s	kinematic viscosity
$R = 3\,500$ km	radius of the core
$M = 1.8 \times 10^{24}$ kg	mass of the core
$\mathcal{P} = 3$ TW	convective power
$\eta = 1$ m <sup>2</sup> /s	magnetic diffusivity

# QG-MAC characteristic times for the Earth



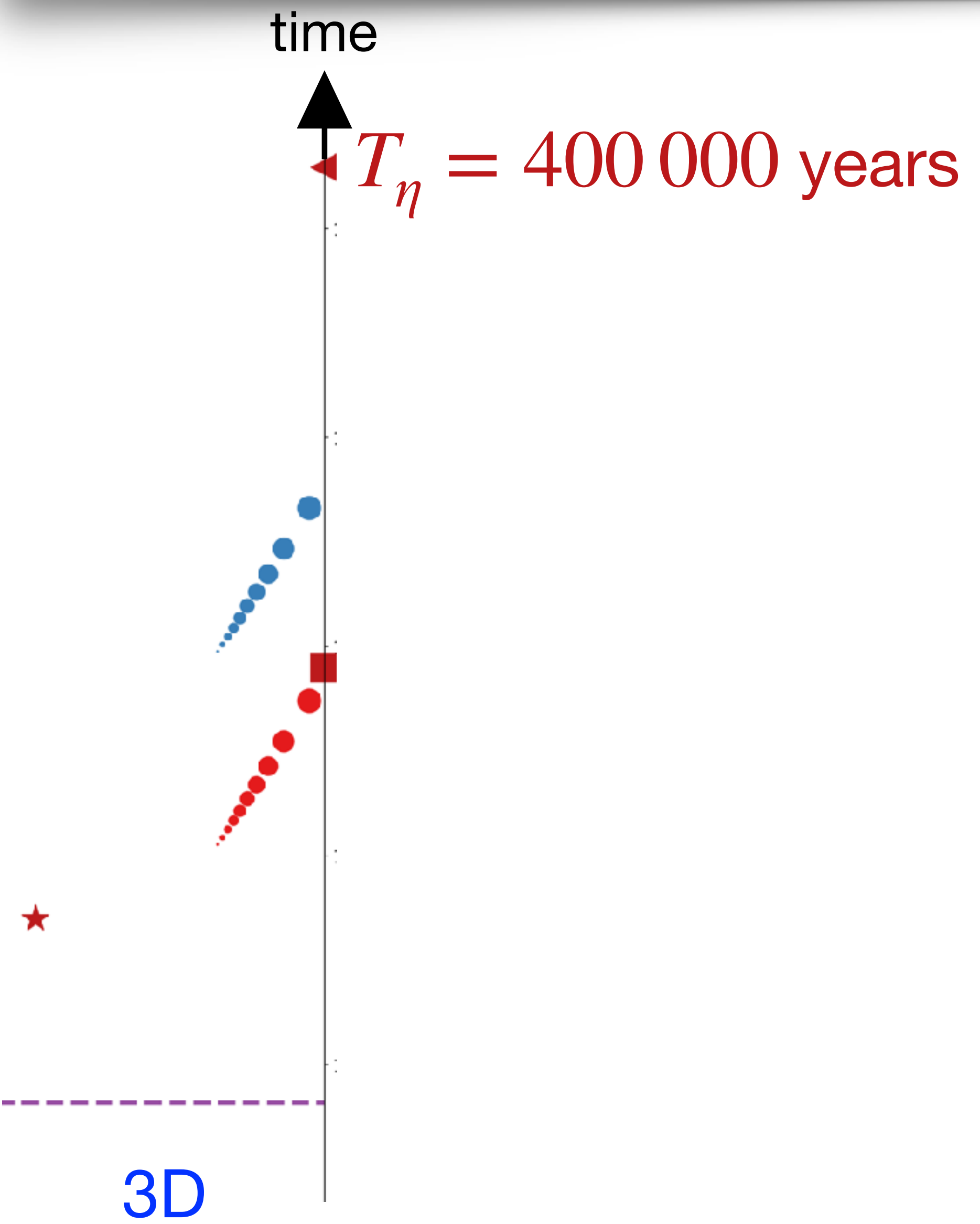
A ‘**power**’ time  $\mathcal{T}$  is obtained from the power  $\mathcal{P}$  and we recover **magnetic time**  $t_B \simeq \mathcal{T}$  in this QG-MAC scenario.

However, Ohmic dissipation occurs at small scales yielding **dissipation time**  $t_*$ .

With  $t_*$  **larger** than  $t_\Omega$  one gets  $t_U \gg t_\Omega$ : the flow is **quasi-geostrophic**.

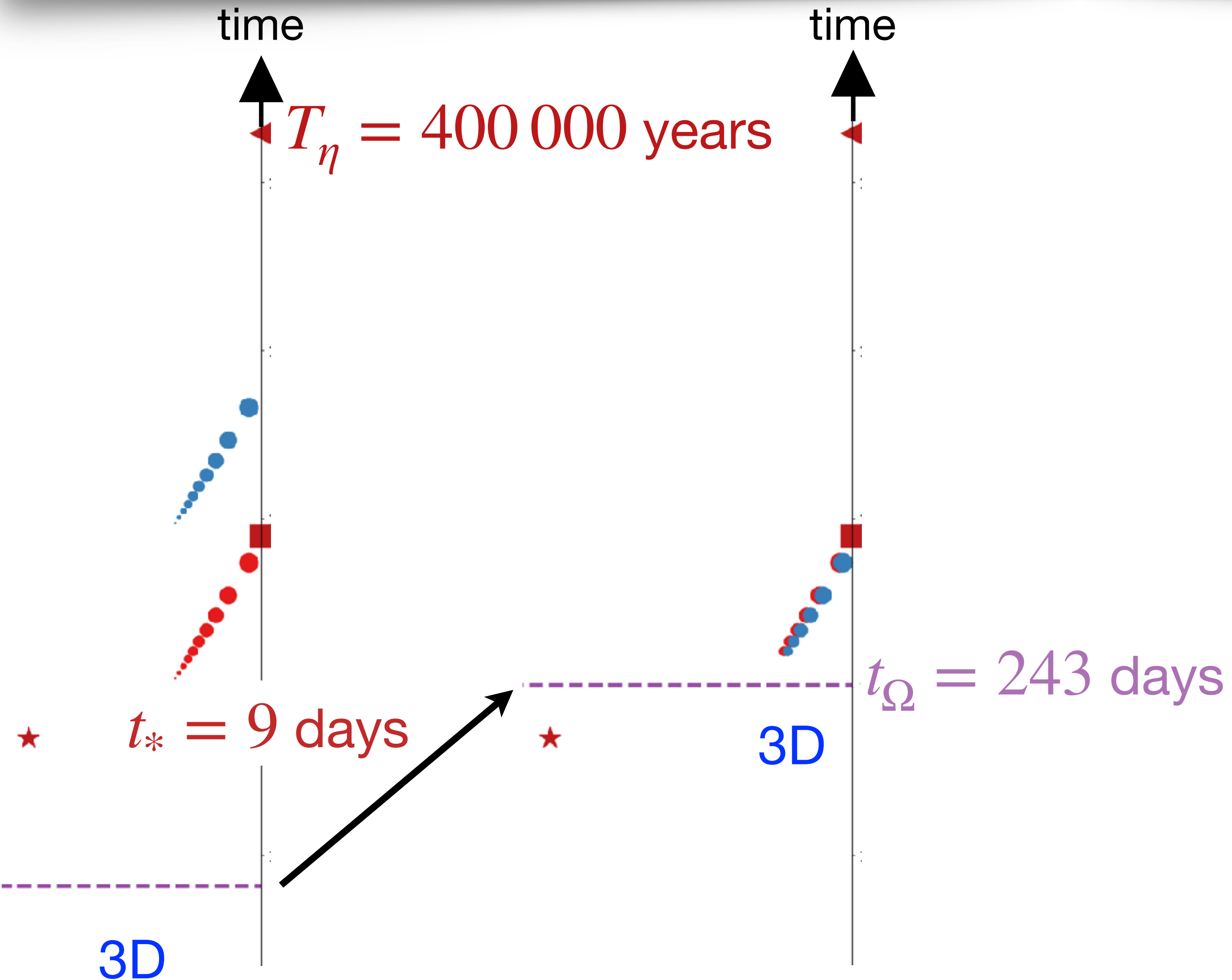
$\nu = 10^{-6}$ m <sup>2</sup> /s	kinematic viscosity
$R = 3\,500$ km	radius of the core
$M = 1.8 \times 10^{24}$ kg	mass of the core
$\mathcal{P} = 3$ TW	convective power
$\eta = 1$ m <sup>2</sup> /s	magnetic diffusivity

# What about Venus?



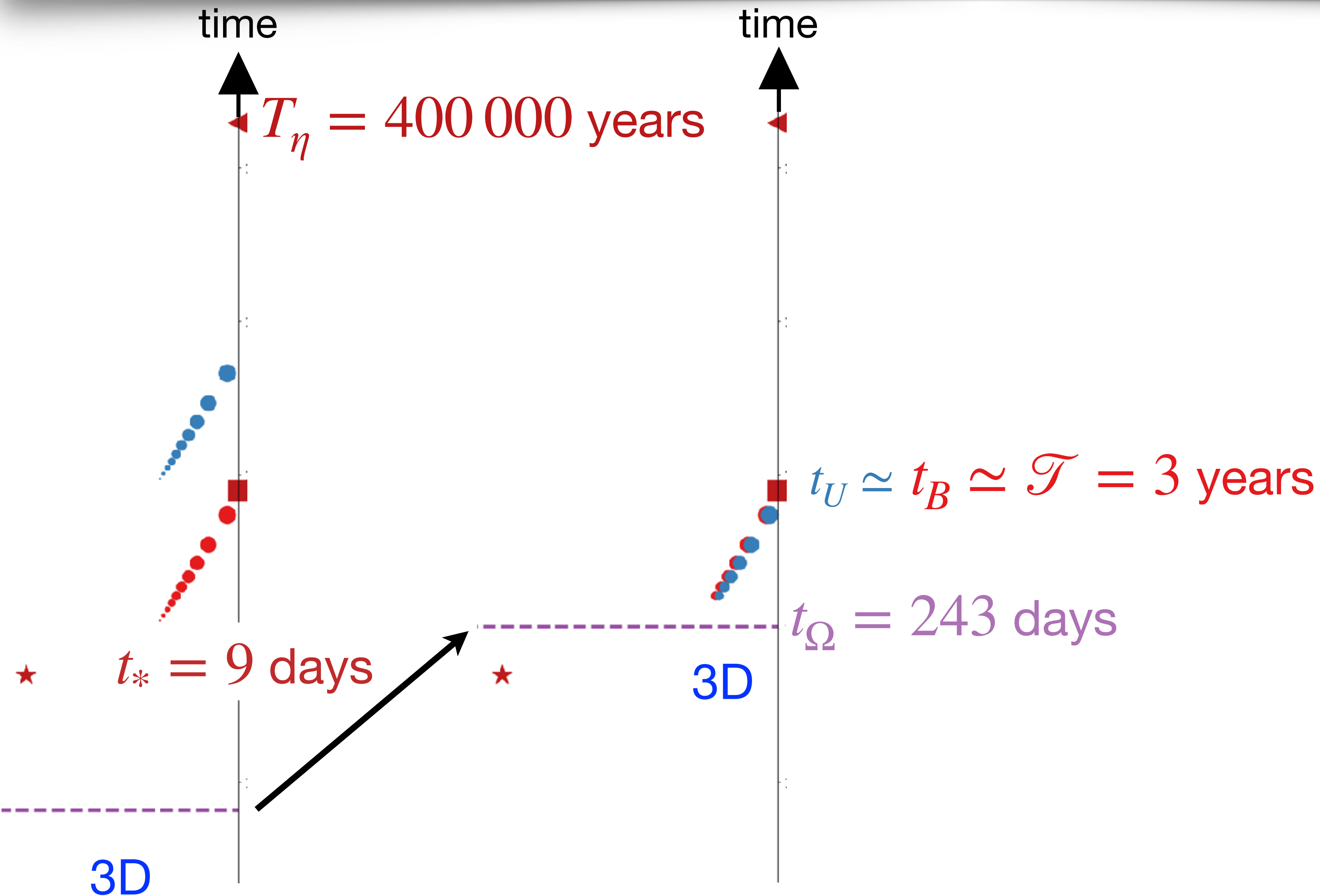
$\nu = 10^{-6}$ m <sup>2</sup> /s	kinematic viscosity
$R = 3\,500$ km	radius of the core
$M = 1.8 \times 10^{24}$ kg	mass of the core
$\mathcal{P} = 3$ TW	convective power
$\eta = 1$ m <sup>2</sup> /s	magnetic diffusivity

# What about Venus?



$\nu = 10^{-6}$ m <sup>2</sup> /s	kinematic viscosity
$R = 3\,500$ km	radius of the core
$M = 1.8 \times 10^{24}$ kg	mass of the core
$\mathcal{P} = 3$ TW	convective power
$\eta = 1$ m <sup>2</sup> /s	magnetic diffusivity

# What about Venus?

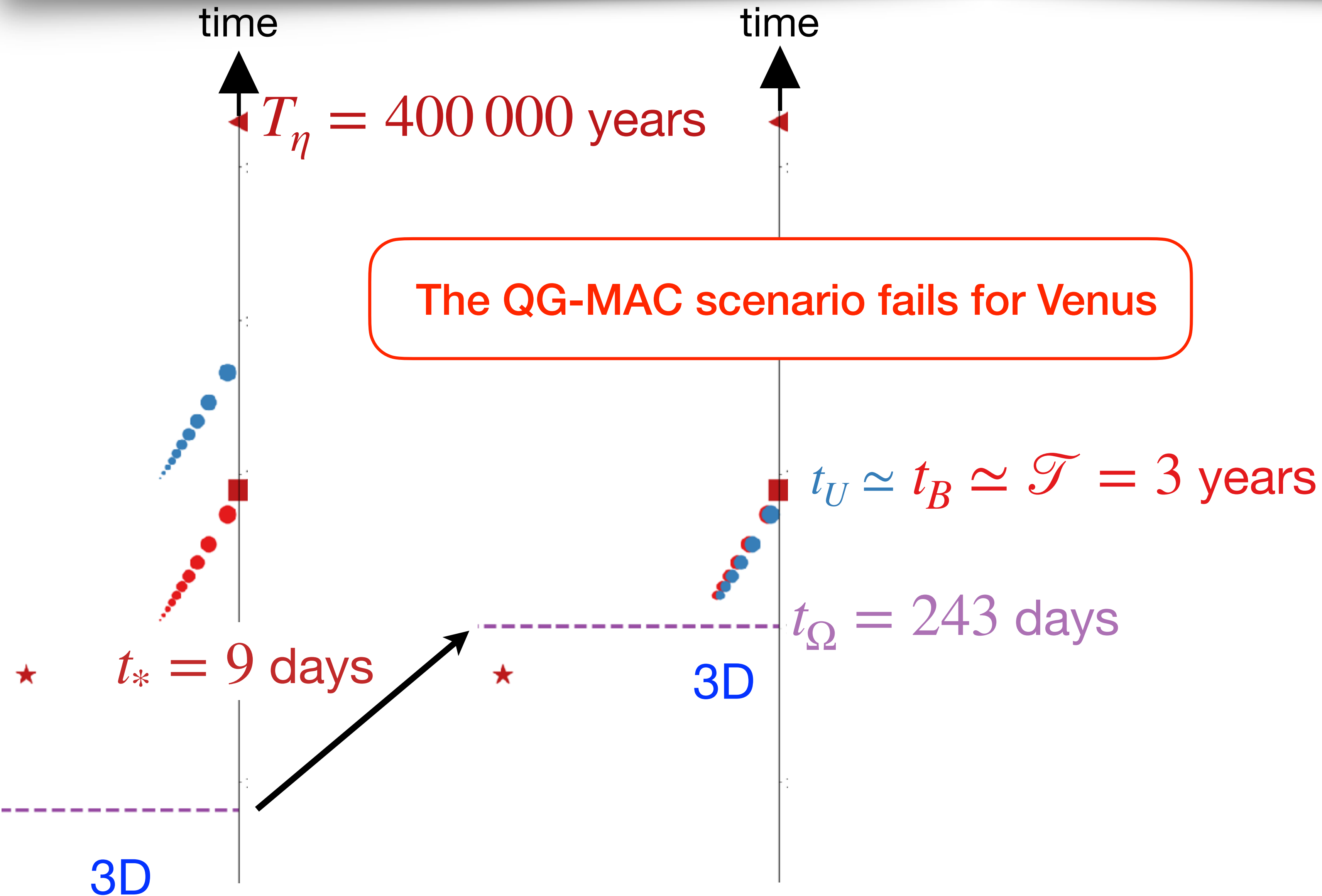


$t_*$  is now **smaller** than  $t_\Omega$ , and a QG-MAC scenario would yield a very rapid flow

$t_U \simeq t_B \sim t_\Omega$ , which would **not be** quasi-geostrophic anymore.

$\nu = 10^{-6}$ m <sup>2</sup> /s	kinematic viscosity
$R = 3\,500$ km	radius of the core
$M = 1.8 \times 10^{24}$ kg	mass of the core
$\mathcal{P} = 3$ TW	convective power
$\eta = 1$ m <sup>2</sup> /s	magnetic diffusivity

# What about Venus?



$t_*$  is now **smaller** than  $t_\Omega$ , and a QG-MAC scenario would yield a very rapid flow

$t_U \simeq t_B \sim t_\Omega$ , which would **not be** quasi-geostrophic anymore.

$\nu = 10^{-6}$ m <sup>2</sup> /s	kinematic viscosity
$R = 3\,500$ km	radius of the core
$M = 1.8 \times 10^{24}$ kg	mass of the core
$\mathcal{P} = 3$ TW	convective power
$\eta = 1$ m <sup>2</sup> /s	magnetic diffusivity

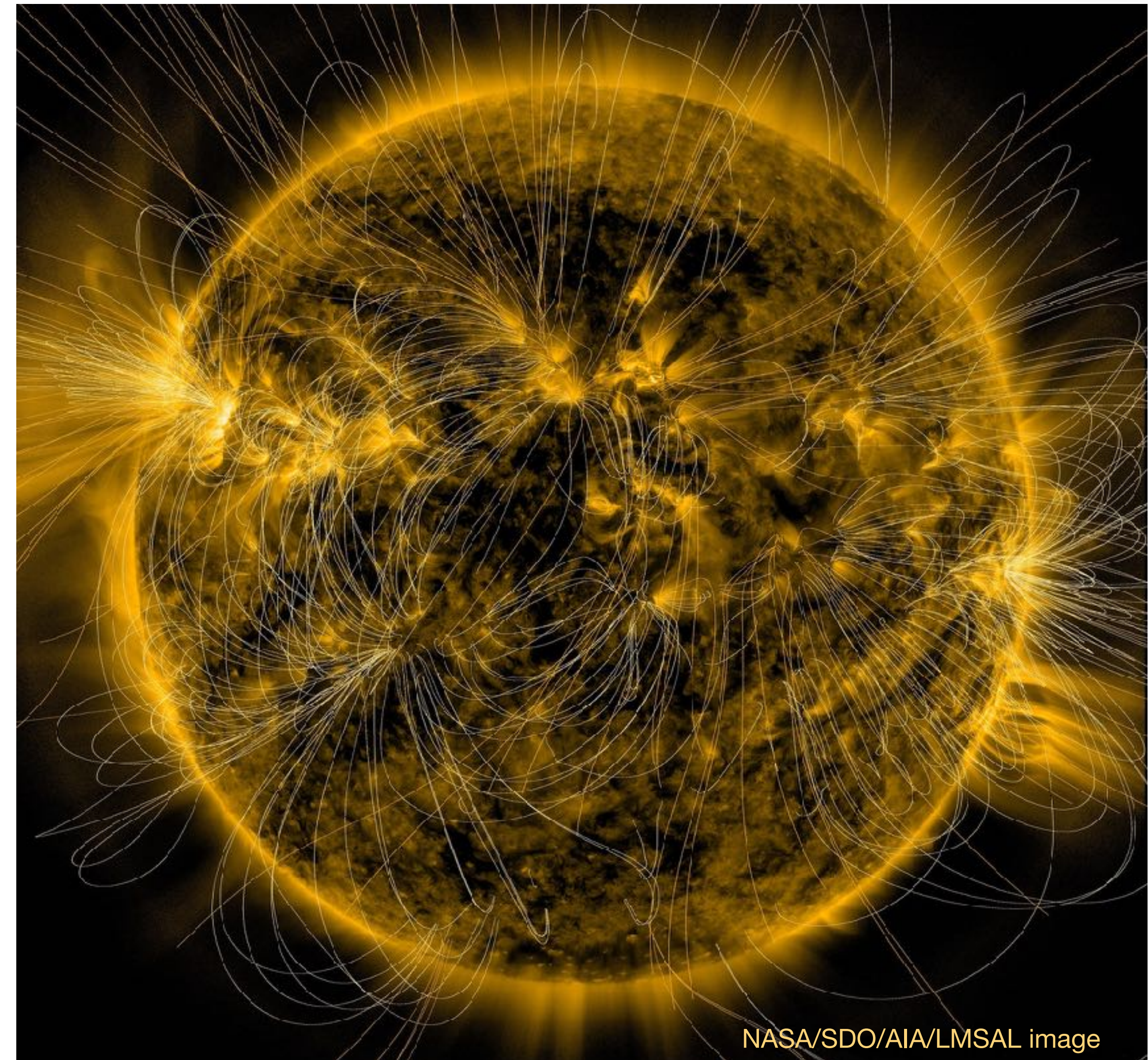
# Another object has $t_*$ smaller than $t_\Omega$ : the Sun

# Another object has $t_*$ smaller than $t_\Omega$ : the Sun

Its **kinetic energy is larger** than its magnetic energy.

**Small-scales dominate** over large-scales.

The Solar dynamo can be modeled with an **IMAC** scenario, in which inertia plays an important role.

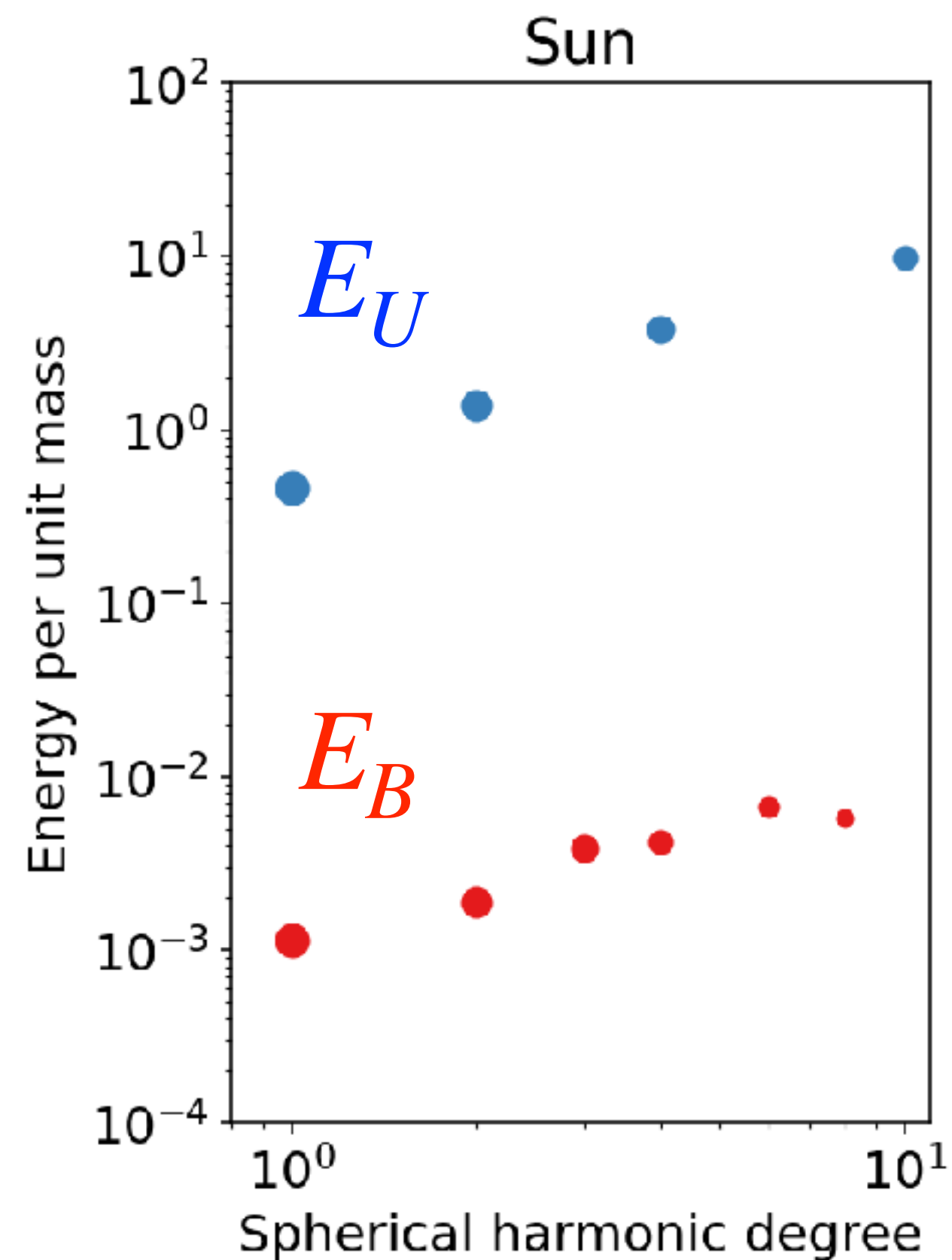


# Another object has $t_*$ smaller than $t_\Omega$ : the Sun

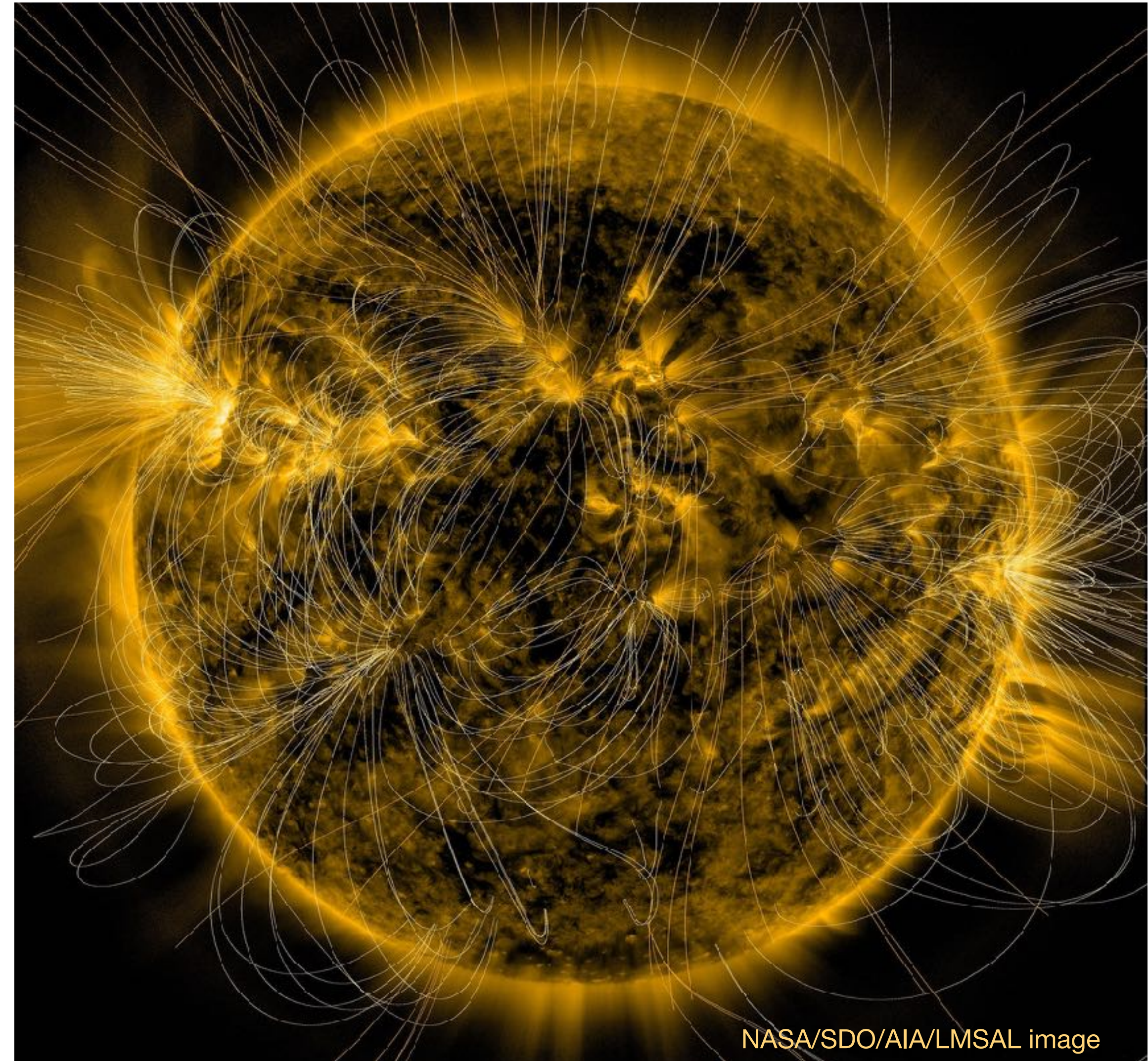
Its **kinetic energy is larger** than its magnetic energy.

**Small-scales dominate** over large-scales.

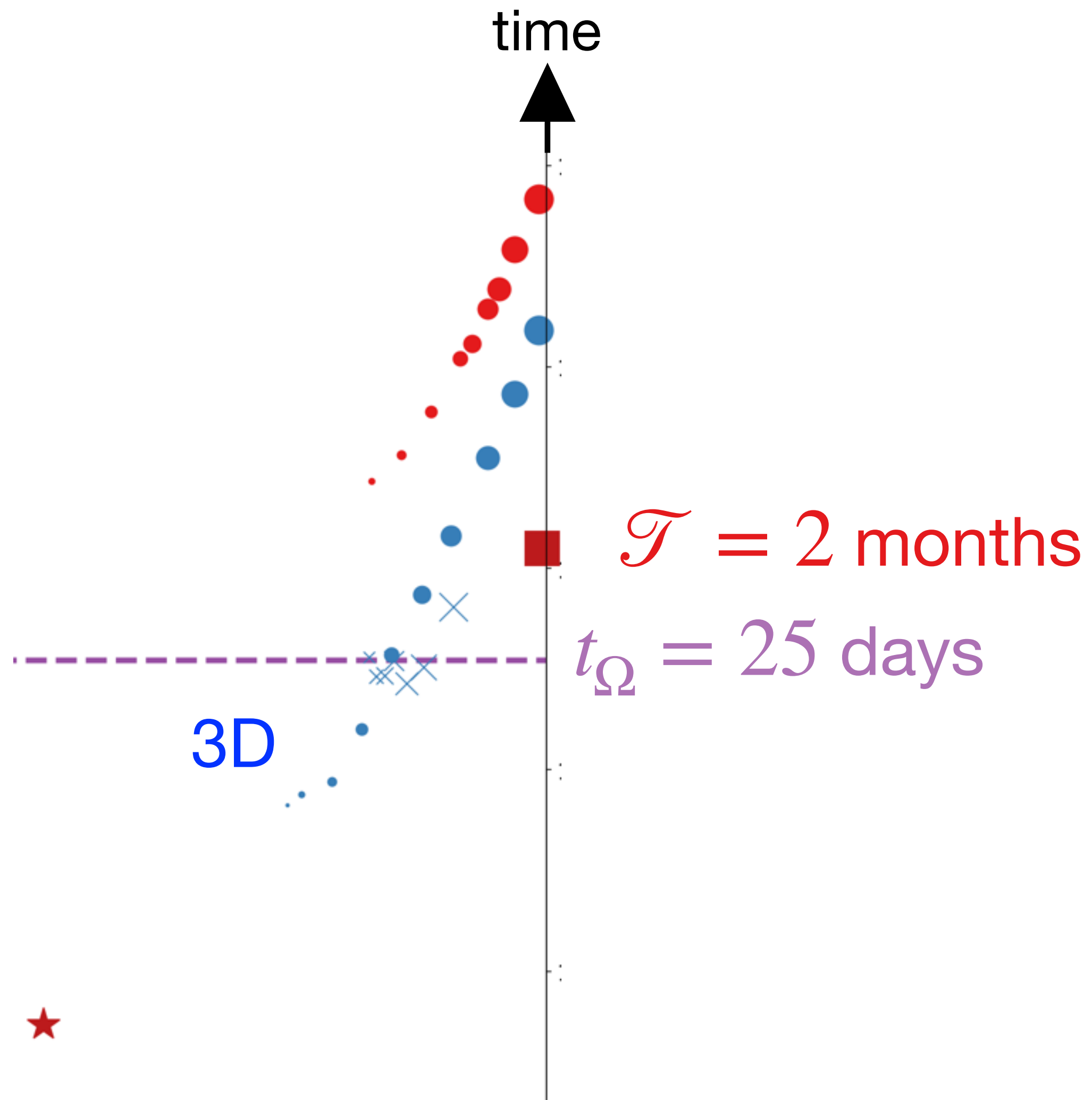
The Solar dynamo can be modeled with an **IMAC** scenario, in which inertia plays an important role.



spectra adapted from  
Proxhauf, 2021  
and Finley & Brun, 2023

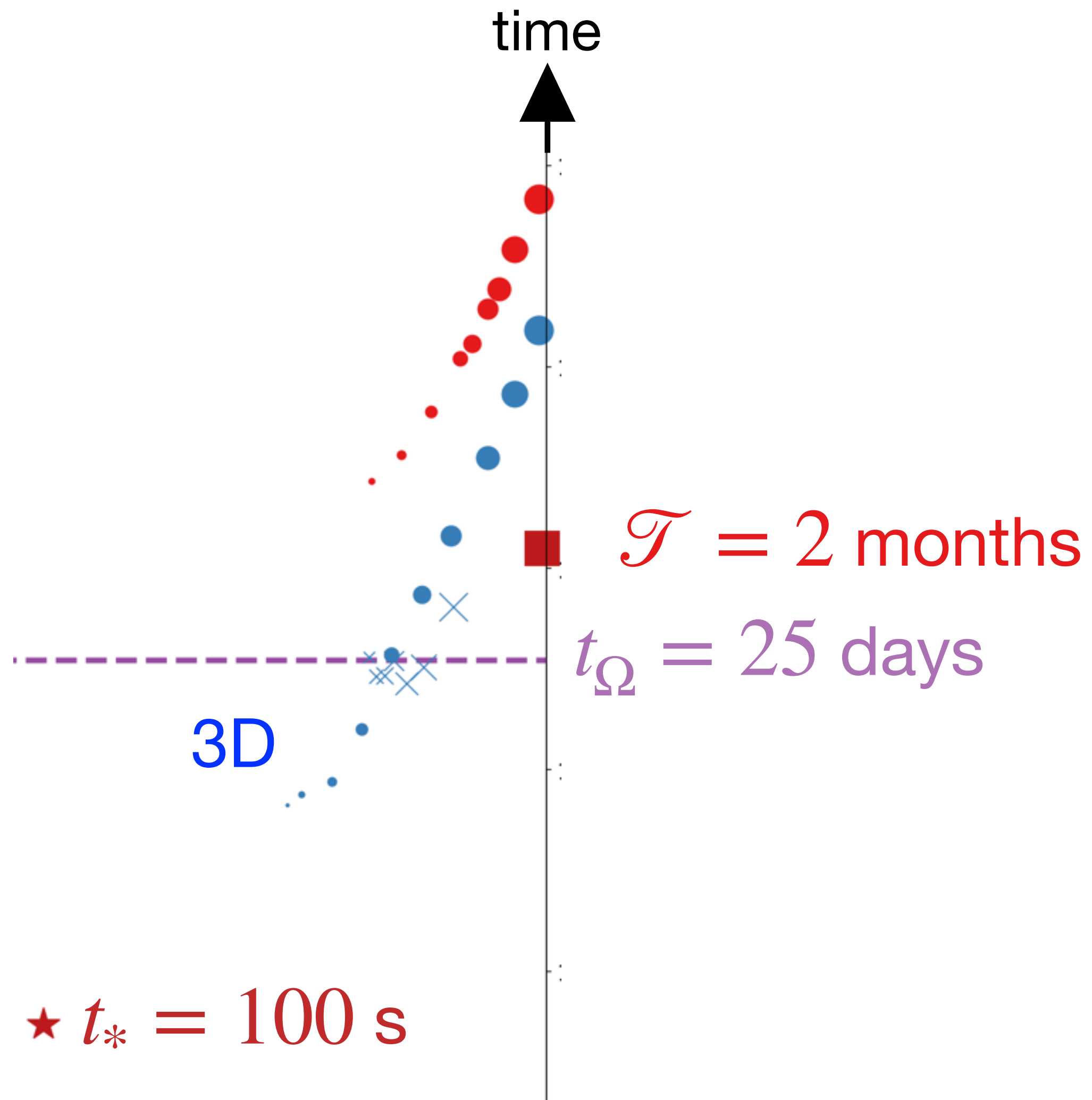


# IMAC characteristic times for the Sun



$\nu = 10^{-3} \text{ m}^2/\text{s}$	kinematic viscosity
$R = 200\,000 \text{ km}$	thickness of conv zone
$M = 5.4 \times 10^{28} \text{ kg}$	mass of convective zone
$\mathcal{P} = 19 \text{ YW}$	convective power 5% $L_{\text{sol}}$
$\eta = 3 \text{ m}^2/\text{s}$	magnetic diffusivity

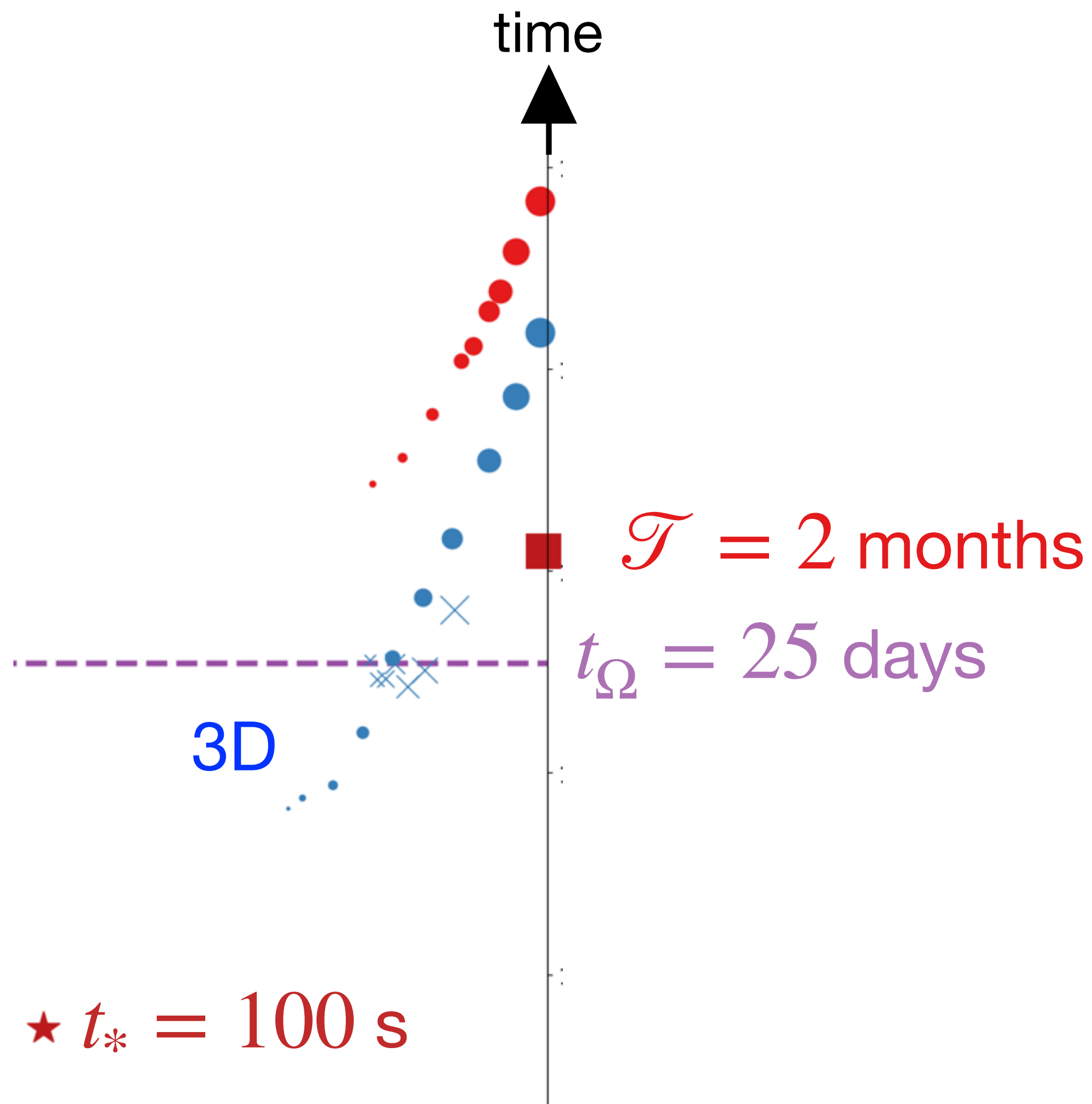
# IMAC characteristic times for the Sun



$\nu = 10^{-3} \text{ m}^2/\text{s}$	kinematic viscosity
$R = 200\,000 \text{ km}$	thickness of conv zone
$M = 5.4 \times 10^{28} \text{ kg}$	mass of convective zone
$\mathcal{P} = 19 \text{ YW}$	convective power 5% $L_{\text{sol}}$
$\eta = 3 \text{ m}^2/\text{s}$	magnetic diffusivity

# IMAC characteristic times for the Sun

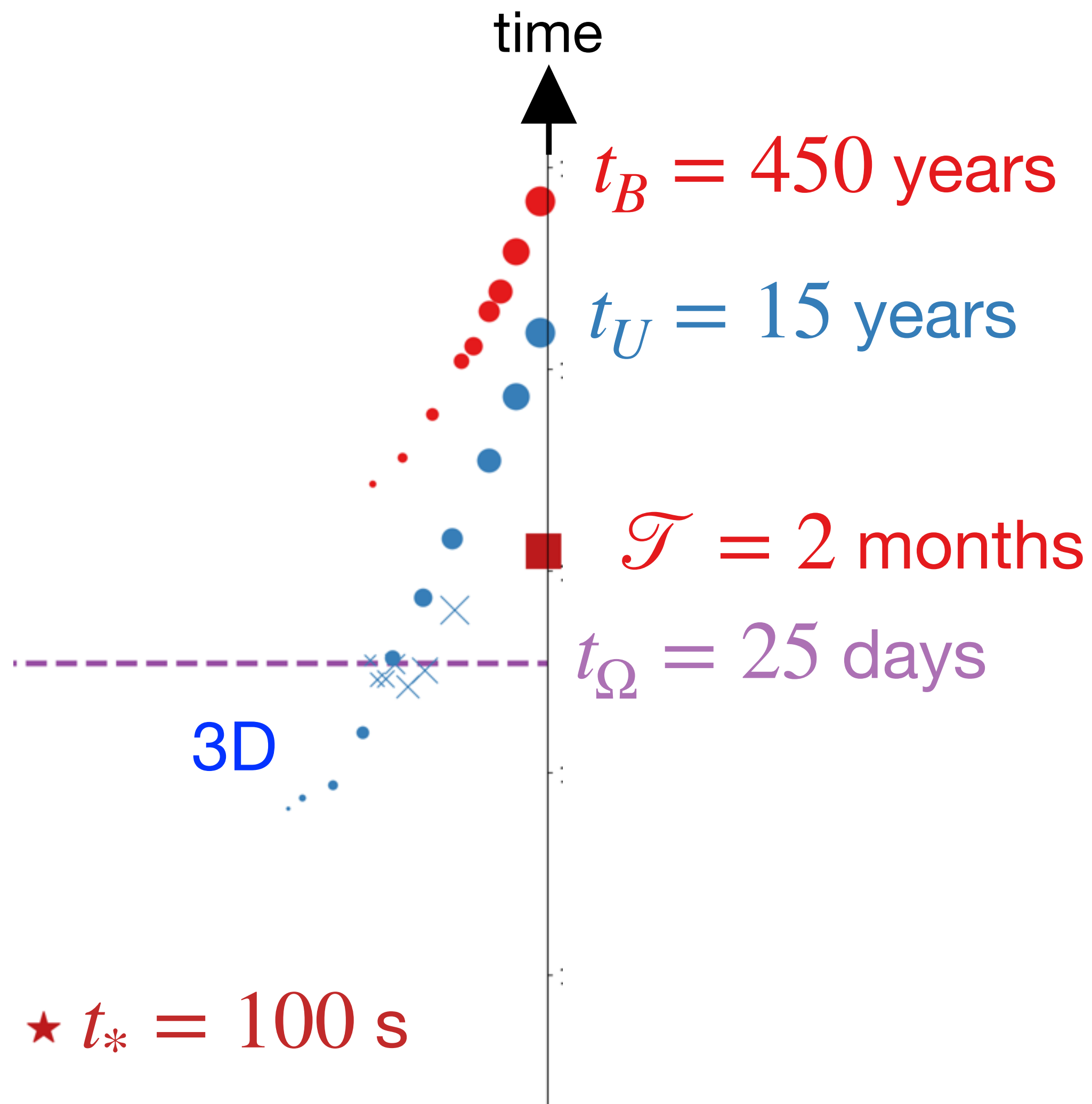
$t_* \ll t_\Omega$  in the convective zone of the Sun.  
 An **IMAC** force balance controls the  
 dynamo and  $t_U \ll t_B$ .



$\nu = 10^{-3}$ m <sup>2</sup> /s	kinematic viscosity
$R = 200\,000$ km	thickness of conv zone
$M = 5.4 \times 10^{28}$ kg	mass of convective zone
$\mathcal{P} = 19$ YW	convective power 5% $L_{\text{sol}}$
$\eta = 3$ m <sup>2</sup> /s	magnetic diffusivity

# IMAC characteristic times for the Sun

$t_* \ll t_\Omega$  in the convective zone of the Sun.  
An **IMAC** force balance controls the dynamo and  $t_U \ll t_B$ .

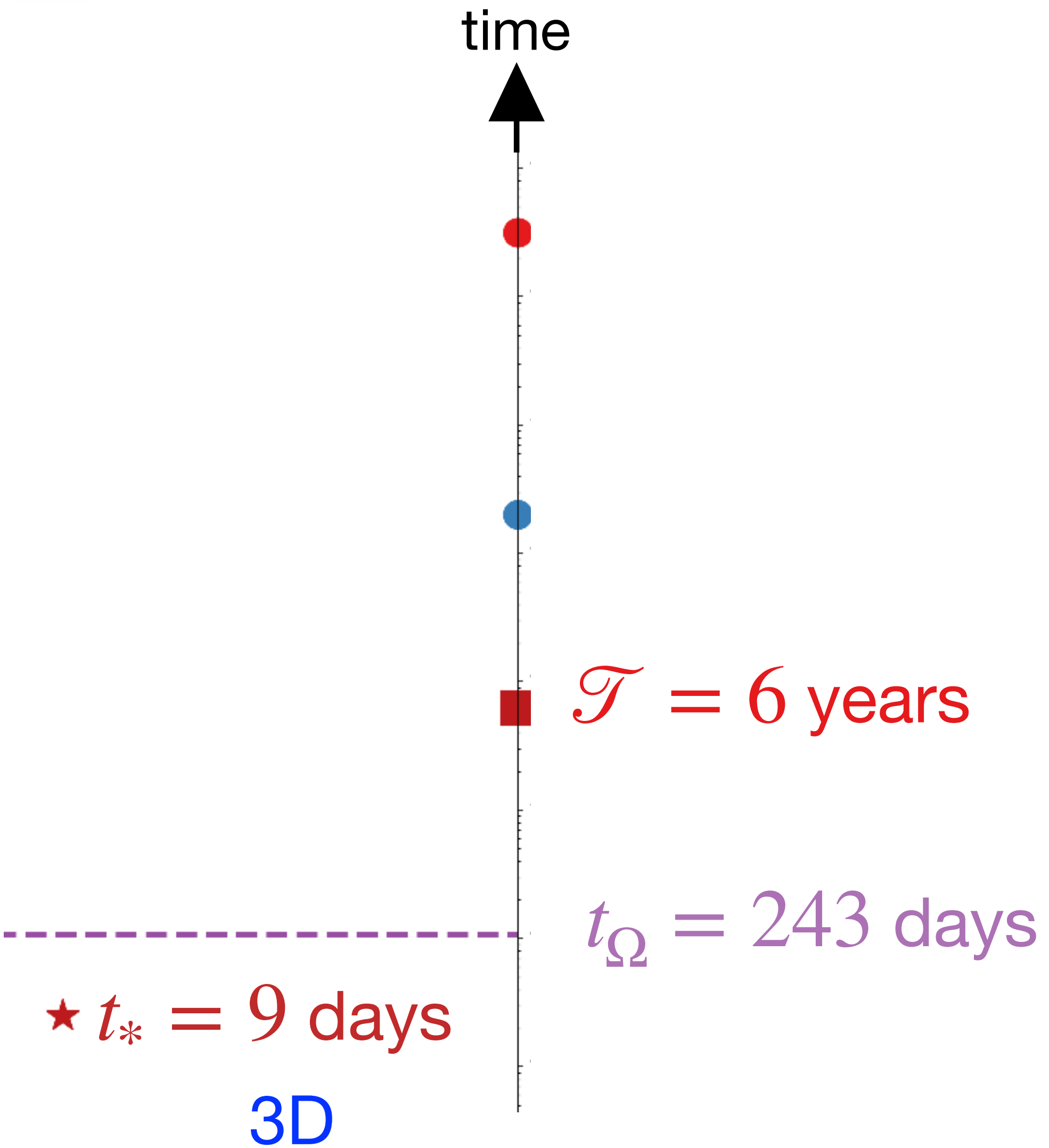


**force balance**

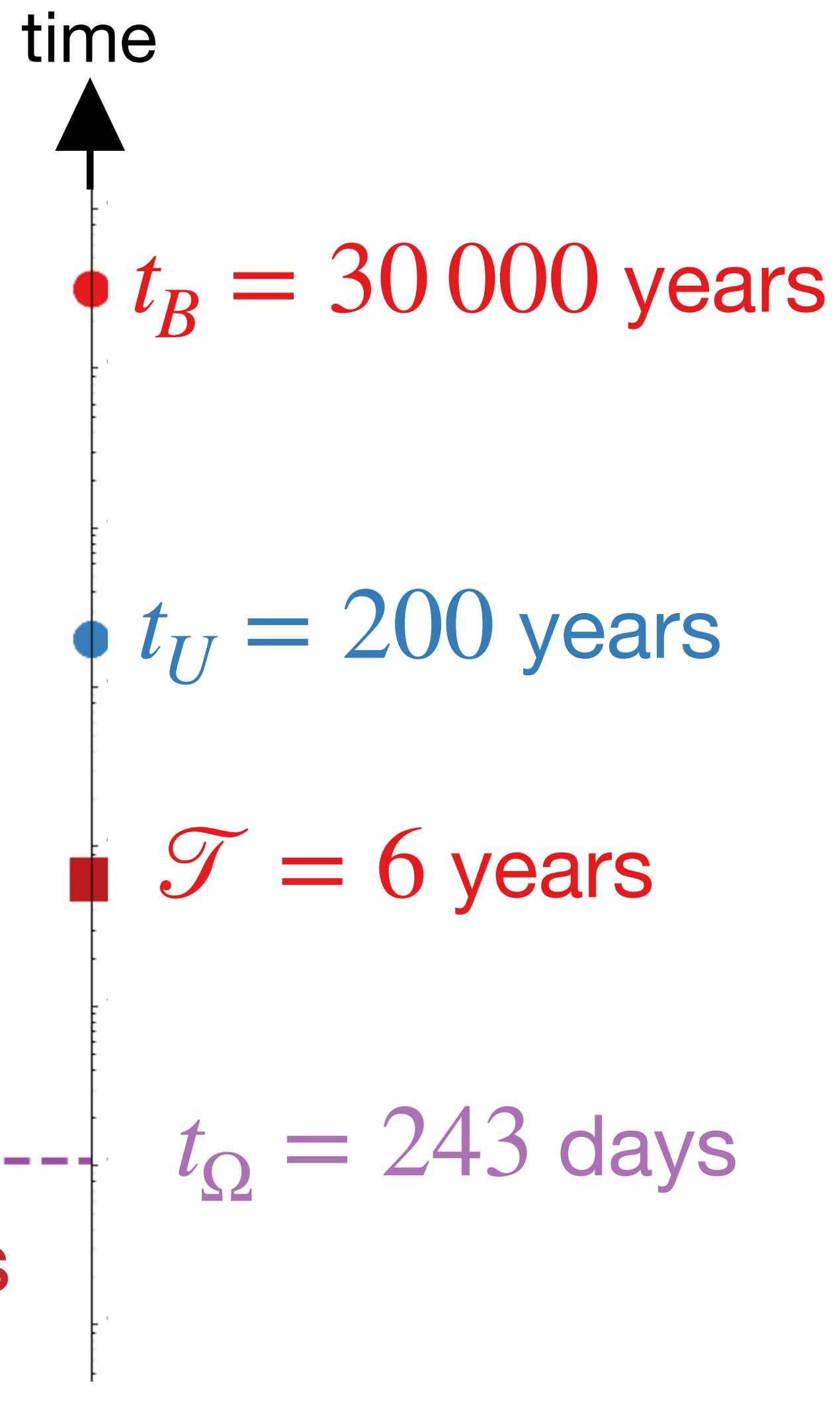
Inertia-Magneto-Archimedean-Coriolis

$\nu = 10^{-3}$ m <sup>2</sup> /s	kinematic viscosity
$R = 200\,000$ km	thickness of conv zone
$M = 5.4 \times 10^{28}$ kg	mass of convective zone
$\mathcal{P} = 19$ YW	convective power 5% $L_{\text{sol}}$
$\eta = 3$ m <sup>2</sup> /s	magnetic diffusivity

# An IMAC force-balance dynamo for Venus



# An IMAC force-balance dynamo for Venus

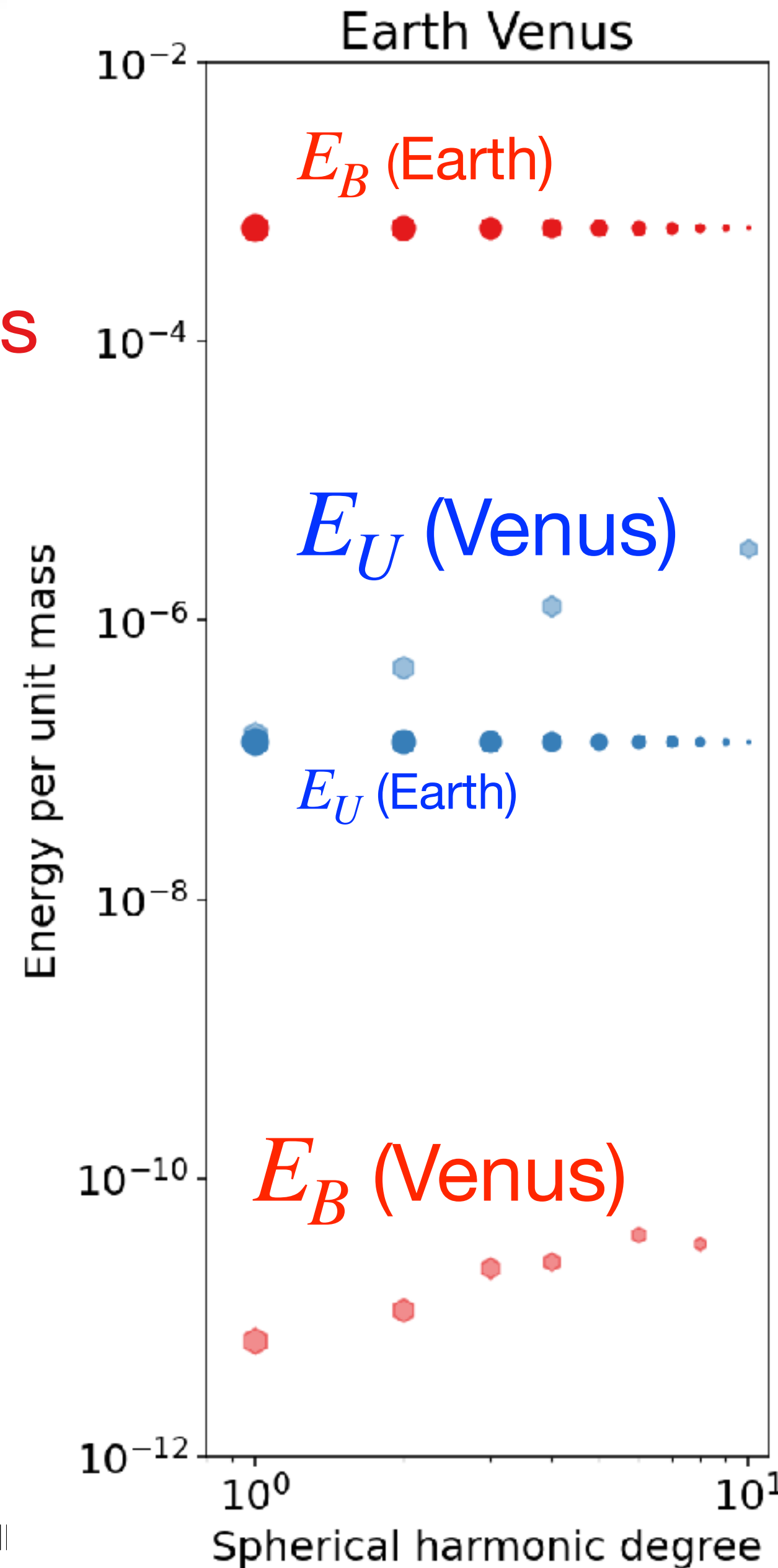
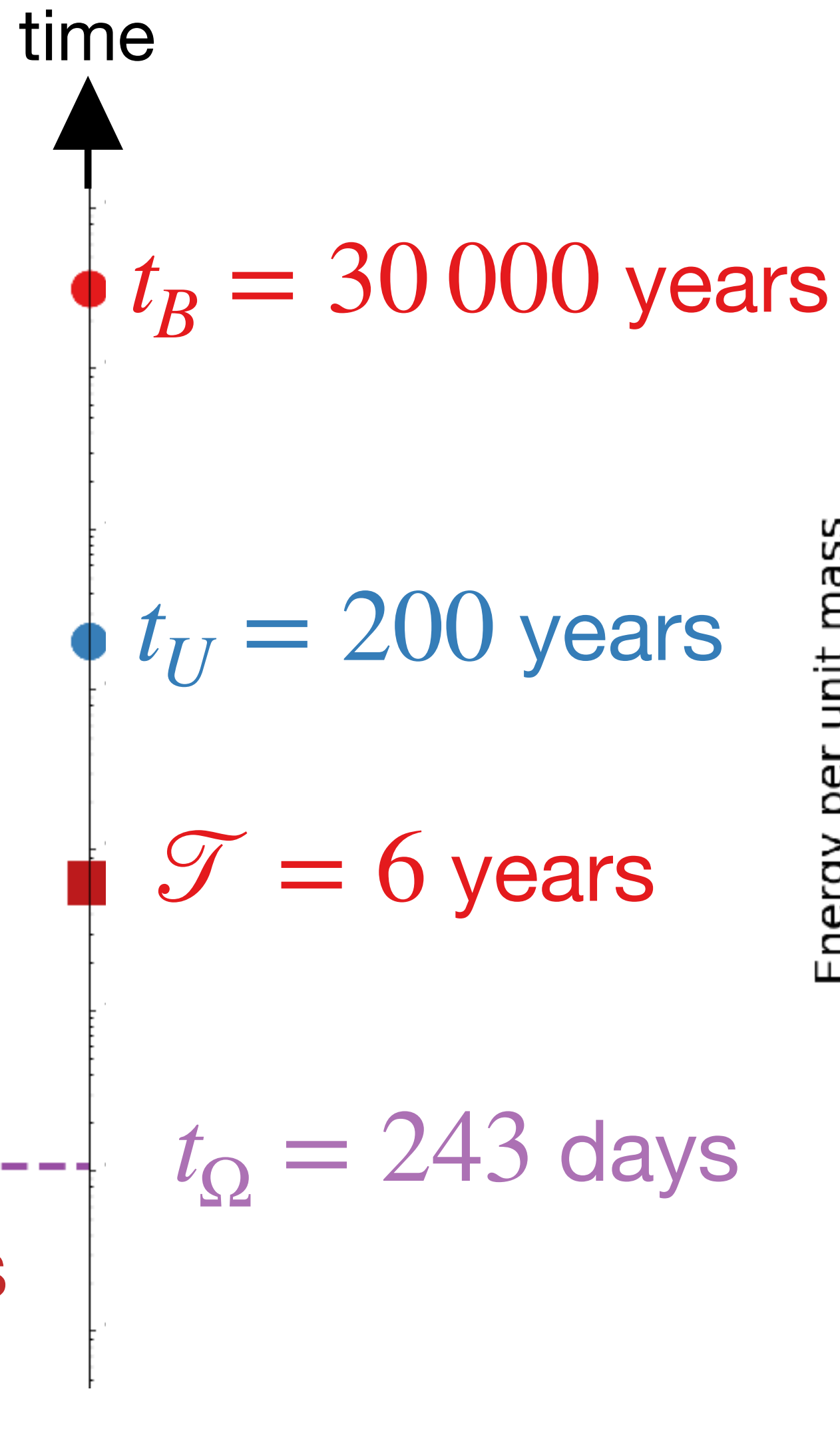


An IMAC force-balance scenario for Venus predicts that its large-scale magnetic field would be **10 000 times smaller** than that of the Earth.

It would be **dominated** by **small-scales**, and might therefore have **escaped detection...**

Zonal flows could be strong.

# An IMAC force-balance dynamo for Venus



An IMAC force-balance scenario for Venus predicts that its large-scale magnetic field would be **10 000 times smaller** than that of the Earth.

It would be **dominated by small-scales**, and might therefore have **escaped detection...**

Zonal flows could be strong.

# Take-home messages

The **classical** view is that both Earth and Venus should belong to the same class of « rapidly rotating dynamos ».

Stevenson, 2003, 2010

# Take-home messages

The **classical** view is that both Earth and Venus should belong to the same class of « rapidly rotating dynamos ».

Stevenson, 2003, 2010

We show that Venus **cannot be** a rapidly rotating dynamo.

# Take-home messages

The **classical** view is that both Earth and Venus should belong to the same class of « rapidly rotating dynamos ».

Stevenson, 2003, 2010

We show that Venus **cannot be** a rapidly rotating dynamo.

The **classical** explanation for the absence of a magnetic field on Venus is simply that its core *would not be convecting* today (or not strongly enough).

Stevenson+, 1983; Nimmo, 2002; Christensen, 2010

# Take-home messages

The **classical** view is that both Earth and Venus should belong to the same class of « rapidly rotating dynamos ».

Stevenson, 2003, 2010

We show that Venus **cannot be** a rapidly rotating dynamo.

The **classical** explanation for the absence of a magnetic field on Venus is simply that its core *would not be convecting* today (or not strongly enough).

Stevenson+, 1983; Nimmo, 2002; Christensen, 2010

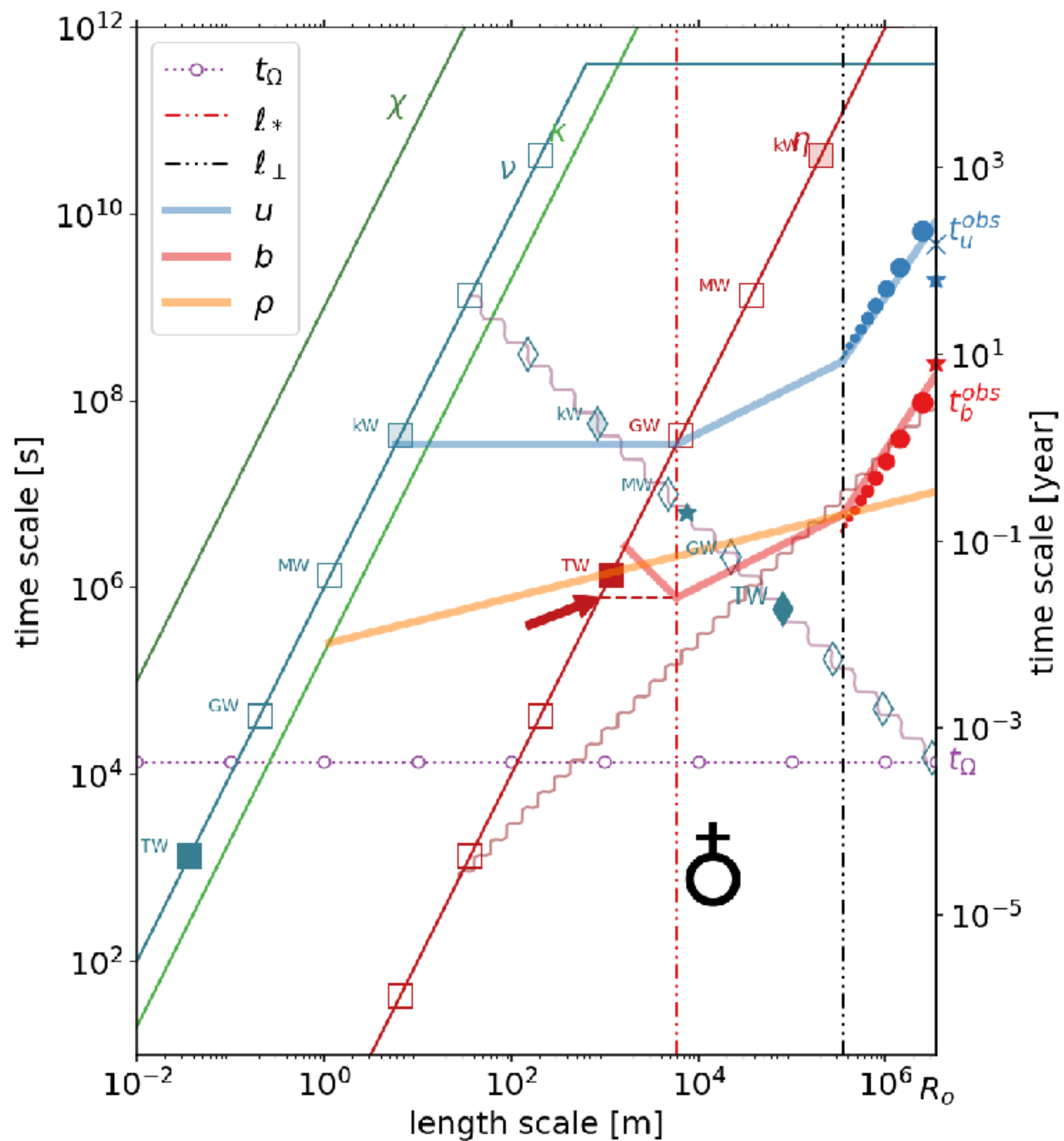
Our analysis suggests an **alternative speculative scenario**: Venus is producing a weak small-scales dominated magnetic field that could have escaped detection...

$t = t_{allocated}$   
Thank you!

- Nataf H-C. & N. Schaeffer, Dynamic regimes in planetary cores:  $\tau - \ell$  diagrams, *Comptes Rendus Géoscience*, **356**, 1-30, [DOI](#), 2024.
- Noraz Q, A.S. Brun & A. Strugarek, Global turbulent solar convection: a numerical path investigating key force balances in the context of the convective conundrum, *The Astrophysical Journal*, **981** 206, [DOI](#), 2025.

# $\tau - \ell$ diagrams

Earth QG-MAC\_JA rotating convective dynamo



Venus QG-MAC\_JA rotating convective dynamo

